

Atmósfera

ISSN: 0187-6236 ISSN: 2395-8812

claudio.amescua@atmosfera.unam.mx

Universidad Nacional Autónoma de México

México

Rostami, Noredin; Fathizad, Hassan **Spatial and temporal changes of land uses and its relationshipwith surface temperature in western Iran** Atmósfera, vol. 35, núm. 4, 2022, Octubre, pp. 701-717

Universidad Nacional Autónoma de México
Ciudad de México, México

DOI: https://doi.org/10.20937/ATM.52985

Disponible en: https://www.redalyc.org/articulo.oa?id=56582396008





Más información del artículo

Página de la revista en redalyc.org



Sistema de Información Científica Redalyc Red de revistas científicas de Acceso Abierto diamante Infraestructura abierta no comercial propiedad de la academia

# Spatial and temporal changes of land uses and its relationship with surface temperature in western Iran

Noredin ROSTAMI<sup>1</sup>\* and Hassan FATHIZAD<sup>2</sup>

Received: October 1, 2020; accepted: March 22, 2021

#### RESUMEN

Se utiliza un algoritmo de ventana dividida en la cuenca de la presa de Ilam para determinar la relación entre la temperatura de la superficie terrestre (LST, por sus siglas en inglés) y los tipos de uso de suelo. Se utilizan imágenes satelitales Landsat del sensor TM para 1990, 1995, 2000, 2005 y 2010 y Landsat 8 (sensor OLI) para 2015 y 2018. Después de las correcciones geométricas y radiométricas de las imágenes satelitales, los mapas de uso del suelo se extraen mediante el método de lógica difusa ARTMAP. Una evaluación de precisión mostró que el valor más alto del coeficiente kappa fue de 94% con una precisión total de 0.95 para 2015, y que su valor más bajo fue de 87% con una precisión total de 0.9 para 1990. Los valores altos de estos coeficientes indican una precisión aceptable en el uso de datos de teledetección Landsat para el uso del suelo. Los cambios más importantes en el uso del suelo están relacionados con bosques densos y bosques dispersos, con una disminución de 20.07 y 17.04%, respectivamente. Los valores mínimos de LST en 1990, 2010 y 2018 en bosques densos son de 21.27, 30.55 y 33.82 °C, respectivamente. Los valores máximos de LST para el uso de tierras forestales dispersas en 1990 y 2010 es 52.48 y 56.09 °C, respectivamente, y de 56.10 °C para el uso de tierras forestales densas en 2018. Como resultado, el LST promedio en tierras agrícolas fue más bajo que en bosques dispersos y pastizales, lo cual se debe principalmente al alto contenido de humedad y a la mayor tasa de evapotranspiración. Las variaciones de uso del suelo/cobertura del suelo (LULC) de 1990 a 2018 muestran que en todos los usos del suelo se ha experimentado un aumento de la LST.

#### ABSTRACT

A split-window algorithm has been used in the Ilam dam watershed to determine the relationship between land surface temperature (LST) and types of land use. Landsat satellite images of the TM sensor for 1990, 1995, 2000, 2005 and 2010 and Landsat 8 (OLI Sensor) for 2015 and 2018 are used. After geometric and radiometric corrections of satellite images, land use maps are extracted by using the fuzzy ARTMAP method. An accuracy assessment showed that the highest value of the kappa coefficient was 94% with a total accuracy of 0.95 for 2015, and the lowest kappa coefficient value was 87% with a total accuracy of 0.9 for 1990. The high values of these coefficients indicate the acceptable accuracy of using Landsat's remote sensing data for land use detection. The most important land use change is related to dense forest and sparse forest land uses, with decreases of 20.07 and 17.04%, respectively. The minimum LST measures in 1990, 2010, and 2018 in dense forest are 21.27, 30.55 and 33.82 °C, respectively. The maximum LSTs for the sparse forest land use in 1990 and 2010 are 52.48 and 56.09, and 56.10 °C for the dense forest land use in 2018. As a result, the average LST in agricultural lands was lower than in sparse forest and rangeland;, which is mainly due to the high moisture content and the greater evapotranspiration rate. Land use/land cover variations from 1990 to 2018 show that all land uses have experienced an increase in LST.

Key words: Landsat satellite, split-window algorithm, fuzzy ARTMAP, kappa coefficient, Ilam dam watershed.

 $<sup>^{</sup>l}D$ epartment of Rangeland and Watershed Management, Faculty of Agriculture, Ilam University, Ilam, Iran.

<sup>&</sup>lt;sup>2</sup>Department of arid and desert regions management, School of Natural Resources and Desert Studies, Yazd University, Yazd, Iran.

<sup>\*</sup>Corresponding author; email: n.rostami@ilam.ac.ir

#### 1. Introduction

Land use and land cover change (LUCC) are among the major driving forces of regional and global climate change (Feddema et al., 2005). However, climatic effects of deforestation and agricultural development vary from one region to another, and studies clearly show that land use changes and land cover play an important biophysical and biochemical role in climate systems at local, regional, and even continental scales (Brovkin et al., 2013; Luyssaert et al., 2014; Mahmood et al., 2014). These also depend on seasonality and land-atmosphere interaction (Halder et al., 2016). Contemporaneous social and economic developments have intensified the impact of human activities on land surface temperature (LST). Urban development, due to the reduction in vegetation cover (Feng et al., 2012) and the increase in impervious materials such as concrete and asphalt (Connors et al., 2013; Hasanlou and Mostofi, 2015; Pal and Ziaul, 2017; Ziaul and Pal, 2018a), increases LST (Amiri et al., 2009). Other factors such as radiation conditions, heat conduction in the upper layer of the Earth's surface, elevation from the surface of the Earth, relief, cloudiness, oceanic flows, and horizontal and vertical air flow are also important to determine LST. Therefore, intensity of land use can reflect the intensity of human activities, providing a basis to assess the relationship between land use change and environmental modifications.

LST is a major parameter in assessing the exchange of surface material, energy balance, and physical and chemical processes, and is now widely used in soil, hydrology, biology, and geochemistry studies (Tomlinson et al., 2011; Hao et al., 2016). Also, LST is an important factor in global studies, and heat stress is considered as a sign of climate change (Srivastava et al., 2009). Because of reflection from surface and roughness of different types of land uses, land use/land cover (LULC) changes are the main causes of LST variations (Hou et al., 2010). The soil LST is sensitive to vegetation cover and soil moisture, so it can be used to track land cover changes and land use, contributing to a better understanding of the desertification phenomena.

Fuzzy ARTMAP is an artificial neural network based on the adaptive resonance theory. The networks which work based on the adaptive resonance theory with supervised learning are known as ARTMAP (Carpenter et al., 1991). Each ARTMAP system consists of two modules (ARTa, ARTb) that create stable recognition classes in response to arbitrary sequences of input patterns. These two modules are connected to each other through an interface module called Fast Appearance-Based Mapping (Fab). The split-window algorithm requires only three parameters (emissivity, atmospheric emission, and average effective air temperature) to calculate LST. This algorithm was developed in 2001 to calculate the surface temperature by using the Landsat 5 Thematic Mapper (TM) sensor. With slight changes in the coefficients of the equations, it was then calibrated for other sensors (Rozenstein et al., 2014). This algorithm is based on mathematical analysis, and calculates LST by using ground data, a thermal infrared sensor (TIRS), land surface emissivity (LSE), and fractional vegetation cover (FVC), obtained from an operational land imager (OLI) (Latif, 2014).

LSTs can be estimated from infrared radiation, which is emitted from the surface, by the Stefan-Boltzmann inverse equation (Reutter et al., 1994). On the other hand, the Normalized Difference Vegetation Index (NDVI) is a good indicator of long-term changes in land cover and its status (Baihua and Isabela, 2015). It should be noted that an increase in temperature might raise the density of vegetation in areas with enough water resources (Xu et al., 2011). It has been demonstrated that there is a logical connection between NDVI and LST (Kustas et al., 2003; Weng et al., 2004; Agam et al., 2007; Inamdar et al., 2008; Wei et al., 2015). While temperature data, recorded by synoptic stations, are not useful for obtaining a good and wide spatial resolution, remote sensing (RS) images are a good source of information for the preparation of thermal maps, which is due to their extensive and continuous coverage, and timeliness (and the ability to obtain information in the reflection and thermal fields of electromagnetic waves) (Jiménez-Muñoz and Sobrino, 2010).

