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## Last millennium environmental changes and climate inferences in the Southeastern Atlantic forest, Brazil

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### ABSTRACT

This study presents paleoenvironmental data based on pollen, elemental and isotopic compositions of organic matter (TOC, N,  $\delta^{13}\text{C}$  and  $\delta^{15}\text{N}$ ) and  $^{14}\text{C}$  dating of 170 cm lake sediment record. Samplings have been made in Lagoa Grande at Parque Estadual Turístico do Alto Ribeira – PETAR, Southern São Paulo State, Southeastern Brazil. The variation in relative frequencies (in percentage) of arboreal pollen along the core range between 40 and 80%. The  $\delta^{13}\text{C}$  values ranged from  $-23\text{‰}$  to  $-30\text{‰}$  and C/N of  $\sim 10$  to 15, indicating the contribution of terrestrial  $\text{C}_3$  plants and algae in the sediment organic matter. The  $\delta^{15}\text{N}$  results presented values from 3 to 4.5‰, also suggesting a mixture of algae and terrestrial  $\text{C}_3$  plants. The  $^{14}\text{C}$  dating indicates modern age for the shallow horizons to  $\sim 1030$  BP at the base of the core. A probable wetter climate in the period of  $\sim 370$  BP to  $\sim 340$  BP was inferred from the data set, which corresponds to a part of the period covered by the Little Ice Age (LIA).

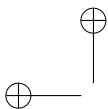
**Key words:** lacustrine sediment, C and N isotopes, pollen, Atlantic Forest, last millennium, Little Ice Age.

### INTRODUCTION

Lake systems are diverse, and the sources and alterations of organic matter are geographically and temporally variable. Nonetheless, useful generalizations can be made about the different kinds of elemental, isotopic and palynological proxies that provide evidence of the origins and depositional histories of sedimentary organic matter and hence paleoenvironmental conditions (Meyers 2003).

Pollen records obtained from lacustrine sediments have been used for paleoenvironmental reconstruction studies during the late Quaternary and the Holocene in

van der Hammen 1976, Absy et al. 1991, Roth and Absy 1993, van der Hammen and Absy 1994, Ledru et al. 1996, 2001, 2002, 2003, Ledru et al. 1995a, b, 1997a, b, 1998, 2002, Salgado-Labrière et al. 1998, Sifeddine et al. 2001, 2003, Saia 2006). Considering the last 3000 BP, most of the records indicate increased moisture conditions, although different climatic signals are recorded according to the site location. However, no data have been produced in high resolution emphasizing the potential use of an interdisciplinary study, involving different proxies and techniques for the last millennium.



general information. In fact, according to Meyers (2003), Sifeddine et al. 2004 and Ortiz et al. 2004, interpretation is not simple, in most cases, organic matter is a mixture of components from many sources and with variable degrees of preservation. Despite diagenesis, partial reworking and alteration of the original organic matter when sinking to the lake bottom, sedimentary organic matter retains important information on its origin, transportation and deposition, which helps paleoenvironmental reconstructions (Ortiz et al. 2004).

The  $\delta^{13}\text{C}$  values of lacustrine sedimentary organic matter may potentially reflect changing proportions of different types of vegetal matter (e.g.  $\text{C}_3$  vs.  $\text{C}_4$  plant abundance), as well as the presence of aquatic organisms, e.g. phytoplankton, which have similar isotopic signatures (Meyers 1997). During photosynthesis, phytoplankton preferentially utilizes the  $^{12}\text{C}$  isotope, leading to typical  $\delta^{13}\text{C}$  values of  $-20\text{‰}$  in its organic matter (Meyers 2003, Schidlowski et al. 1994). Thus, when the phytoplankton utilizes the dissolved  $\text{CO}_2$  in equilibrium with the atmosphere, it is usually an isotope indistinguishable from organic matter produced by terrestrial  $\text{C}_3$  plants. The correlation with C/N (total organic carbon and total organic nitrogen ratio) values is useful. C/N ratio of lacustrine sediments can be used to distinguish two main types of organic matter: (i) organic matter without cellulosic structure, originated from algae and phytoplankton, characterized by ratios between 4 and 10; and (ii) organic matter with cellulosic structure, produced by terrestrial plants, characterized by ratios  $\geq 20$  (Meyers and Ishiwatari 1993, Meyers 1994). This correlation should thus lead to discrimination between organic matter from freshwater algae and organic matter from  $\text{C}_3$  or  $\text{C}_4$  type plant sources. Analysis of stable carbon isotopes and the C/N ratio of organic matter preserved within lake sediments can provide important information about paleoenvironments, which can serve to complement pollen-based paleovegetation studies.

