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WHEAT YIELD AND PHYSICAL PROPERTIES OF A BROWN LATOSOL UNDER NO-TILLAGE IN SOUTH-CENTRAL PARANÁ[1]

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SUMMARY

Soil management influences the chemical and physical properties of soil. Chemical conditions have been thoroughly studied, while the role of soil physical conditions regarding crop yield has been neglected. This study aimed to analyze the wheat yield and its relationship with physical properties of an Oxisol under no-tillage (NT). The study was carried out between 2010 and 2011, in Reserva do Iguaçu, State of Paraná, Brazil, on the Campo Bonito farm, after 25 years of NT management. Based on harvest maps of barley (2006), wheat (2007) and maize (2009) of a plot (150 ha), zones with higher and lower yield potential (Z1 and Z2, respectively) were identified. Sampling grids with 16 units (50 x 50 m) and three sampling points per unit were established. The wheat grain yield (GY) and water infiltration capacity (WIC) were evaluated in 2010. Soil samples with disturbed and undisturbed structure were collected from the 0.00-0.10 and 0.10-0.20 m layers. The former were used to determine soil organic carbon (Corg) levels and the latter to determine soil bulk density (BD), total porosity (TP), macroporosity (Mac), and microporosity (Mic). Soil penetration resistance (PR) and water content (SWC) were also evaluated. The wheat GY of the whole plot was close to the regional average and the yield between the zones differed significantly, i.e. 22 % higher in Z1 than in Z2. No significant variation in Mic was observed between zones, but Z1 had higher Corg levels, SWC, TP and Mac and lower BD than Z2 in both soil layers, as well as a lower PR than Z2 in the 0.00-0.10 m layer. Therefore, soil physical conditions were more restrictive in Z2, in agreement with wheat yield and zone

yield potential defined *a priori*, based on the harvest maps. Soil WIC in Z1 was significantly higher (30 %) than in Z2, in agreement with the results of TP and Mac which were also higher in Z1 in both soil layers. The correlation analysis of data of the two layers showed a positive relationship between wheat GY and the soil properties TP, SWC and WIC.

Index terms: soil physical quality, wheat yield, correlation.

RESUMO: *PRODUTIVIDADE DO TRIGO E ATRIBUTOS FÍSICOS DE UM LATOSSOLO BRUNO SOB PLANTIO DIRETO NO CENTRO-SUL DO PARANÁ*

O manejo influencia atributos químicos e físicos do solo; entretanto, os químicos têm sido mais estudados, negligenciando-se a relevância da física do solo para a produtividade das culturas. Os objetivos deste trabalho foram avaliar a produtividade do trigo e caracterizar a sua relação com os atributos físicos de um Latossolo Bruno, sob sistema plantio direto (NT). O estudo foi realizado entre 2010-2011 em Reserva do Iguaçu, Paraná, em um dos talhões da fazenda Campo Bonito, com 25 anos de NT. Os mapas de colheita de cevada (2006), trigo (2007) e milho (2009) permitiram identificar as zonas Z1 (potencial produtivo maior) e Z2 (potencial produtivo menor), em que se estabeleceram malhas amostrais com 16 unidades de 50 x 50 m e três pontos de amostragem por unidade. Em 2010, avaliaram-se a produtividade de grãos (GY) do trigo e a capacidade de infiltração de água no solo (WIC). Amostras de solo com estrutura deformada e indeformada foram coletadas nas camadas de 0,00-0,10 e 0,10-0,20 m. As primeiras serviram para determinar o teor de carbono orgânico (Corg) e as últimas para determinar: densidade do solo (BD), porosidade total do solo (TP), macroporosidade (Mac) e microporosidade (Mic). Avaliaram-se, também, a resistência do solo à penetração (PR) e o conteúdo de água do solo (SWC). A GY do trigo no talhão foi próxima da média regional e houve diferença significativa de rendimento entre zonas, sendo 22 % superior em Z1 em relação a Z2. Não houve variação significativa da Mic entre as zonas do talhão; Z1 apresentou teores de Corg, SWC, TP e Mac maiores e BD menor que Z2, nas duas camadas de solo avaliadas, além de PR menor que Z2, na camada de 0,00-0,10 m, havendo, portanto, condição física de solo mais restritiva em Z2, condizente com os resultados de produtividade do trigo e potencial produtivo delimitado a priori, a partir dos mapas de colheita. A WIC foi significativamente maior (30 %) em Z1 que em Z2, em acordo com os resultados de TP e Mac, também maiores em Z1, nas duas camadas de solo. Considerando-se essas duas camadas, a análise de correlação permitiu destacar relações positivas entre a GY do trigo e TP, SWC e WIC.