Due to the spatial correlation between data, conventional statistical methods are not suitable to analyze environmental data (Ripley, 1977). In this regard, various studies have been conducted to measure LST using RS technology (Herb et al., 2008; Feizizadeh et al., 2012. Qian et al., 2015; Isaya and Avdan, 2016; Fathizad et al., 2017; Deng et al., 2018;

Li and Jiang, 2018; Ziaul and Pal, 2018 a, b; Tariq and Shu, 2020; Tariq et al., 2020).

Over the past decades, droughts, increased human intervention in soil and natural resources, degradation, rapid population growth, and increasing need for food and new energy resources have caused overexploitation of natural resources in Iran. Undoubtedly, the land use in the Ilam dam watershed has undergone some changes and can be a threat for the residents of the region. So, this study investigates the relationship between land use changes and LST over 28 years (1990-2018) using Landsat satellite imagery, and identifies land uses with the highest surface temperature in the study region. Due to the limited access to ground-based data, the satellite imagery has been used to study the temperature patterns.

#### 2. Materials and methods

#### 2.1. Materials

# 2.1.1 Study area

This research was conducted in the Ilam dam watershed in Iran (Fig. 1), with a total area of 47 652 ha, located at 46° 16′ 50"-46° 38′ 56" E, 33° 23′ 24"-33° 38′ 58" N, and 936-2584 masl. Figure 1 shows the location of the study area.

The study area forms a part of the folded Zagros zone. Zagros Mountains are a complex chain of ridges and mountains in SW Iran, extending NW-SE from the border areas of eastern Turkey and northern Iraq to the Strait of Hormuz. The Zagros range is about 1600 and 240 km long and wide, respectively. This mountain forms the extreme western boundary of the Iranian plateau and its foothills extend into adjacent countries. It divides the region between Iran's dry inland plateau to the east and the fertile plains of Mesopotamia and the Persian Gulf lowlands to the west.

The main kind of vegetation in the Zagros forest habitat is oak trees. At altitudes above the forest border (about 2300 masl) there are dense grasslands and shrubs. In the west of Iran, especially in the Zagros region, oak is the most important and abundant tree species. The Zagros Mountains are the largest and most important habitat of various oak species in Iran, and therefore this region is of special importance.

# 2.1.2 Remote sensing data

In this research, Landsat Thematic Mapper (TM) imagery for 1990, 1995, 2000, 2005 and 2010 (pass/row: 167/37) with six spectral lines with a resolution of 30 m, and Landsat 8 (Operational Land Imager [OLI] Sensor) for 2015 and 2018 (pass/row: 167/37)

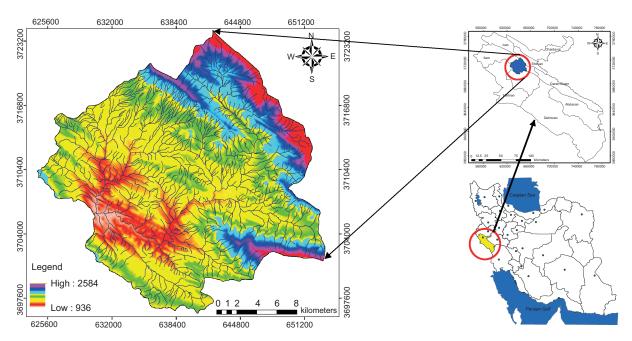


Fig. 1. Location of the study area in Iran and Ilam province.

bands 1 to 7 with a resolution of 30 m, band 8 with a resolution of 15 m and bands 9, 10 and 11 with a 30 m resolution were (Table I). Landsat images, published by the United States Geological Survey (USGS), were retrieved from the Earth Explorer website (http://earthexplorer.usgs.gov).

# 2.1.3 Software packages

ENVI 4.8, ArcGIS10.3 and Excel have been used in this study. The pre-processing works, including geometric and radiometric corrections, and post-processing tasks such as land use classification and LST layers have been extracted with the ENVI 4.8 software. ArcGIS 10.3 is used to provide the output of the maps and Excel is used to perform the statistical analysis.

# 2.2 Methodology

# 2.2.1 Pre-processing of images

The initial raw images of satellite data have incorrect geometry for reasons such as Earth orbits and changes in satellite elevation, in which case these images are not matched with the other satellite data. So, 35 ground control points from topographic maps were gathered for processing and interpreting multi-temporal satellite data. Further, the geometry of images was corrected in the ENVI 4.8 software environment by using the Global Positioning System (GPS).

Radiometric correction is necessary in remote sensing. Removing the undesirable effects of atmosphere is more important when the goal is comparing multi-temporal images (Chavez, 1988). In the present study, the Chavez method, which diminishes the value of dark object subtraction, is used for radiometric correction, and the value of dark object subtraction in the image is reduced to make the classification process highly accurate. For geometrical correction, topographic maps with a scale of 1.55000, prepared by the national geographical organization of Iran, were used. Images used in this research have been corrected by using ground control points and re-sampling equations. So, 45 ground control points with suitable distribution were used at the intersection of roads, waterways, etc., leading to a more accurate mathematical model. The average error obtained for the image sensor used was equal to 0.54 pixels, which is acceptable.

There are two types of radiometric corrections, absolute radiometric correction and relative radiometric correction. The absolute radiometric correction method requires data of atmospheric properties and sensor calibration. This type of correction is often very difficult, especially for older data (Du et al., 2002). In contrast, relative radiometric corrections are done with the aim of reducing the expected atmospheric variables, etc., between multi-time images. The *dark-object subtraction technique* is a simple relative radiometric correction method widely used in many cases (Chavez and MacKinnon, 1994). For radiometric correction, digital values are converted into spectral radiance in the first step by using the calibration coefficients of the sensor and the following equation:

Table I. Applied Landsat satellite imagery.

| Sensor | Data  | Pass/row                             | Resolution (m)       | Number of bands  |
|--------|---|--------------------------------------|----------------------|------------------|
| MSS    | 1985. 07. 22  | 167/37                               | 60                   | 4                |
| TM     | 1990. 08. 29<br>1995. 08. 19<br>2000. 08. 16<br>2005. 10. 1 | 167/37<br>167/37<br>167/37<br>167/37 | 30<br>30<br>30<br>30 | 7<br>7<br>7<br>7 |
|        | 2010. 07. 11  | 167/37                               | 30                   | 7                |
| OLI    | 2015. 08. 26<br>2018. 08. 18                                | 167/37<br>167/37                     | 30<br>30             | 11<br>11         |

MSS: Multispectral Scanner System; TM: Thematic Mapper; OLI: Operational Land Imager.

$$L = gain \times DN + offset \tag{1}$$

where L is the spectral radiance (Wem<sup>-2</sup> Ster<sup>-1</sup>  $\mu$ m<sup>-1</sup>), DN the digital value of the pixel (0 to 255), and *gain* and *offset* are the calibration coefficients of the sensor. In the next step, according to Eq. (2), the amount of spectral radiance is converted into spectral reflectance (Richards, 2013; Lillesand et al., 2015):

$$p = \frac{\pi L d^2}{ESUN.COS(SZ)}$$
 (2)

where p is the spectral reflectance without units, from 0 to 1; L is the spectral radiance in the sensor;  $d^2$  is the square of the distance between the Earth and the Sun based on astronomical units; ESUN is the height of the Sun, and SZ is the angle of the Sun when radiating while recording a satellite image.

By converting spectral radiance values to spectral reflectance, the effects of changes in the Sun light, season, latitude, and weather conditions on the images are eliminated. The outcome is relatively standard and can be used to directly compare the reflections of phenomena between different images and an image at different times. In this study, the method of reducing the darkness of the phenomenon, which is implemented in ENVI software, has been used for radiometric correction. This process is to reduce the effects of atmospheric diffusion on the image.