Although not widely used as a paleolimnological proxy as  $\delta^{13}\text{C}$  values, nitrogen isotopic compositions ( $\delta^{15}\text{N}$ ) can similarly help to identify sources of organic matter of lakes and to reconstruct past productivity rates

and from  $\text{C}_3$  terrestrial plants ( $+0.5\text{‰}$ ) (Peterson and Howarth 1987).

This study shows the first lake sediment record regarding the last millennium in Brazil. The main objective is the reconstruction of the vegetation changes in high resolution, with climate inferences that occurred in Southeastern Brazil during the last millennium. The data were obtained from a sediment record using pollen, geochemistry (total organic carbon – TOC and total nitrogen – TN) and isotope ( $\delta^{13}\text{C}$  and  $\delta^{15}\text{N}$ ) analysis, in addition to a detailed chronological frame based on radiocarbon dates. We did inferences with the period of the Little Ice Age (LIA) that occurred among 14<sup>th</sup> and 19<sup>th</sup> centuries and the Maunder Minimum (1645-1715 AD).

#### STUDY AREA

Lagoa Grande is a small lake located at PETAR (*Parque Estadual Turístico do Alto Ribeira*;  $24^\circ 31' 59.1''\text{S}$ ,  $48^\circ 39' 45.0''\text{W}$ , 364 m elevation), Vale do Ribeira, South of São Paulo State (Fig. 1). This park is a conservation unit on the left bank of the Ribeira River, approximately 350 km from the city of São Paulo, and 100 km from the Atlantic coastline. The lake developed within a centripetal drainage basin on karst topography. The catchment area is about 10 000  $\text{m}^2$  and the lake itself occupies an area of about 150  $\text{m}^2$ , with a present maximum water depth of 1 m. The lake, which is situated in a remote region, contains largely untouched natural sediments. Bedrock geology consists of a sequence of low-grade metasedimentary rocks of the Late Proterozoic Açungui Group, which is formed by fine-graded limestones interbedded mainly with metasiltites and phyllites (Campanha and Sadowski 1999). The major minerals in these rocks are calcite, quartz, mica and chlorite. The clayey soils are classified as cambisols.

The mean annual temperature is  $20^\circ\text{C}$ , with minimum and maxima of 14 and  $27^\circ\text{C}$ , respectively. The average annual precipitation is around 1600 mm (Karmann 1994), with low rainfall from June to August.

The natural vegetation around Lagoa Grande is the Atlantic Dense Ombrophilous Forest (forest hillside) and lies within the Atlantic Forest domain of vegeta-



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Fig. 1 – Location of the study area.

Melastomataceae, Meliaceae, Monimiaceae, Moraceae, Myristicaceae, Myrtaceae, Rubiaceae, Sapindaceae and Sapotaceae.

## MATERIALS AND METHODS

### SAMPLING AND ANALYTICAL PROCEDURES

A sediment core was collected using a vibracorer (Martin et al. 1995) and a 3-inch diameter aluminum tube with 3 m length. The 170 cm long core was collected close to the lake margin and sub-sampled in 2 to 10 cm intervals for analysis.