Termos de indexação: qualidade física do solo, rendimento do trigo, correlação.

INTRODUCTION

Brazil has the world's second largest area under no-tillage (NT) agriculture, with a total of more than 25.5 million hectares under this management (Derpsch et al., 2010). Considered one of the greatest advances achieved in Brazilian agriculture, NT is primarily a conservation system, because it significantly reduces water erosion and enhances the organic carbon stocks in soils (Corg) (Lopes et al., 2004), also increasing the efficiency in terms of storage and recycling of water and nutrients and biological activity (Ceretta et al., 2002; Resende, 2011).

On the other hand, soil management with NT may cause a reduction in soil porosity and an increase in soil bulk density (BD) (Tormena et al., 2002), greater penetration resistance (PR) and decreased soil permeability (Silva et al., 2009), which can reduce the availability of water and nutrients to plants and affect the crop performance (Giarola et al., 2009). Therefore, the soil physical quality needs to be monitored, to identify the best management practices to maintain the sustainability of agricultural systems (Beutler et al., 2009).

Most efforts regarding monitoring, however, have been based on soil chemical properties, because the collection of information related to soil physics is more difficult. In this case, according to Santi (2007), sampling would be more reasonable if it were directed or if it addressed areas with high and low crop yield, which should be as small and homogeneous as the technical level and costs allow (Dodermann & Ping, 2004). To characterize the interrelation between the yield potential of a wheat crop and soil properties, Mulla & Bhatti (1997) divided a study area in units of low, moderate and high yield.

According to Guimarães (2011), the loss of soil physical quality is evidenced by a reduction of water

infiltration in the soil, increased erosion susceptibility and mechanical impediments to root penetration. As it is related to plant growth and is easily and rapidly determined, PR has been used as soil compaction indicator (Mercante et al., 2003), with critical values for plants established between 1.5 and 4.0 MPa, depending on the species, while a value of 2.0 MPa is generally accepted as the limit above which there is an impediment to root growth of agricultural crops (Imhoff et al., 2000).

In turn, BD is directly related to soil water retention and infiltration, root development, gas exchange, and soil susceptibility to erosion, and constitutes the most direct quantitative measurement of soil compaction (Reichert et al., 2007). Critical values of BD for different textural soil classes were proposed by Reinert & Reichert (2001): 1.30 Mg m^{-3} for soils above 55 % clay; 1.50 Mg m^{-3} for soils between 20 and 55 % clay; and 1.80 Mg m^{-3} for soils below 20 % clay.

Water infiltration into the soil is related to the structural arrangement of the particles and, therefore, to the BD and total porosity (TP) of the soil. Measuring the water infiltration capacity (WIC) into the soil, in turn, determines the amount of water (irrigation, rain) an area can tolerate without occurrence of surface flow (Pott et al., 2005), so WIC is an important indicator of soil conditions for plant growth and of water, nutrient and soil loss. According to Islan & Weil (2000), permeability to water reflects the structural quality and stability of a soil. Studying no-tillage fields, Santi (2007) concluded that water infiltration into the soil was one of the soil physical properties with greatest influence on soybean, wheat and maize yield.

Inversely proportional to BD, soil TP is divided into microporosity (Mic), representing pores in which water is retained for plant absorption, and macroporosity (Mac), representing pores from which water drains and, therefore, gas exchange occurs (Kiehl, 1979). Increasing BD values result in modifications of the soil porous spaces, arising from TP reduction, normally at the expense of Mac, which may reduce gas exchange between the soil and atmosphere and increase soil PR (Letey, 1985; Blainski et al., 2009).