#### 2.2.2 Post-processing of images

After geometric and radiometric corrections of the satellite image and cutting it in the study area, the land use map was extracted by using the fuzzy ART-MAP supervised classification method to six land use types of fair rangeland (26-50% of climax condition), poor rangeland (0-25% of climax condition), dense forest (canopy cover less than 25%), sparse forest (canopy cover more than 25%), agricultural land, and a water body. The main advantage of the fuzzy ARTMAP is that it requires lower training data for accuracy analysis compared with other methods. In addition, this method does not depend on the statistical distribution of data and does not require specific statistical variables. In order to verify the accuracy of the classification, a comparison was made with existing land use maps and field observations. In this way, the reference map or reality map from all parts of the study area was prepared by using other

methods. A random sampling method was used to assess the accuracy of the obtained maps. Samples were recorded by using a GPS method in a number of polygons, by using a land use map and local views from the study area. In order to classify and separate the land uses of previous years from each other, the vegetation map of Ilam province (developed by the Forest, Range and Watersheds Management Organization of Iran [FRWO]) was used together with aerial photographs (1:20000) prepared by the National Cartographic Center of Iran (NCC).

Also, a split-window algorithm, which removes atmospheric effects, was developed and applied to estimate the emissivity and LST. In fact, LST is obtained by using the corrected thermal radiance. To calculate the corrected thermal radiance, it is necessary to determine the emissivity in the thermal bond.

# 2.2.3 Fuzzy ARTMAP classification method

The fuzzy ARTMAP method is a remote sensing classification based on neural network analysis and the Adaptive Resonance Theory (ART). The fuzzy ARTMAP supervised classification method consists of four layers: input layer  $(F_1)$ , category layer  $(F_2)$ , field layer, and output layer. The input layer represents the imported images, so there are some neurons to measure each criterion. The input layer for the infinity of the criterion is as follows:

$$I = (a, a^c) = (a_1, a_2...a_n, 1 - a_1, 1 - a_2...1 - a_n)$$
 (3)

In this method, the number of  $F_2$  layer neurons is automatically determined. The field layer and output layer are constructed with the ARTb model. Each of these two layers has m neurons. There is a one-to-one connection between these two layers.

2.2.4. Split-window algorithm for LST calculation With this algorithm, LST is obtained by using corrected thermal radiance (Waters et al., 2002). To calculate the corrected thermal radiance, it is necessary to calculate the emissivity in the thermal bond. In fact, every land-related phenomenon is characterized by a specific emissivity, which was indicated by Snyder (1998). After detecting the minimum and maximum values, the NDVI is obtained; therefore, the average of soil and vegetation emissivity and the distribution of other areas can be calculated from Eq. (4).

$$\mathcal{E} = \mathcal{E}V \times PV + \mathcal{E}S \times (1 - PV) + d\mathcal{E} \tag{4}$$

Where  $\varepsilon v$  and  $\varepsilon s$  are the emissivity of areas with complete vegetation and areas with dry soil, respectively, and  $d\varepsilon$  is the effect of surface distribution, which is calculated by using Eq. (5).

$$d\varepsilon = (1 - \varepsilon s)(1 - F)\varepsilon v \tag{5}$$

Where F is the form factor and its mean value is equal to 0.55, and Pv is the vegetation percentage, which is calculated by Eq. (6).

$$PV = \left(\frac{NDVI_{\text{max}} - NDVI}{NDVI_{\text{max}} - NDVI_{\text{min}}}\right)^{a}$$
 (6)

#### 2.2.5. NDVI index

This index is one of the most practical indicators for vegetation studies. It has a simple computational process and has the best dynamic power compared to other indicators. This index is more sensitive to vegetation changes and less susceptible to atmospheric and soil effects, except in cases where vegetation is low. The NDVI index is derived from Eq. (7).

$$NDVI = \frac{NIR - RED}{NIR + RED} \tag{7}$$

where NIR is the reflection in the near infrared band and RED is the reflection in the red band. From a theoretical point of view, the value of this index is in the range of  $\pm 1$ . The values of this indicator for dense vegetation cover tend to be  $\pm 1$ . Clouds, snow and water are identified with negative values. Rocks and bare lands, with similar spectral responses in two bands, are seen at values close to zero. In this index, the common soil is equal to 1. Density of vegetation is determined by the distance between the indices of one pixel higher than the soil value (Allison, 1989).

# 2.2.6. Correction of LST

In this method, LST is obtained in °K by using Eq. (8) (Artis and Carnahan, 1982).

$$LST = \frac{T_B}{1 + \left(\lambda \times \frac{T_B}{\rho}\right) Ln\varepsilon}$$
 (8)

Where LST is land surface temperature in  ${}^{\rm o}{\rm K}$ ,  $\lambda$  is the wavelength of the desired band (11.5  $\mu{\rm m}$ ) (Weng et al., 2004),  $\rho$  is the Boltzmann constant (Eq. 9), h is the Planck's constant (6.626  $\times$  10<sup>-34</sup>), c is the speed of light (2.998  $\times$  10<sup>8</sup> m s<sup>-1</sup>), and  $T_{\rm B}$  is the brightness temperature in  ${}^{\rm o}{\rm K}$  obtained from Eq. (10).

$$\rho = \frac{h \times c}{1.438 \times 10^{-2} mk} \tag{9}$$

$$T_B = \frac{K_2}{\ln\left(\frac{K_1}{L\lambda} + 1\right)} \tag{10}$$

K<sub>1</sub> and K<sub>2</sub> are correction coefficients with values of 666.09 and 1287.71, respectively (for Landsat images).

In ENVI, after obtaining the temperature of the black body and multiplying it by the coefficients of any phenomenon (emissivity), temperature can be calculated in °C or °K. It is also possible to calculate LST by using ENVI. Finally, the statistical parameters of each land use and LST for the years 1985, 1990, 1995, 2000, 2005, 2010, 2015 and 2018 are plotted on ArcGIS.

The final part of this procedure is the accuracy evaluation of the results. So, by using the ground control points, the correctness of calculations is evaluated by using the error matrix and statistical parameters of the total accuracy, kappa coefficient, and user's and producer's accuracy (Lu et al., 2004).

#### 2.2.7. Accuracy assessment

Estimating accuracy is very important to understand the results and make decisions. The most common parameters for estimating accuracy are the overall accuracy, producer's accuracy, user's accuracy, and kappa coefficient (Lu et al., 2004). From a probability theory point of view, overall accuracy (Eq. 11) cannot be a good criterion for evaluating classification results; randomness may play a significant role in this index.

$$OA = \frac{1}{N} \sum P_{ii} \tag{11}$$

where OA is the overall accuracy, n is the number of experimental pixels, and  $P_{ii}$  are the elements of the original diameter of the error matrix.

Due to the problems of overall accuracy, the kappa coefficient is often used in executive tasks where the comparison of classification accuracy is considered (Mesgari, 2002).

$$Kappa = \frac{p_{o-}p_c}{1 - p_c} \times 100 \tag{12}$$

where  $P_{\rm o}$  stands for correctly observed and  $P_{\rm c}$  for expected agreement.

The producer's accuracy is the probability that a pixel in the classification image will be placed on the ground in the same class, and the user's accuracy is the probability that a specific class will be placed on the ground in the same class on the classified image (Eqs. 13 and 14)

$$UA = \frac{ta}{n_1} \times 100 \tag{13}$$

$$PA = \frac{ta}{ga} \times 100 \tag{14}$$

where PA is the percentage accuracy of a class for producer's accuracy, ta is the number of correct pixels classified as class a, ga is the number of class a pixels in ground reality, UA is the percentage accuracy of a

class for user's accuracy, and  $n_1$  is the number of pixels in class a as a result of classification.

By cutting the classified maps with the ground reality map obtained from the field survey, an error matrix was formed to evaluate the accuracy of the classified maps, and based on that, the overall accuracy and the kappa coefficient were calculated.