### C AND N ISOTOPES

The sediment samples separated of leaves, roots, etc., were treated with HCl 4% to eliminate carbonates, washed with distilled water up to pH 6, dried at 50°C and homogenized. These samples were used for total organic carbon and nitrogen analyses, carried out at the

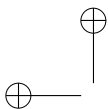
0.09 and 0.07%, respectively. The  $^{13}\text{C}$  and  $^{15}\text{N}$  were expressed as  $\delta^{13}\text{C}$  and  $\delta^{15}\text{N}$  with respect PDB standard and atmospheric air, respectively the conventional notations:

$$\delta^{13}\text{C}_{\text{Sample}} = \frac{R_{\text{S1}} - R_{\text{PDB}}}{R_{\text{PDB}}} \times 1000$$

$$\delta^{15}\text{N}_{\text{Sample}} = \frac{R_{\text{S2}} - R_{\text{Air}}}{R_{\text{Air}}} \times 1000$$

where  $R_{\text{S1}}$  and  $R_{\text{S2}}$  are, respectively, the  $^{13}\text{C}/^{12}\text{C}$  and  $^{15}\text{N}/^{14}\text{N}$  ratios in the sample,  $R_{\text{PDB}}$  the  $^{13}\text{C}/^{12}\text{C}$  ratio of the international standard (PDB) and  $R_{\text{Air}}$  the  $^{15}\text{N}/^{14}\text{N}$  ratio for the atmospheric air. The results were expressed in delta *per mil* notation, with analytical precision better than 0.2‰ (Pessenda et al. 2004, 2005).

For radiocarbon analyses, wood fragments and bulk sediment samples were subjected to physical treatment (removal of vegetable remains, etc.) and then treated with HCl 4% for 4 h at 60°C to remove carbonates.



sent to the Isotracer Laboratory, University of Toronto, Canada, for Accelerator Mass Spectrometry (AMS). Radiocarbon ages (Table II) are expressed as  $^{14}\text{C}$  BP (years Before Present) and in cal BP (calibrated years Before Present) (Stuiver et al. 1998) and with precision of  $\pm 2\sigma$ .

#### PALINOLOGY

The pollen analyses were performed each at 10 cm interval along the 160 cm long core. Chemical treatment of approximately 5 g of sediment was performed for pollen extraction based on Faegri and Iversen (1989). Sample treatment followed the Lüber technique using 70% HF for 18 h, followed by 50% HCl in hot water bath and by 10% KOH solution. Palynomorphs were separated using  $\text{ZnCl}_2$  and the residue fraction of 50  $\mu\text{L}$  mounted on microscope slides in silicon oil; pollen counting was performed under  $1000\times$  magnification. Approximately 300 pollen grains of tree and herb taxa were counted *per* sample. The reference collection of Dr. Marie-Pierre Ledru and pollen atlas were used for pollen identification (Absy 1975, Markgraf and D'Antoni 1978, Colinvaux et al. 1999). Frequencies are expressed as percentages of the total sum of arboreal pollen (AP), non-arboreal pollen (NAP) and undetermined types, with aquatic taxa, fern spores and Cyperaceae excluded from the total sum. Relative frequencies of spores and aquatic taxa were calculated in relation to the total AP and NAP sums. The determination of pollen concentration was based on the method of Cour (1974), following the mathematical equation:

$$f(\%) = \frac{v}{V} \times \frac{I}{L} \times 100$$

where:

- $v$  = volume of the residue used in the feature of pollen slides;
- $V$  = total volume of the residue after a chemical treatment;
- $I$  = total length observed in the slides during the counting of grains;

The residue percentage in each sample is obtained from this calculation. From the fraction observed in each residue ( $f\%$ ), it is possible to calculate the grain concentration *per* gram of sediment. This calculation is obtained from the next equation:

$$r = \frac{n \times (100/f)}{m}$$

where:

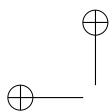
- $n$  = total sum of grains in one slide;
- $f$  = fraction observed in the residue;
- $m$  = original weight used in the chemical treatment of the sample.