Soil Mic is highly influenced by the texture and Corg level and weakly influenced by traffic of agricultural machinery (Silva & Kay, 1997). Soil Mac, which is altered by compaction, may be used to evaluate the performance of soil management systems in relation to crop yield, since values less than 0.10 m^3 m^{-3} of air-filled porosity limit root growth (Tormena et al., 1998). According to Araújo (2004), porosity is an indicator of soil alterations caused by agricultural use, reflecting the soil quality, for directly affecting the water infiltration velocity, gas exchange, microbial life, and root growth.

This study was carried out on a Brown Latosol (Oxisol) under no-tillage management, on a wheat field in south-central Paraná, to investigate wheat yield and its relation with properties of soil physical quality.

MATERIAL AND METHODS

The study was conducted between 2010 and 2011 on the Campo Bonito farm, a 5,000 ha property in the Reserva do Iguaçu, Paraná State. For at least 25 years, the no-tillage management had been used and as of 2005, precision agriculture techniques were adopted, based on yield maps. The climate is Cfb (Köppen), with a mild summer and severe frosts in winter (IAPAR, 2000). Pluvial precipitation data from the meteorological station of the Instituto Agronômico do Paraná (IAPAR, 2011), located in Guarapuava (100 km from the field), are shown in figure 1, together with the precipitation recorded on the farm (pluviometer) in 2010. According to the soil map of the State of Paraná (Bhering & Santos, 2008), the area is located in the mapping unit LBd5, with dystroferric Brown Latosols (Oxisols).

In May 2010, one of the plots (150 ha) was selected and the soil sampled (0.00-0.20 m layer), to determine the chemical fertility and texture (Table 1). Considering the difficulties with soil sampling and laboratory analysis of soil physical properties for areas of many hectares, and in agreement with Mulla & Bhatti (1997) and Santi (2007), the available yield maps were used to separate two sampling zones in the plot, Z1 and Z2 (Table 2), with distinct yield potential, also aiming to avoid the restriction of range (restriction of variability) of the variables, which could impair the correlation analysis (Lira, 2004).

To allow comparisons, the selection of Z1 and Z2 considered similarities in terms of landscape (shoulder position) and slope (≤ 5 %). The crop rotations, managed homogeneously in both zones, consisted of: two soybean crops for each crop of maize in summers; in the winter, wheat and barley were interspersed with oats for grazing. The occurrence of these phases is variable, depending on the target profitability. A rotation grazing system was used in the winter, with 7 days of grazing and 28 days of pasture regrowth in paddocks subdivided by electric fencing. When the plants had sufficient biomass, the animals grazed again for 7 days, rotating to neighboring paddocks at the end of each grazing period.

As the crop to be evaluated in 2010 was wheat, only the yield maps of grass crops were used to delimit Z1 and Z2 (Table 2): barley (2006), wheat (2007) and maize (2009). Barley (BRS 195) and wheat (cv. Ônix) were sown in a row spacing of 0.17 m, in stands of 280 plants m^{-2} and fertilized with 280 kg ha^{-1} of the NPK 08-30-20 fertilizer and 100 kg ha^{-1} of nitrogen (N), with urea topdressing at tillering. Maize (Pioneer 30R50) was sown in a row spacing of 0.70 m, 75,000 plants ha^{-1} and fertilized with 350 kg ha^{-1} of NPK

Table 1. Soil fertility and particle size analysis* of the plot at the beginning of the study

Corg	P[1]	pH (CaCl$_2$)	Al^{3+}	H+Al	Ca^{2+}	Mg^{2+}	K$^+$	Sand	Silt	Clay
g dm^{-3}	mg dm^{-3}				cmol$_c$ dm^{-3}				g kg^{-1}	
32	6.0	5.0	0.27	6.20	1.92	0.80	0.38	45	234	721

[1] Mehlich-1. *According to Pavan et al. (1992) and Embrapa (1997), respectively.