#### 3. Results

#### 3.1. Land use map

In this research, a neural network fuzzy ARTMAP algorithm was used to determine and plot land use maps for 1990, 1995, 2000, 2005, 2010, 2015 and 2018 by using Landsat satellite images. Classes of Landsat images in the study area included rangeland, poor rangeland, dense forest, sparse forest, agricultural, and water body land use types. The land use map of the study area for 1990, 1995, 2000, 2005, 2010, 2015 and 2018 is shown in Figure 2.

The results of the accuracy evaluation of the classified images are given in Table II, in which total, producer's and user's accuracy, as well as

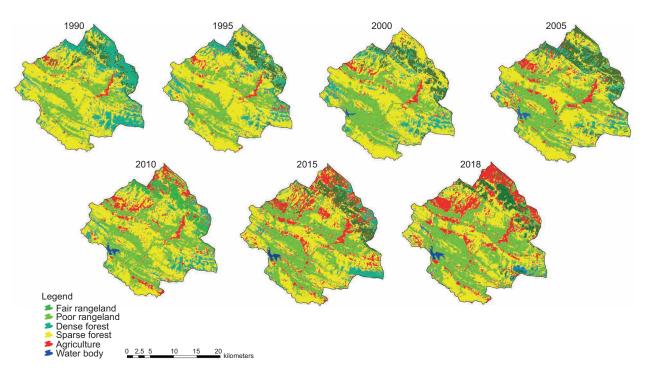


Fig. 2. Land use map of the study area based on the neural network fuzzy ARTMAP classification method for the years 1990, 1995, 2000, 2005, 2010, 2015 and 2018.

| Table II: Results of accuracy e | valuation of the classified l | land use images for the period 1985 |
|---------------------------------|-------------------------------|-------------------------------------|
| to 2018.                        |                               |                                     |

| Vacadalasa         | User's accuracy |      |      |      |      |      |      |  |
|--------------------|-----------------|------|------|------|------|------|------|--|
| Year/class         | 1990            | 1995 | 2000 | 2005 | 2010 | 2015 | 2018 |  |
| Fair rangeland     | 0.67            | 0.73 | 0.78 | 0.75 | 0.74 | 0.81 | 0.85 |  |
| Poor rangeland     | 0.95            | 0.95 | 0.95 | 0.94 | 0.94 | 0.99 | 0.97 |  |
| Dense forest       | 0.8             | 0.94 | 0.96 | 0.92 | 0.88 | 0.95 | 1    |  |
| Sparse forest      | 0.84            | 0.91 | 0.91 | 0.88 | 0.86 | 0.94 | 0.83 |  |
| Agricultural       | 0.94            | 0.95 | 0.88 | 0.91 | 0.90 | 0.94 | 0.93 |  |
| Water body         |                 |      | 0.99 | 0.99 | 0.99 | 1    | 1    |  |
| Total accuracy (%) | 0.90            | 0.94 | 0.92 | 0.92 | 0.91 | 0.95 | 0.93 |  |

| Year/class            | Producer's accuracy |      |      |      |      |      |      |  |
|-----------------------|---------------------|------|------|------|------|------|------|--|
| real/class            | 1990                | 1995 | 2000 | 2005 | 2010 | 2015 | 2018 |  |
| Fair rangeland        | 0.81                | 0.99 | 0.93 | 0.96 | 0.94 | 0.98 | 0.99 |  |
| Poor rangeland        | 0.96                | 0.93 | 0.96 | 0.94 | 0.94 | 0.97 | 0.94 |  |
| Dense forest          | 0.85                | 0.88 | 0.88 | 0.85 | 0.84 | 0.93 | 0.69 |  |
| Sparse forest         | 0.80                | 0.92 | 0.86 | 0.89 | 0.85 | 0.90 | 0.96 |  |
| Agricultural          | 0.89                | 0.91 | 0.88 | 0.89 | 0.88 | 0.95 | 0.94 |  |
| Water body            |                     |      | 0.92 | 0.96 | 0.95 | 0.98 | 0.93 |  |
| Kappa coefficient (%) | 0.87                | 0.92 | 0.90 | 0.90 | 0.89 | 0.94 | 0.91 |  |

kappa coefficient, are also reported. High values of these coefficients indicate the acceptable accuracy for the land use identification by using RS data of Landsat images. It is observed that the classification accuracy for all the mentioned years is above 85%, which is an appropriate precision for classification of land uses

#### 3.2. Land use changes

Results about changes in land use types of the study area during the investigated time periods are shown in Figure 3. They show that during the 28-year period (1990-2018), the area of agricultural land use has increased by 8650.6 ha (18.155%), highlighting an rise in population as well as in human pressure in the studied area. The most important land use change is related to dense and sparse forest land uses, with a decrease of 9867.14 and 8125.11 ha (20.07 and 17.04%), respectively. The study area has been conserved by governmental organizations in recent years. However, a decline in the extent of forest land use in the region is still seen. In a study that examined oak decline in the province of Ilam, results indicated a kind of ailment called charcoal

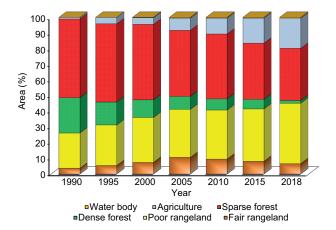


Fig. 3. Area of land use classes (1990-2018).

disease (*Biscogniauxia mediterranea*) and borer beetles, which cause trees to die and fall. This disease has developed in recent years due to climatic conditions including rainfall reduction, drought and moisture stress which provide a suitable field for disease outbreaks (Mirabolfathi, 2013). Indeed, the rising demand for wood, timber, shelter and agricultural products has led to the destruction of natural

land cover, especially forests that are becoming agricultural land at an alarming rate. The pressure from population growth has also led to uncontrolled alterations in exploitation, especially forestry in a wide area of the region. Hence, erosion and destruction of land are two of the most important current problems of this region. Negative effects of land use change in the study area include declining soil fertility, reduced vegetation and animal diversity, flood events, and increased landslide risk. The area of fair and poor rangeland land use types has increased by 1399.77 and 7652.16 ha (2.93 and 16.05%), which is due to the change of forest land to rangeland. Water bodies were not included in land use categories in 1985 and they were added in 2000. This land use is related to the Ilam dam, constructed in 2000 to supply drinking water.

# 3.3. LST maps

According to LST maps (Fig. 4), the central and northwestern parts of the study area have a higher temperature because of agricultural activities, dense forest, and poor rangeland land use. LST maps (Fig. 4) show minimum LSTs of 21.27, 20.88, 25.37, 13.34, 28.85, 26.95 and 30.71 °C and maximum of 54.18, 50.65, 52.85, 46.16, 57.51, 52.04 and 56.10 °C for

1990, 1995, 2000, 2005, 2010, 2015 and 2018, respectively. In this study, the lowest temperature was due to dense forest land use in 1990, when the study area had a cool median temperature. However, overtime increases in the LST indicate climate change effects in the region, which could have devastating outcomes in the future. The maximum temperature drop of 2005 was caused by increased rainfall (above 600 mm) in the region, which led to an increase in vegetation cover in the study area.

# 3.4. Relationship between NDVI, LULC and LST The LST is affected by different surface conditions, such that areas with vegetation accumulation tend to have lower LST than vegetation-free areas. By absorbing sunlight and transpiration of water through its leaves, vegetation creates a natural air-conditioning system. Changes from forest to rangeland and rainfed farming land uses reduce the vegetation cover, remove the cooling system of natural surfaces and increase LST. Figure 5 shows the NDVI and LST differences between 1990 and 2018 in the study area. As it is obvious, wherever the vegetation has

The results of the survey on average LST of the land uses in the study area for different years are

increased, LST has decreased and vice versa.