#### RESULTS

The core sediment consists of grey dark clay with massive structure at the bottom, changing upward to plane-parallel lamination of brown and grey clay mud, to light brown mud with massive structure and red spots (iron/hydroxide/oxide), to plane-parallel lamination of brown and grey clay mud, to light brown mud with massive structure and red spots, and to light brown and massive structure without bioturbation at the surface (Fig. 2, Table I).

The ages of terrestrial macrofossils (wood fragments) and sediment samples ranged from  $1030 \pm 60$  BP at 171-168 cm to modern ages (8-6 cm sediment depth), with increasing ages downward (Fig. 2, Table II). At 126-116 cm a wood fragment presented an age of  $680 \pm 80$  BP, and at 124-122 cm the organic sediment sample an age of  $550 \pm 50$  BP. Similar results were obtained for the wood fragment ( $250 \pm 60$  BP) at 65-63 cm, and for the organic sediment sample ( $320 \pm 60$  BP) at 64-62 cm. In order to provide age estimation relative to sediment depth, a continuous age scale based on interpolation of  $^{14}\text{C}$  ages *versus* sediment depth is presented in Figure 3.

The results of the pollen analysis are presented in Figures 2 and 4. Figure 2 shows changes in arboreal pollen and spores frequencies and the concentration values of number of pollen grains *per* gram of wet sediment



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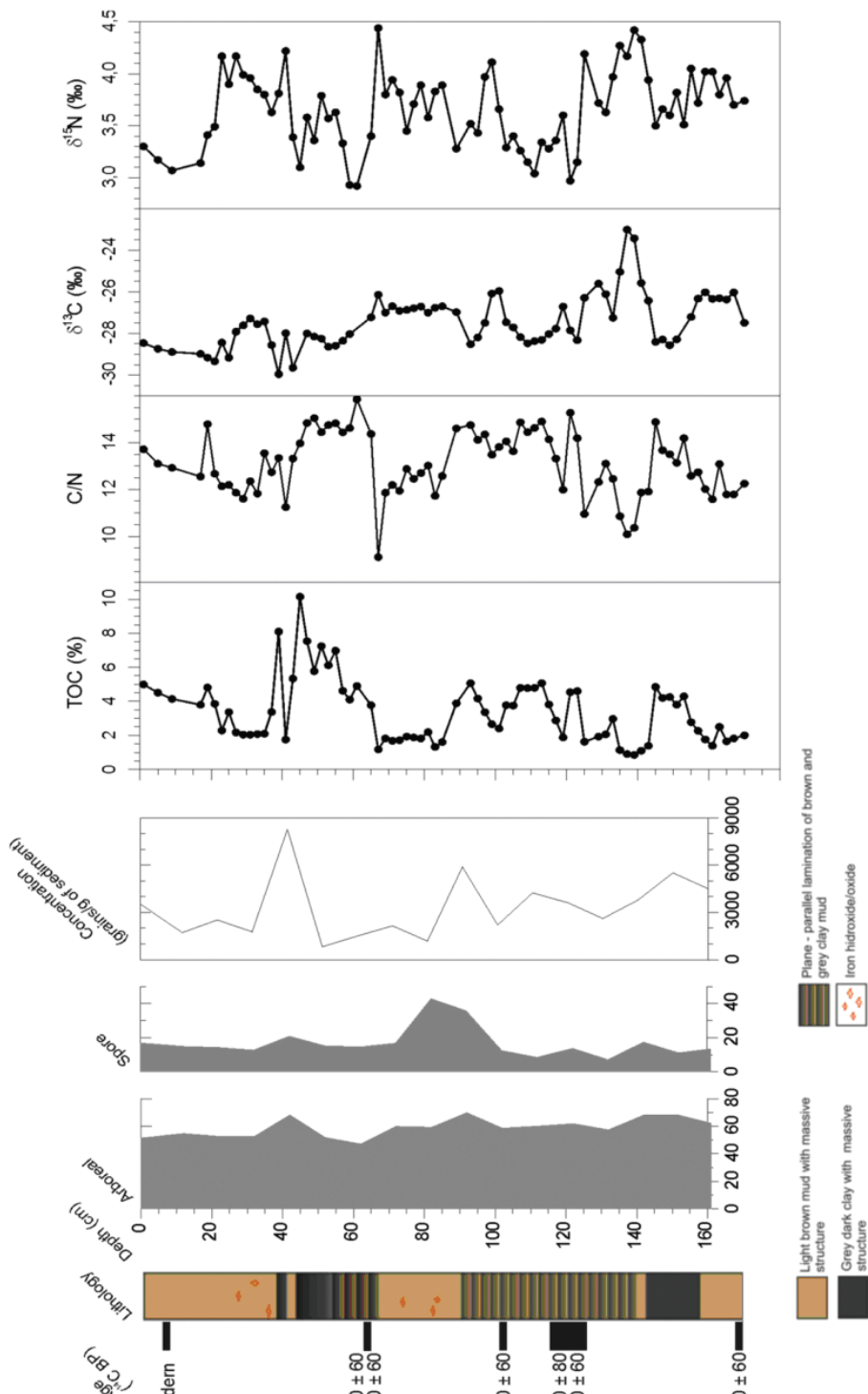