Table 2. Descriptive statistics of the data from the yield maps of barley (2006), wheat (2007) and maize (2009) in the zones of the plot selected for the study

Zone	Mean	Minimum	Maximum	SD[1]	CV[2]
			kg ha^{-1}		
		Barley 2006			
Z1	3,462	1,750	4,995	650	19
Z2	3,126	1,752	4,998	687	22
		Wheat 2007			
Z1	3,312	1,750	4,999	761	23
Z2	1,491	1,000	3,808	327	22
		Maize 2009			
Z1	10,218	4,038	13,709	1,811	18
Z2	7,532	4,857	13,349	1,150	15

[1]SD: standard deviation; [2]CV: coefficient of variation.

10-26-24 fertilizer and topdressed with 200 kg ha^{-1} N (urea) at the eight-leaf stage. The phytosanitary control and fertilization were performed according to the official recommendations for the crops (Embrapa, 2008a,b; 2009).

A regular grid with 16 sampling units of 50 x 50 m was established in each zone, covering 8 ha of the total plot. In each unit, a diagonal line was set up with three equidistant points defined for sampling, making up 96 points in the plot. The evaluated wheat crop (cv. BRS Guamirim) was sown in June 2010, in a row spacing of 0.17 m, in stands of 280 plants m^{-2} fertilized with 330 kg ha^{-1} of 09-25-24 NPK fertilizer and topdressed with 100 kg ha^{-1} N (urea) at tillering. Yield was determined in November 2010 after manual harvest of an area of 1 m^2 per sampling point. After drying and threshing, the grains were weighed and the moisture content was corrected to 130 g kg^{-1}.

The soil was evaluated at each point. Using a constant head permeameter (model IAC), which operates on the Mariotte siphon principle, WIC was determined using a controlled hydraulic head of 0.06 m at a depth of 0.10 m, in holes opened with a 0.06 m diameter auger. Measures were taken at 1 min intervals until constant flow was achieved, i.e. after five consecutive readings of the same value. WIC was calculated using equation 1 (Pott & De Maria, 2003):

$$WIC = q \times 60 \times \left(\frac{Dp^2}{Do^2 + (4 \times Do \times H)} \right) \quad (1)$$

where q is the constant water flow of the permeameter (mm min^{-1}), Dp is the permeameter diameter (mm), Do is the soil hole diameter (mm) and H is the hydraulic head (mm).

The soil was also sampled at each point, with disturbed (auger) and undisturbed soil structure (volumetric rings of 0.0001 m^3) in the 0.00-0.10 and 0.10-0.20 m layers. The disturbed samples were used to determine the Corg levels, according to Pavan et al. (1992). The undisturbed samples were prepared with nylon cloth fixed to the lower end of the rings and saturated in water for 48 h, weighed and left to drain on a tension table at -0.006 MPa. After drainage, they were weighed again and the volumetric water content was calculated to determine the soil Mic (Embrapa, 1997). Then, the soil was dried to constant weight at ±105 °C. The soil dry mass and volume (ring volume) were used to determine BD (Blake & Hartge, 1986).

The TP was obtained using the relation between BD and soil particle density (Dp) (Danielson & Sutherland, 1986), while soil Mac was obtained by the difference between TP and Mic. Soil Dp was determined by the volumetric flask method with ethyl alcohol (Embrapa, 1997), using the soil from the core ring samples. In July 2011, soil PR was evaluated with a Falker® penetrometer at each 0.01 m (Tormena & Roloff, 1996) to a depth of 0.20 m, and the means were calculated for the 0.00-0.10 and 0.10-0.20 m layers. Concomitantly, soil samples were collected with an auger from the same layers to determine the soil water content (SWC) (Embrapa, 1997). All determinations were performed at the Soil and Plant Nutrition Laboratory of the South-Central State University - UNICENTRO.

Data were subjected to descriptive statistics and to a normality test. Averages per zone and depth were compared by the *t*-test for independent samples, using the ASSISTAT (2011) package. Simple linear correlation analysis between variables was performed using the SPSS18 package (SPSS, 2009).