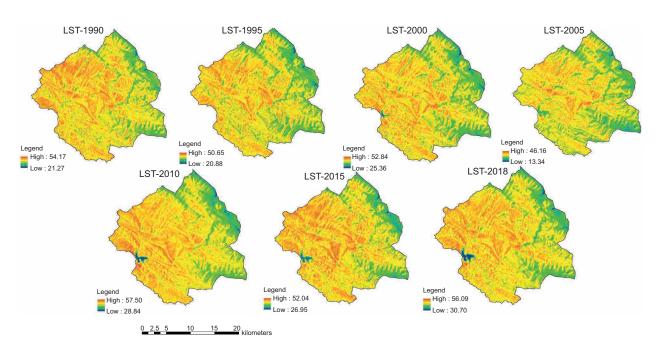


Fig. 4. LST maps obtained with the split-window algorithm (in °C).

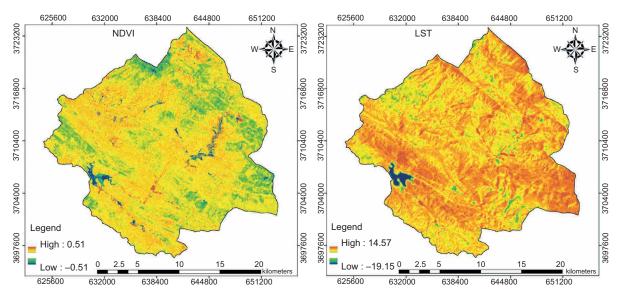


Fig. 5. Role of NDVI and LST (in °C) differences between 1990 and 2018 in the study area.

shown in Table III. According to this table, the minimum temperature in 1990, 2010, and 2018 in dense forest was 3.5, 7.3 and 7.9 °C, respectively, while the maximum surface temperature for sparse forest in 1990 and 2010 was 9.5, 11.8 °C, and 10.9 °C for dense forest in 2018. During most of the studied period, agricultural land had a lower average temperature than sparse forest and rangeland, which is mainly due to the high moisture content in agricultural land and the greater degree of evapotranspiration. Water bodies have the lowest average temperature due to the high-water heat capacity. Changes between 1990 and 2018 show that all land use types were subject to an increase in average temperature, which can be attributed to the increasing trend in temperature in the study area.

In order to study the effect of land use changes on surface temperature, an LST map for different time periods between 1990 and 2018 was prepared and compared to the land use change map for the same time period (Fig. 6). The results show that there was an increase in average LST in areas where land use change increased, indicating an intensification in heat-producing human-based activities, such as conversion of forest land use to agricultural use.

# 3.5. The nature of LST changes and land use The crosstab method was used to investigate the land use and LST changes. In this method, classes of two

classified maps are compared one by one. As a result, it is possible to determine the changes occurring in each class relative to the other. Figure 9 shows the crossed classified land use maps for 1990 and 2018, which exhibit the 1990 LULC variations compared to 2018. The difference in LST maps between 1990 and 2018 was used to examine LST changes (Fig. 5). The amount of various land use changes and the LST changes are shown in Figure 7 and Table IV. As Table IV shows, average temperature for all land use changes was positive and increasing, with the exception of dense forest, sparse forest and agricultural land, where LST dropped to -12.4 °C.

3.6. Relationship between LST and NDVI index
For better analysis of the relationship between LST and Land use, the correlation coefficients between LST and the NDVI index were calculated in 2018, based on randomized control points from each land use (Fig. 8). The highest correlation coefficient was obtained in dense forest and sparse forest land uses at 0.833 and 0.814, respectively. The lowest correlation coefficient (0.288) was related to water body land use, which confirms the lack of vegetation cover in this site compared to other land uses. Comparing the LST values of 2018 (Table III) and their correlation with the NDVI index (Fig. 8), it can be stated that in each land use where the average temperature is higher, the dependence between LST

Table III: Temperature variations in different land uses (1990-2018).

| Class          | 1990  |       |       |      | 1995  |       |       |      |
|----------------|-------|-------|-------|------|-------|-------|-------|------|
|                | Min.  | Max.  | Mean  | Std. | Min.  | Max.  | Mean  | Std. |
| Fair rangeland | 33.17 | 51.11 | 42.93 | 3.34 | 27.99 | 43.09 | 35.65 | 2.70 |
| Poor rangeland | 28.80 | 54.18 | 44.02 | 3.10 | 22.70 | 50.65 | 40.68 | 2.97 |
| Dense forest   | 21.27 | 51.46 | 38.25 | 4.15 | 20.88 | 46.54 | 35.61 | 4.29 |
| Sparse forest  | 33.17 | 52.48 | 43.89 | 2.86 | 22.70 | 48.42 | 39.72 | 3.20 |
| Agriculture    | 32.00 | 52.48 | 42.91 | 5.42 | 27.99 | 49.54 | 39.36 | 4.62 |
| Water body     |       |       |       |      |       |       |       |      |
| Class          |       | 20    | 00    |      |       | 20    | 05    |      |
|                | Min.  | Max.  | Mean  | Std. | Min.  | Max.  | Mean  | Std. |
| Fair rangeland | 30.13 | 49.17 | 39.73 | 3.22 | 20.88 | 40.75 | 30.75 | 3.26 |
| Poor rangeland | 30.13 | 52.85 | 43.92 | 2.56 | 13.34 | 46.16 | 34.41 | 3.97 |
| Dense forest   | 30.98 | 50.65 | 41.30 | 3.20 | 18.11 | 37.97 | 28.57 | 3.22 |
| Sparse forest  | 25.37 | 51.02 | 42.01 | 3.81 | 19.51 | 42.71 | 32.72 | 3.75 |
| Agriculture    | 31.40 | 50.28 | 42.92 | 3.65 | 22.25 | 44.25 | 35.52 | 3.80 |
| Water body     | 27.55 | 40.36 | 32.09 | 3.00 | 23.15 | 31.40 | 24.85 | 1.76 |
| Class          | 2010  |       |       |      | 2015  |       |       |      |
|                | Min.  | Max.  | Mean  | Std. | Min.  | Max.  | Mean  | Std. |
| Fair rangeland | 35.95 | 49.91 | 43.49 | 2.39 | 33.66 | 49.56 | 41.54 | 2.72 |
| Poor rangeland | 29.28 | 57.15 | 48.13 | 3.39 | 28.66 | 52.04 | 44.11 | 2.71 |
| Dense forest   | 30.55 | 57.51 | 44.39 | 4.60 | 26.95 | 50.23 | 37.31 | 4.25 |
| Sparse forest  | 32.24 | 56.09 | 48.20 | 3.17 | 31.46 | 51.39 | 43.85 | 2.85 |
| Agriculture    | 32.24 | 57.15 | 46.17 | 4.60 | 29.24 | 51.41 | 43.92 | 3.70 |
| Water body     | 28.85 | 45.02 | 33.20 | 3.53 | 29.65 | 43.28 | 32.09 | 2.42 |
| Class          |       | 2018  |       |      |       |       |       |      |
|                | Min.  | Max.  | Mean  | Std. |       |       |       |      |
| Fair rangeland | 34.83 | 51.12 | 43.20 | 2.77 |       |       |       |      |
| Poor rangeland | 30.71 | 54.92 | 46.90 | 3.26 |       |       |       |      |
| Dense forest   | 33.82 | 56.10 | 44.75 | 4.43 |       |       |       |      |
| Sparse forest  | 37.71 | 54.57 | 47.87 | 2.25 |       |       |       |      |
| Agriculture    | 33.22 | 55.20 | 46.86 | 4.08 |       |       |       |      |
| Water body     | 31.17 | 45.71 | 33.60 | 2.49 |       |       |       |      |

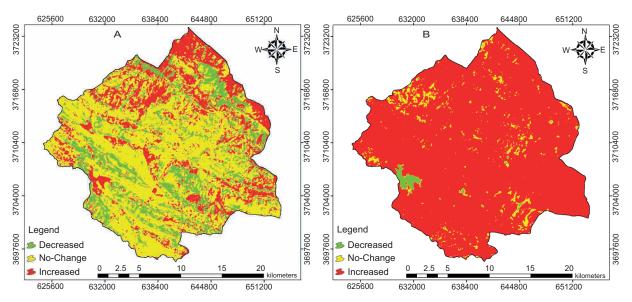


Fig. 6. Change detection map (1990-2018). (a) Land use. (b) LST.