TABLE I  
Stratigraphic description of the Lagoa Grande core.

Depth (cm)	Description
0-25	Light brown mud with massive structure without bioturbation
25-41	light brown mud with massive structure; presence of iron/hydroxide/oxide
41-67	plane – parallel lamination of brown and grey clay mud
67-88	light brown mud with massive structure; presence of iron/hydroxide/oxide
88-137	plane – parallel lamination of brown and grey clay mud
137-170	grey dark clay with massive structure

TABLE II  
<sup>14</sup>C ages of organic matter obtained in the Lagoa Grande core.

Sample	Depth (cm)	Laboratory number	<sup>14</sup> C Age (BP)	<sup>14</sup> C Age (cal BP), 2σ
Organic matter	6-8	TO-11372	Modern	1958-1959*
Organic matter	62-64	TO-11373	320 ± 60	1448-1665
Wood fragment	63-65	TO-10874	250 ± 60	1468-1695
Organic matter	100-102	TO-11374	370 ± 60	1440-1643
Wood fragment	116-126	TO-10875	680 ± 80	1209-1424
Organic matter	122-124	TO-11375	550 ± 60	1296-1443
Organic matter	168-171	TO-11376	1030 ± 60	890-1155

TO – Isotrace Laboratory, Toronto, Canada; (\*) – cal AD.

*Cecropia*, Myrtaceae and Moraceae, and *Podocarpus* and *Weinmannia*, which are elements of Atlantic Om-brophilous forest and of forest of altitude, respectively.

The predominance of arboreal taxa is observed along all the sediment core (Fig. 2). From 170-140 cm depth (~1000 BP to ~750 BP), arboreal pollen frequencies are high, with ~70% and a peak of fern spores of ~20% is being observed. *Cecropia*, Moraceae, Myrtaceae and Poaceae are also recorded. The frequency of *Weinmannia* and *Alchornea* is high (Fig. 4). *Weinmannia* shows the highest frequency of this interval, with 7%, and *Podocarpus* appears around 150 cm. In the interval 155 cm to 145 cm (~950 to 800 BP), the TOC content increased from 1.8 to 5.0%, the C/N ranged from 12.6 to 14.9, the  $\delta^{13}\text{C}$  values stayed around  $-28.5\text{‰}$  and  $\delta^{15}\text{N}$  values ranged from 3.3‰ to 3.8‰ (Fig. 2). The results are indicative of the predominance of organic matter from terrestrial C<sub>3</sub> plants. From 145 cm to 135 cm (~800 BP to 720 BP), the TOC values decreased from