RESULTS AND DISCUSSION

The yield of the wheat crop evaluated in 2010 is shown in table 3. According to Warrick & Nielsen (1980), the coefficient of variation (CV) can be classified into low (≤ 12 %), moderate (12 ≤ CV ≤ 62 %) or high

(> 62 %); thus, the variability of the wheat yield on the plot was considered moderate (27 %). At the regional level of Guarapuava-PR, were the farm is located, the average wheat yield is 3,100 kg ha^{-1} (Maggian et al., 2010), i.e., the mean plot yield was very close to the regional average. There was a significant difference between Z1 and Z2, confirming the difference observed *a priori* on the yield maps (Table 2). The yield in Z1 was 22 % higher than in Z2. The zones had similar minimum, but differed in the maximum values.

The results for soil Corg, PR, SWC, BD, TP, Mic and Mac are shown in table 4. Low variability was recorded for Corg in both layers (0.00-0.10 and 0.10-0.20 m) and mean values were considered very high, according to Serrat et al. (2006), indicating an

adequate maintenance of soil organic matter levels in the plot by the management system, also consistent with the cold and humid climate conditions of the region (Cfb). There was a significant difference between Z1 and Z2 in both soil layers, with 6.1 and 12.6 % higher Corg levels, respectively, in 0.00-0.10 and 0.10-0.20 m of Z1. In both zones, Corg was higher in the upper layer, which is explained by the use of NT for more than 25 years, without straw incorporation, which is left as mulch on the soil surface.

There was low variability in PR in both soil layers, contrasting with the observations of Cherubin et al. (2011), who reported a CV > 62 % for PR in a dystrophic Red Latosol (Oxisol) under NT, but close to the values observed by Tavares Filho et al. (2012) in a eutroferric Red Latosol under NT, with a CV = 22 % for PR. The low variability in this study may be due to the high number of observations (96) in the plot. In comparison, Tavares Filho & Ribon (2008) found that the number of representative samples required to ensure data accuracy (acceptable variability) of PR was equal to or greater than 15 under NT and perennial crops and equal to or greater than 20 under conventional soil management.

Regarding the 0.00-0.10 m layer, PR was higher in Z2 than Z1 and although soil SWC was lower in this zone, which could overestimate PR (Busscher et al., 1997), BD was also higher in Z2. According to Busscher (1990) and Busscher et al. (1997), inappropriate management may increase PR and BD,

Table 3. Descriptive statistics and Student's *t*-test for wheat yield, evaluated in 2010 in the zones of the plot selected for the study

Zone	Mean	Minimum	Maximum	SD[1]	CV[2]
			kg ha^{-1}		
Z1	3,431A*	1,326	5,887	762	22
Z2	2,676B	1,379	4,799	721	27
Plot mean	3,054	1,326	5,887	829	27

[1]SD: standard deviation; [2]CV: coefficient of variation. *Means followed by different letters indicate significant difference (p<0.05) by Student's t-test.

Table 4. Descriptive statistics and Student's *t*-test for organic carbon (Corg), soil penetration resistance (PR), soil water content (SWC), soil bulk density (BD), soil total porosity (TP), microporosity (Mic) and macroporosity (Mac) in the zones of the plot selected for the study

Zone	Corg	PR	SWC	BD	TP	Mic	Mac
	g dm^{-3}	MPa	kg kg^{-1}	Mg m^{-3}		m^3 m^{-3}	
				0.00-0.10 m			
Z1	34.94 Aa*	1.50 Bb	0.46 Aa	1.05 Ba	0.61 Ab	0.51 Ab	0.09 Aa
Z2	32.80 Ba	1.65 Ab	0.42 Ba	1.12 Aa	0.58 Bb	0.52 Ab	0.06 Ba
Plot Mean	33.87	1.58	0.44	1.09	0.59	0.52	0.08
Minimum	22.78	0.98	0.38	0.93	0.54	0.43	0.02
Maximum	39.41	2.40	0.52	1.29	0.66	0.59	0.22
SD[1]	1.82	0.24	0.02	0.06	0.02	0.02	0.03
CV (%)[2]	6.49	9.65	3.55	3.52	2.49	2.50	24.45
				0.10-0.20 m			
Z1	32.66 Ab	2.33 Aa	0.47 Aa	1.01 Bb	0.64 Aa	0.56 Aa	0.08 Aa
Z2	28.54 Bb	2.32 Aa	0.43 Ba	1.13 Aa	0.61 Ba	0.55 Aa	0.06 Ba
Plot Mean	30.60	2.33	0.45	1.07	0.62	0.55	0.07
Minimum	19.84	1.63	0.39	0.86	0.55	0.47	0.01
Maximum	38.47	3.12	0.54	1.27	0.70	0.62	0.14
SD	1.50	0.24	0.02	0.06	0.02	0.03	0.03
CV (%)	6.34	6.71	4.22	4.07	3.21	4.32	21.62