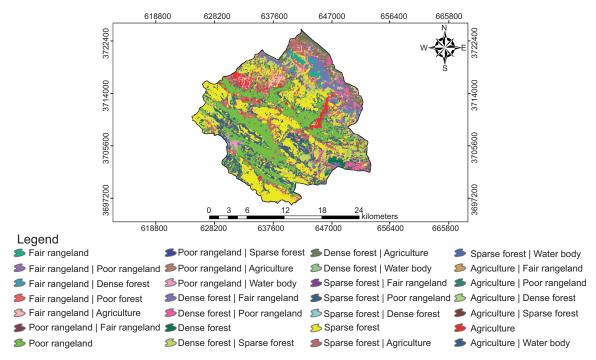


Fig. 7. Land use change map (1990 to 2018). In the map legend, each color indicates how the land use has changed over time. Some colors show two land uses, meaning that the area has changed from one land use to another. Some colors show only one land use, indicating no change in the land use.

Table IV: Results of land use and LST changes.

| LST (°C)<br>Class               | Min.  | Max.  | Mean  | Std. |
|---------------------------------|-------|-------|-------|------|
| Fair rangeland   Agriculture    | -4.6  | 8.4   | 2.9   | 1.4  |
| Dense forest   Poor rangeland   | -1.0  | 10.2  | 6.0   | 1.4  |
| Dense forest   Fair rangeland   | -0.5  | 9.0   | 5.3   | 1.2  |
| Agriculture                     | -5.1  | 14.6  | 3.7   | 2.5  |
| Poor rangeland   Agriculture    | -4.7  | 8.9   | 2.9   | 2.0  |
| Poor rangeland   Fair rangeland | 3.0   | 6.4   | 4.4   | 1.1  |
| Sparse forest   Agriculture     | -4.7  | 9.5   | 3.3   | 1.6  |
| Dense forest   Agriculture      | -3.3  | 9.4   | 4.2   | 1.7  |
| Poor rangeland                  | -10.6 | 9.5   | 4.1   | 1.6  |
| Sparse forest   Fair rangeland  | -0.4  | 7.4   | 3.3   | 1.3  |
| Sparse forest   Poor rangeland  | -10.2 | 9.9   | 4.9   | 1.5  |
| Fair rangeland                  | -1.4  | 7.7   | 3.9   | 1.1  |
| Dense forest   Sparse forest    | -1.4  | 9.8   | 5.0   | 1.6  |
| Sparse forest                   | -2.8  | 9.9   | 4.0   | 1.4  |
| Fair rangeland   Sparse forest  | -1.2  | 8.0   | 3.6   | 1.3  |
| Dense forest                    | -0.4  | 9.1   | 4.9   | 1.4  |
| Agriculture   Sparse forest     | -0.8  | 10.4  | 4.6   | 2.4  |
| Poor rangeland   Sparse forest  | 0.0   | 8.8   | 3.5   | 1.4  |
| Fair rangeland   Poor rangeland | -2.0  | 7.8   | 3.3   | 2.2  |
| Sparse forest   Dense forest    | 0.9   | 7.5   | 4.0   | 1.2  |
| Fair rangeland   Dense forest   | 1.7   | 6.1   | 4.3   | 0.8  |
| Agriculture   Poor rangeland    | -9.0  | 14.1  | 4.9   | 2.8  |
| Agriculture   Fair rangeland    | 4.3   | 7.5   | 5.8   | 0.8  |
| Agriculture   Dense forest      | 2.7   | 5.2   | 4.4   | 1.0  |
| Poor rangeland   Water body     | -19.2 | 0.2   | -10.8 | 3.5  |
| Sparse forest   Water body      | -16.8 | 6.0   | -9.4  | 4.4  |
| Agriculture   Water body        | -17.8 | -4.3  | -12.4 | 3.0  |
| Dense forest   Water body       | -12.4 | -12.3 | -12.3 | 0.1  |

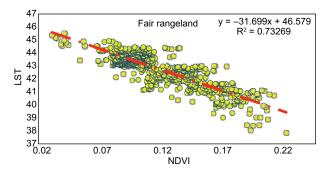


Fig. 8. Correlation chart between LST and NDVI index based on land use.

and NDVI is lower, indicating the low density of vegetation in these land uses. As shown in Figure 8, the correlation of the rangelands with the forest is lower, and their surface temperatures are higher than those of the forest.

# 3.7. Spatiotemporal distribution of LST

To study the spatial distribution of LST in the region, thermal images from different years were normalized by using the highest and lowest amount of surface temperature. Then, using average values and standard deviation, the normalized thermal images were classified into five temperature classes of low, medium,

fairly high, high and very high limit. Figure 9 shows the thermal maps, and Figure 10 shows the area of temperature classes in the desired time periods. The results of the differences in areas between 1990 and 2018, show that the low, medium and fairly high limit classes decreased by 5, 15 and 20%, respectively, and the high and very high limit classes increased by 25 and 15%, respectively.

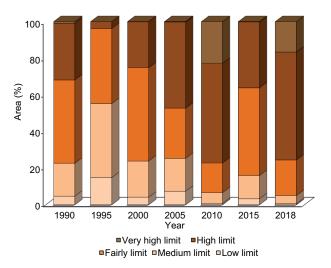


Fig. 10. Area of temperature classes in the studied period (%).

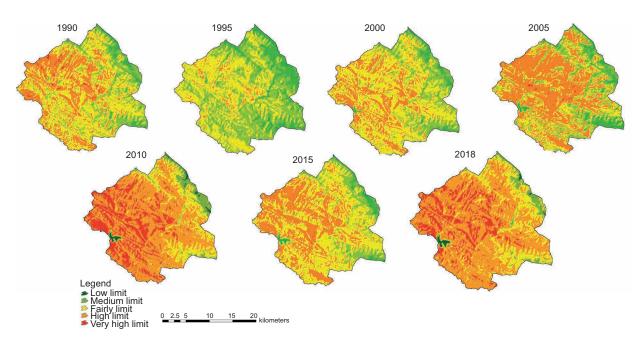


Fig. 9. LST classification maps using average values and standard deviation of thermal images (1990-2018) (in °C).

Due to the decrease in forest area and the increase in human-induced land uses, the area of lower temperature class has decreased, and the upper temperature class has experienced a notable increase during the time analyzed. The reduced area of dense forest and the increasing trend of agricultural and rangeland land uses indicate the replacement and conversion of the region's natural coverage to lower value land uses. According to Figure 8, it can be said that the increase in LST is directly related to the increas in population density within the region and, consequently, the increas in agricultural area. The anticipated continuation of this trend provides grounds to expect a future (continued) rise in temperature (LST).

#### 4. Discussion and conclusions

One of the important results of this research is the demonstration of the ability of fuzzy ARTMAP neural networks to classify land use types. Fuzzy classification includes methods that can provide results better suited to ground reality. In these methods, different values are calculated as the membership grade of each pixel based on land cover variation, while in definitive categorization methods each image pixel is attributed to only one class. The accuracy of the kappa coefficient for land use maps, derived from the satellite data classification by using the fuzzy ARTMAP classification for 1990, 1995, 2000, 2005, 2010, 2015, and 2018 is approximately equal to 87, 92, 90, 90, 89, 94 and 91%, respectively, which represents the high reliability of this algorithm in the classification of satellite data.

The results of the trend in land use change show that, in the period 1990-2018, most changes are related to agricultural lands and low-density (poor) rangeland, which have increased by 18.15 and 16.05%, respectively. On the other hand, land use types of dense forest and sparse forest have decreased by 20.70 and 17.04%, respectively, highlighting the destruction and change of natural and vegetated lands to cultivated lands. Another manmade land use that has many negative effects on nature is the construction of dams. From the time of construction of the Ilam dam (2005) until 2018, the water body behind the dam has increased by about 0.60%.

From LST maps, the range of surface temperature changed from 21.27 to 54.18 °C in 1990 and from

30.07 to 56.09 °C in 2018. Many scholars have focused on the relationship between the effects of land use change and LST. Results of Setturu et al. (2013) in the Uttara Kannada district (India), Pal and Ziaul (2017) in the English Bazar urban center of West Bengal (India), Fathizad et al. (2017) in southwest Iran, and Choudhury et al. (2019) in the Asansol-Durgapur Development show that land use changes are directly related to the increase in surface temperature. By examining the trend of temperature changes in the study area during the period 1990-2018, we observe an increase of 1.92 °C, which is in line with the IPCC report of 2018.

The results of this study show that the LST has a high sensitivity to vegetation cover; land uses with higher vegetation density have lower surface temperatures. Hence, it can be used to detect changes in land use over time. With regard to the effects of land use changes, such as urbanization, establishment of communication roads, agricultural development and soil erosion in the study area, it can be predicted that surface temperature will increase in the future. It is evident that land use changes will result in changes in LST.

Further studies are required in this area under different seasons. Also, to achieve better results, particularly to more accurately estimate the LST, we suggest the use of image sensors with higher spatial resolution of the thermal bonding. In the future, more attention is needed to the modeling of land use changes, specifically by considering climate factors and detecting changes with different types of satellite images. This can reduce some degree of uncertainty in order to support management decisions. The results are potentially useful for various applications, including climatology, hydrology, ecology, geology, design and improvement of transport and agriculture networks.

#### Acknowledgments

This study is related to a project supported by Ilam University, Iran. Also, the experiments comply with the current laws of Iran. We also thank Pedro J. M. Costa, Postdoctoral Research Fellow and Invited Lecturer, Instituto D. Luiz, and Departamento de Geologia, Faculdade de Ciências, Universidade de Lisboa, Portugal, for comments that greatly improved the manuscript.

#### References

- Agam N, Kustas WP, Anderson MC, Li F, Neale CMU. 2007. A vegetation index-based technique for spatial sharpening of thermal imagery. Remote Sensing of Environment 107: 545-558. https://doi.org/10.1016/j.rse.2006.10.006
- Allison EW. 1989. Monitoring drought affected vegetation with AVHRR. Digest International Geoscience and Remote Sensing Symposium 4, 1965-1967.
- Amiri R, Weng Q, Alimohammadi A, Alavipanah SK. 2009. The spatial-temporal dynamics of land surface temperatures in relation to fractional vegetation cover and land use/cover in the Tabriz urban area, Iran. Remote Sensing of Environment 113: 2606-2617. https://doi.org/10.1016/j.rse.2009.07.021
- Artis DA, Carnahan WH. 1982. Survey of emissivity variability in thermography of urban areas. Remote Sensing of Environment 12: 313-329. https://doi.org/10.1016/0034-4257(82)90043-8
- Baihua F, Isabela B. 2015. Riparian vegetation NDVI dynamics and its relationship with climate, surface water and groundwater. Journal of Arid Environments 113: 59-68. https://doi.org/10.1016/j.jaridenv.2014.09.010
- Brovkin V, Boysen L, Arora VK, Boisier JP, Cadule P, Chini L, Claussen M, Friedlingstein P, Gayler V, Van Den Hurk BJJM, Hurtt GC, Jones CD, Kato E, Noblet-Ducoudré N, Pacifico F, Pongratz J, Weiss M. 2013. Effect of anthropogenic land-use and land-cover changes on climate and land carbon storage in CMIP5 projections for the twenty-first century. Journal of Climate 26: 6859-6881. https://doi.org/10.1175/JCLI-D-12-00623.1
- Carpenter GA, Grossberg S, Reynolds JH. 1991. ART-MAP: Supervised real-time learning and classification of nonstationary data by a self-organizing neural network. Neural Networks 4: 565-588. https://doi.org/10.1016/0893-6080(91)90012-T
- Chavez PS. 1988. An improved dark-object subtraction technique for atmospheric scattering correction of multispectral data. Remote Sensing of Environment 24, 459-479. https://doi.org/10.1016/0034-4257(88)90019-3
- Chavez PS, MacKinnon DJ. 1994. Automatic detection of vegetation changes in the southwestern United States using remotely sensed images. Photogrammetric Engineering and Remote Sensing 60: 571-583.
- Choudhury D, Das K, Das A. 2019. Assessment of land use land cover changes and its impact on variations of land surface temperature in Asansol-Durgapur

- Development Region. The Egyptian Journal of Remote Sensing and Space Science 22: 203-218. https://doi.org/10.1016/j.ejrs.2018.05.004
- Connors JP, Galletti CS, Chow WTL. 2013. Landscape configuration and urban heat island effects: Assessing the relationship between landscape characteristics and land surface temperature in Phoenix, Arizona. Landscape Ecology 28: 271-283. https://doi.org/10.1007/s10980-012-9833-1
- Deng Y, Wang S, Bai X, Tian Y, Wu L, Xiao J, Chen F, Qian Q. 2018. Relationship among land surface temperature and LUCC, NDVI in typical karst area. Scientific Reports 8: 641. https://doi.org/10.1038/s41598-017-19088-x
- Du Y, Teillet PM, Cihlar J. 2002. Radiometric normalization of multitemporal high-resolution satellite images with quality control for land cover change detection. Remote Sensing of Environment 82: 123-134. https://doi.org/10.1016/S0034-4257(02)00029-9
- Fathizad H, Tazeh M, Kalantari S, Shojaei S. 2017. The investigation of spatiotemporal variations of land surface temperature based on land use changes using NDVI in southwest of Iran. Journal of African Earth Sciences 137: 249-256. https://doi.org/10.1016/j.iafrearsci.2017.06.007
- Feddema JJ, Oleson KW, Bonan GB, Mearns LO, Buja LE, Meehl GA, Washington WM. 2005. The importance of land-cover change in simulating future climates. Science 310: 1674-1678. https://doi.org/10.1126/science.1118160
- Feizizadeh B, Blaschke T, Nazmfar H, Akbari E, Kohbanani HR. 2012. Land surface temperature relationship to land use/land cover from satellite imagery in Maraqeh County, Iran. Journal of Environmental Planning and Management 56: 1290-1315. https://doi.org/10.1080/09640568.2012.717888
- Feng JM, Wang YL, Ma ZG, Liu YH. 2012. Simulating the regional impacts of urbanization and anthropogenic heat release on climate across China. Journal of Climate 25: 7187-7203. https://doi.org/10.1175/JCLI-D-11-00333.1
- Halder S, Saha SK, Dirmeyer PA, Chase TN, Goswami BN. 2016. Investigating the impact of land-use land-cover change on Indian summer monsoon daily rainfall and temperature during 1951-2005 using a regional climate model. Hydrology and Earth System Sciences 20: 1765-1784. https://doi.org/10.5194/hess-20-1765-2016