*Means followed by different letters indicate significant difference (p<0.05) by Student's *t*-test. Upper case letters compare zones in the same layer and lower case letters compare layers at same zone. [1]SD: standard deviation; [2]CV: coefficient of variation.

properties which are directly related. In the 0.10-0.20 m layer, there was no difference of PR between zones, but PR increased from the 0.00-0.10 to the 0.10-0.20 m layer in both. In areas where NT is consolidated, as in the case of this study, the soil tends to compact in the subsurface (0.07-0.15 m) (Silva et al., 2000; Stone & Silveira, 2001). On the surface, the mobilization of the soil by furrower mechanisms during sowing contributes to lower PR values, as does the higher Corg levels observed (Table 4), since they are inversely correlated to PR (Melo et al., 2008; Schiavo & Colodro, 2012).

Although some papers cited PR thresholds for soybean and rice root growth close to 3.0 MPa (Mielniczuk et al.,1985 - for rice; Beutler, 2003 - for soybean and rice), the most commonly cited values are between 2.0 and 2.5 MPa (Taylor et al., 1966), and 2.0 MPa is the most cited threshold (Imhoff et al., 2000; Lapen et al., 2004). Therefore, the PR in Z1 and Z2 was potentially limiting for root growth in the 0.10-0.20 m layer, with means close to 2.3 MPa. This was aggravated by the fact that soil SWC was high during the PR evaluations for Z1 and Z2 (0.43 and 0.47 kg kg^{-1}), since a lower SWC would raise the PR values. Indeed, soil limitations must have been more restrictive in Z2 than in Z1 for the wheat crop in 2010, since PR in the 0.00-0.10 m layer, measured in moist soil, was higher in Z2 than Z1. But figure 1 shows below-average rainfall in the months of wheat cultivation, leading to dry soil, and consequently, greatly increased PR.

Low variability was also recorded for BD in both soil layers, and the values decreased from 0.00-0.10 to 0.10-0.20 m in Z1, but did not change significantly in Z2, which in turn had significantly higher values than Z1 in both layers. However, none of the BD values reached the critical limit of 1.30 Mg m^{-3} proposed by Reinert & Reichert (2001) for clayey soils (as that investigated in this study). Accordingly, the data dispersion for TP and Mic of soil porosity was low and moderate for Mac (Table 4).

In both soil layers, TP and Mac were lower in Z2 than Z1, but no significant difference between zones was observed for Mic. With regard to the depth ranges, the Mic values in both zones were higher in 0.00-0.10 than in 0.10-0.20 m, with no changes in Mac values, showing that the increase in TP observed in the deeper layer occurred due to increased Mic. The data also revealed a high Mic/Mac ratio, with Mic values always above 0.50 m^3 m^{-3}, which is 5.5 to 9.2 times higher than those of Mac, as is typical for NT and more compacted soils. According to Tormena et al. (2002), Mac is restrictive to crops when below 0.10-0.15 m^3 m^{-3}, due to its role in soil aeration, but this limit cannot be seen as ultimate, but depends on the soil species and biological activity. In this study, the highest mean value observed for Mac was 0.09 m^3 m^{-3}, in Z1 (0.00-0.10 m layer), while the lowest was 0.06 m^3 m^{-3}, in the same soil layer in Z2.

A combined increase in BD and PR was found with increasing soil depth, followed by decreased TP at the expense of Mac. Z2 was the zone where the soil physical conditions were potentially more restrictive to crops. Although the BD values were below the critical limit proposed by Reinert & Reichert (2001), Mac remained below 0.10-0.15 m^3 m^{-3}, highlighting the poor distribution of pores in the different size classes. These results are in agreement with those of Amado et al. (2009), also in zones with contrasting yield, who confirmed that soil compaction and available water capacity are the most relevant soil physical properties for crop yield variation.

The variability of WIC data (Table 5) was moderate, but CV values were lower than those found by Miguel et al. (2009) in a Red Yellow Argisol (65-118 %), and by Scherpinski et al. (2010) in a dystrophic Red Latosol (91 %). The results are in agreement with the results in table 4, since the WIC in Z1 was almost 30 % higher than in Z2. This is probably related to the higher Mac in Z1, also almost 30 % higher than in Z2, in the mean of both soil layers.
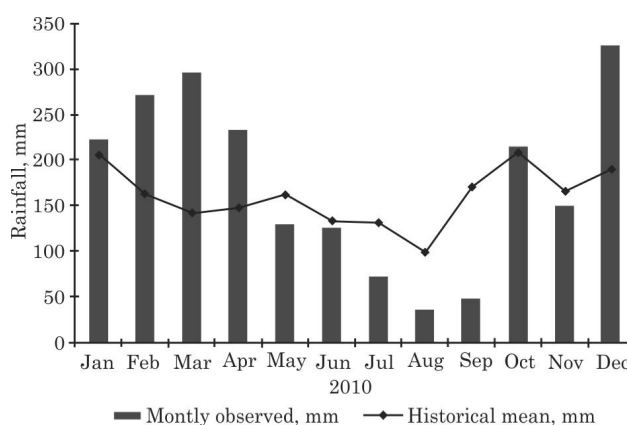

Figure 1. Historical mean of monthly rainfall (1976-2010) at the IAPAR weather station in Guarapuava, PR, and monthly observed (pluviometer) rainfall at the farm (Reserva do Iguaçu, PR) in 2010. Source: IAPAR (2011).

Table 5. Descriptive statistics and Student's *t*-tests for water infiltration capacity (WIC) of the soil in the zones of the plot selected for the study

Zone	Mean	Minimum	Maximum	SD[1]	CV[2]
		mm h^{-1}			%
Z1	56.11 A*	30.26	90.79	18.03	32
Z2	38.97 B	30.26	70.62	13.21	34
Plot mean	47.54	30.26	90.79	21.05	34

[1]SD: standard deviation; [2]CV: coefficient of variation. *Means followed by different letters indicate significant difference (p<0.05) by Student's t-test.

Although no direct correlation was observed between WIC and Mac, Pott & De Maria (2003), using a permeameter in different textured soils, detected a positive correlation of WIC with TP and negative correlation of WIC with BD. According to Silva et al. (2008), decreased Mac has a great effect on the water infiltration capacity and on root development, especially in clay soils. Santi (2007), on the other hand, working with a soil under NT, concluded that the water infiltration of the soil was one of the most relevant soil physical properties for crop yield.

The reasons for the more restrictive soil physical conditions to plants in Z2 than in Z1 are possibly the same reasons that resulted in higher Corg levels in Z1 in both soil layers (Table 3), despite the small distance between zones (about 1,500 m). Assuming a homogeneous field management (equal for Z1 and Z2), the lower Corg content in Z2 may have been the result of the lower residue production by crops in this zone, consistent with the lower crop yields, as shown by the yield maps (Tables 1 and 2). One reason is that Al^{3+}, analyzed *a posteriori*, was significantly higher in the subsurface of Z2 (0.79 $cmol_c$ dm^{-3}) than of Z1 (0.17 $cmol_c$ dm^{-3}), possibly restricting root growth and water and nutrient uptake, and consequently reducing plant growth in Z2 (unpublished data). In the NT management, lime is applied to the soil surface without incorporation, explaining why the difference in Al^{3+} between zones was observed in the surface layer only, since the conditions for lime dissolution and Al^{3+} neutralization in the deeper layer may have been less favorable in Z2, considering a soil with lower Corg, higher BD, lower porosity and lower water infiltration.