- Hao X, Li W, Deng H. 2016. The oasis effect and summer temperature rise in arid regions-case study in Tarim Basin. Scientific Reports 6: 35418. https://doi.org/10.1038/srep35418
- Hasanlou M, Mostofi N. 2015. Investigating urban heat island estimation and relation between various land cover indices in Tehran City using Landsat 8 imagery. In: 1st International Electronic Conference on Remote Sensing.
- Herb WR, Janke B, Mohseni O, Stefan HG. 2008. Ground surface temperature simulation for different land covers. Journal of Hydrology 356: 327-343. https://doi.org/10.1016/j.jhydrol.2008.04.020
- Hou GL, Zhang HY, Wang YQ, Qiao ZH, Zhang ZX. 2010. Retrieval and spatial distribution of land surface temperature in the middle part of Jilin province based on MODIS data. Scientia Geographica Sinica 30: 421-427. https://doi.org/10.13249/j.cnki.sgs.2010.03.421
- Inamdar AK, French A, Hook S, Vaughan G, Luckett W. 2008. Land surface temperature retrieval a high spatial and temporal resolution over the southwestern United States. Journal of Geophysical Research: Atmospheres 113: D07107. https://doi.org/10.1029/2007JD009048
- Isaya Ndossi M, Avdan U. 2016. Application of open source coding technologies in the production of land surface temperature (LST) maps from Landsat: A PyQGIS plugin. Remote Sensing 88: 413. https://doi.org/10.3390/rs8050413
- Jiménez-Muñoz JC, Sobrino JA. 2010. A single-channel algorithm for land-surface temperature retrieval from ASTER data. IEEE Geoscience and Remote Sensing Letters 7: 176-179. https://doi.org/10.1109/ LGRS.2009.2029534
- Kustas WP, Norman JM, Anderson MC, French AN. 2003. Estimating subpixel surface temperatures and energy fluxes from the vegetation index-radiometric temperatures relationship. Remote Sensing of Environment 85: 429-440. https://doi.org/10.1016/S0034-4257(03)00036-1
- Latif MS. 2014. Land surface temperature retrieval of Landsat-8 data using split window algorithm A Case study of Ranchi District. International Journal of Engineering Development and Research 2: 3840-3849.
- Li S, Jiang GM. 2018. Land surface temperature retrieval from Landsat-8 data with the generalized split-window algorithm. IEEE Access 6: 18149-18162. https://doi.org/10.1109/ACCESS.2018.2818741

- Lillesand T, Kiefer RW, Chipman J. 2015. Remote sensing and image interpretation. John Wiley & Sons, 736 pp.
- Lu D, Mausel P, Brondizio E, Moran E. 2004. Change detection techniques. International Journal of Remote Sensing 25: 2365-2407. https://doi. org/10.1080/0143116031000139863
- Luyssaert S, Jammet M, Stoy PC, Estel S, Pongratz J, Ceschia E, Churkina G, Don A, Erb K, Ferlicoq M, Gielen B, Grunwald T, Houghton RA, Klumpp K, Knohl A, Kolb T, Kuemmerle T, Laurila T, Lohila A, Loustau D, McGrath MJ, Meyfroidt P, Moors EJ, Naudts K, Novick K, Otto J, Pilegaard K, Pio CA, Rambal S, Rebmann C, Ryder J, Suyker AE, Varlagin A, Wattenbach M, Dolman AJ. 2014. Land management and land-cover change have impacts of similar magnitude on surface temperature. Nature Climate Change 4: 389-393. https://doi.org/10.1038/nclimate2196
- Mahmood R, Pielke RA, Hubbard KG, Niyogi D, Dirmeyer PA, McAlpine C, Carleton AM, Hale R, Gameda S, Beltrán-Przekurat A, Baker B, McNider R, Legates DR, Shepherd M, Du J, Blanken PD, Frauenfeld OW, Nair U, Fall S. 2014. Land cover changes and their biogeophysical effects on climate. International Journal of Climatology 34; 929-953. https://doi.org/10.1002/joc.3736
- Mesgari S. 2002. Investigation of the trend of forest area change using GIS and RS, Tehran. Research Project, Faculty of Engineering, K.N. Toosi University of Technology, Tehran, Iran.
- Mirabolfathi M. 2013. Outbreak of charcoal disease on *Quercus* spp. and *Zelkova carpinifolia* trees in forests of Zagros and Alborz mountains in Iran. Iranian Journal of Plant Pathology 49: 257-263.
- Pal S, Ziaul Sk. 2017. Detection of land use and land cover change and land surface temperature in English Bazar urban center. The Egyptian Journal of Remote Sensing and Space Sciences 20: 125-145. https://doi.org/10.1016/j.ejrs.2016.11.003
- Qian C, Deyong Y, Matei G, Zhe H, Jianguo W. 2015. Impacts of land use and land cover change on regional climate: a case study in the agro-pastoral transitional zone of China. Environmental Research Letters 10: 124025. https://doi.org/10.1088/1748-9326/10/12/124025
- Reutter HF, Olesen S, Fischer H. 1994. Distribution of the brightness temperature of land surfaces determined from AVHRR data. International Journal of Remote Sensing 15: 95-104. https://doi.org/10.1080/01431169408954053

- Ripley BD. 1977. Modelling spatial patterns. Journal of the Royal Statistical Society: Series B (Methodological) 39: 172-192. https://doi.org/10.1111/j.2517-6161.1977. tb01615.x
- Richards JA. 2013. Remote sensing digital image analysis. An introduction. 5th ed. Springer, 516 pp. https://doi.org/10.1007/978-3-642-30062-2
- Rozenstein O, Qin Z, Derimian Y, Karnieli A. 2014. Derivation of land surface temperature for Landsat-8 TIRS using a split window algorithm. Sensors 14: 5768-5780. https://doi.org/10.3390/s140405768
- Setturu B, Rajan KS, Ramachandra TV. 2013. Land surface temperature responses to land use land cover dynamics. Geoinformatics & Geostatistics: An Overview 1: 4. https://doi.org/10.4172/2327-4581.1000112
- Snyder WC. 1998. Classification based emissivity for land surface temperature measurement from pace. International Journal of Remote Sensing 19: 2753-2774. https://doi.org/10.1080/014311698214497
- Srivastava PK, Majumdar TJ, Bhattacharya AK. 2009. Surface temperature estimation in Singhbhum shear zone of India using Landsat-7 ETM + thermal infrared data. Advances in Space Research 43: 1563-1574. https://doi.org/10.1016/j.asr.2009.01.023
- Tariq A, Shu H. 2020. CA-Markov chain analysis of seasonal land surface temperature and land use land cover change using optical multi-temporal satellite data of Faisalabad, Pakistan. Remote Sensing 12: 3402. https://doi.org/10.3390/rs12203402
- Tariq A, Riaz I, Ahmad Z, Yang B, Amin M, Kausar R, Andleeb S, Farooqi MA, Rafiq M. 2020. Land surface temperature relation with normalized satellite indices for the estimation of spatio-temporal trends in temperature among various land use land cover classes of an arid Potohar region using Landsat data. Environmental Earth Sciences 79: 40. https://doi.org/10.1007/s12665-019-8766-2

- Tomlinson CJ, Chapman L, Thrones JE, Baker C. 2011. Remote sensing land surface temperature for meteorology and climatology: A review. Meteorological Applications 18: 296-306. https://doi.org/10.1002/met.287
- Waters W, Allen R, Masahiro T, Trezza R, Bastiaanssen W. 2002. SEBAL. Surface energy balance algorithms for land. Version 1.0. NASA EOSDIS/Raytheon Company/ Idaho, Department of Water Resources, 97 pp.
- Wei L, Jean-Daniel MS, Thomas WG. 2015. A comparison of the economic benefits of urban green spaces estimated with NDVI and with high-resolution land cover data. Landscape and Urban Planning 133: 105-117. https://doi.org/10.1016/j.landurbplan.2014.09.013
- Weng Q, Lu D, Schubring J. 2004. Estimation of land surface temperature-vegetation abundance relationship for urban heat island studies. Remote Sensing of Environment 89: 467-483. https://doi.org/10.1016/j. rse.2003.11.005
- Xu W, Gu S, Zhao X, Xiao J, Tang Y, Fang J, Zhang J, Jiang S. 2011. High positive correlation between soil temperature and NDVI from 1982 to 2006 in alpine meadow of the Three-River Source Region on the Qinghai-Tibetan Plateau. International Journal of Applied Earth Observation and Geoinformation 13: 528-535. https://doi.org/10.1016/j.jag.2011.02.001
- Ziaul Sk, Pal S. 2018a. Anthropogenic heat flux in English Bazar Town and its surroundings in West Bengal, India. Remote Sensing Applications: Society and Environment 11: 151-160. https://doi.org/10.1016/j.rsase.2018.06.003
- Ziaul Sk, Pal S. 2018b. Analyzing control of respiratory particulate matter on land surface temperature in local climatic zones of English Bazar Municipality and surroundings. Urban Climate 24: 34-50. https://doi.org/10.1016/j.uclim.2018.01.006