Conditions prior to soil use or details of the use itself may also have contributed to these differences. Despite the similarities between Z1 and Z2 in relation to slope and landscape position (shoulder position), Z2 is located in a more extreme position not only of the plot, but also of the farm, in a transition from prominent higher to lower positions, where the farm border is marked by eucalyptus. Therefore, although the soils were similar in the upper layers, there may be chemical, physical and/or hydrological differences in the deeper profile layers. Moreover, in addition to what has already been said, the border areas of large farms, farther away from the administrative headquarters, are negatively affected by operational problems, e.g., these areas are the last to be sown, so if all fertilizer is consumed before, fertilization will be incomplete, as similarly occurs with limestone or other inputs (seeds, pesticides etc).

A simple linear correlation matrix was established between the variables studied on the plot (Table 6). In the 0.00-0.10 m layer, except for Mic, all soil physical properties were significantly correlated to wheat GY. According to Callegari-Jacques (2003), the correlation was considered weak (r < 0.30) for GY x Mac (0.21*) and moderate (0.30 d″ r < 0.60) in all other cases: GY x PR (-0.33**), GY x SWC (0.30**), GY x BD (-0.32**), GY x TP (0.33**) and GY x WIC (0.31**). The results were similar in the 0.10-0.20 m layer: a weak

Table 6. Simple linear correlation matrix between variables studied in the plot, considering the evaluated soil layers

	GY[1]	Corg	PR	SWC	BD	TP	Mic	Mac	WIC
				0.00-0.10 m					
Corg	-0.14	1							
PR	-0.33**	-0.21*	1						
SWC	0.30**	0.13	-0.28**	1					
BD	-0.32**	0.28**	0.45**	-0.28**	1				
TP	0.33**	-0.26**	-0.44**	0.28**	-0.99**	1			
Mic	-0.10	0.07	0.13	-0.10	0.20	-0.19	1		
Mac	0.21*	-0.22*	-0.33**	0.21*	-0.52**	0.50**	-0.32**	1	
WIC	0.31**	0.05	-0.44**	0.09	-0.39**	0.39**	-0.03	0.28**	1
				0.10-0.20 m					
Corg	0.05	1							
PR	-0.34**	-0.09	1						
SWC	0.30**	0.02	-0.25*	1					
BD	-0.29**	0.05	0.60**	-0.21*	1				
TP	0.28**	-0.05	-0.59**	0.21*	-0.99**	1			
Mic	0.09	0.10	0.37**	0.15	-0.51**	0.51**	1		
Mac	0.20*	-0.16	-0.24*	0.06	-0.51**	0.51**	-0.48**	1	
WIC	0.33**	0.10	-0.45**	0.17	-0.65**	0.65**	0.38**	0.28**	1

[1]GY: wheat grain yield; Corg: organic carbon; PR: soil penetration resistance; SWC: soil water content; BD: soil density; TP: soil total porosity; Mic: soil microporosity; Mac: macroporosity; WIC: water infiltration capacity of the soil. ** and * indicate significance at 1 and 5 % probability, respectively.

correlation also for GY x Mac (0.20*), GY x BD (-0.29**) and GY x TP (0.28**), and moderate for GY x PR (-0.34**), GY x SWC (0.30**) and GY x WIC (0.33**), whereas the correlation of these last three soil physical properties in the 0.00-0.20 m layer with crop yield is noteworthy, in agreement with Imhoff et al. (2000), Hoad et al. (2001), Santi (2007) and Corrêa et al. (2009).

The results also show a moderate correlation for PR x BD (0.45** and 0.60**) and PR x WIC (-0.44** and -0.45**) in the 0.00-0.10 and 0.10-0.20 m layers, respectively, confirming that the variations in soil bulk density occur due to changes in soil structure, with a directly proportional influence on soil mechanic resistance and an inversely proportional influence on water infiltration into the soil.

CONCLUSIONS

1. The soil physical conditions were more limiting to crops in the zone of lower yield potential.

2. The wheat yield was negatively correlated with soil penetration resistance and bulk density, and positively correlated with the gravimetric water content in the field, total porosity, macroporosity and water infiltration capacity water of the soil.

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