creased from 3.5‰ to 4.5‰, which is indicative of more significant presence of phytoplankton (algae) (Meyers 1994). This could be associated with the increase of the lake level due to more humid conditions than the previous period. From 135 cm to 70 cm depth (~720 BP to ~320 BP), the elemental and isotopic analyses (TOC ranged from 0.8 to 5.1%, the C/N ratio from 10.1 to 15.3,  $\delta^{13}\text{C}$  values from  $-23.7\text{‰}$  to  $-29\text{‰}$ , and  $\delta^{15}\text{N}$  from 4.25‰ to 3‰) indicated a mixture of organic matter (C<sub>3</sub> plants and phytoplankton), with higher influence of terrestrial C<sub>3</sub> plants. From 135 cm to 80 cm (~720 BP to ~350 BP), a decrease of *Alchornea* (~2%) and *Cecropia* (~0.5%) was observed, and the *Podocarpus* disappeared till ~110 cm (~420 BP). From 80 to 70 cm (~350 BP to ~320 BP), the spore frequencies presented the highest value (up to ~50%) (Fig. 4), which is coincident with the significant increase of *Weinmannia* (4%), *Podocarpus* (0.8%), Cyperaceae (~4 to 5%) and Myrtaceae (28%) at 90 cm (~360 BP). From 70 to 50 cm



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depleted  $\delta^{13}\text{C}$  ( $-29\text{‰}$ ), and  $\delta^{15}\text{N}$  ranged between  $\sim 3\text{‰}$  and  $3.5\text{‰}$ , indicating a mixture of organic matter ( $\text{C}_3$  plant and algae). A decrease of arboreal pollen frequency from  $\sim 60\%$  to  $55\%$  was observed at  $\sim 60$  cm ( $\sim 300$  BP), simultaneously with *Alchornea* (4%), Myrtaceae (10%), *Podocarpus* (0.25%) and *Weinmannia* (2%). The spore frequency was around 10% (Fig. 4).

From 50 cm to top ( $\sim 280$  BP to the present) the TOC contents ranged from 10 to 1.7%, C/N values from 15 to 11,  $\delta^{13}\text{C}$  ( $-27$  to  $-30\text{‰}$ ), and  $\delta^{15}\text{N}$  values from 3 to  $4.25\text{‰}$  (Fig. 2), suggesting a mixture of organic matter originated from  $\text{C}_3$  terrestrial plants and phytoplankton (Meyers 1994). From 50 to 25 cm ( $\sim 280$  BP to  $\sim 180$  BP), an increase was observed in the arboreal pollen (up to 70%), spores (25%), *Alchornea* (10%), Poaceae (20%), Moraceae (8%) and in the grains/g of sediment. Simultaneously, this increase is accompanied by a decrease in the frequencies of *Cecropia* (0.5%), *Podocarpus* (0.25%) and *Weinmannia* (1%). From  $\sim 25$  cm to top ( $\sim 180$  BP to the present), the total arboreal pollen frequencies are around 60% and the spore frequencies of  $\sim 15\%$  at 15 cm. *Alchornea* (from 4 to 1%) and *Weinmannia* (1%) remained low, Myrtaceae (10%) and Poaceae (20%) remained constant, and *Podocarpus* (up to 1.5%), *Cecropia* (up to 2.5%) and Moraceae (up to 11%) increased.

The distribution of C/N versus  $\delta^{13}\text{C}$  values (Fig. 5) indicates the presence of two sources of organic matter. Very few samples indicate a more significant presence of algae and most of the samples represented a mixture of algae and terrestrial  $\text{C}_3$  plants, probably associated with the variation of the lake level.

#### DISCUSSION

The pollen record of Lagoa Grande indicates no drastic changes in the arboreal vegetation, suggesting that the climate of the area, at least in the last millennium, remained humid in most of the period, probably without longer dry seasons. The pollen types found from  $\sim 1000$  BP to  $\sim 400$  BP are characteristic of dense and humid forest represented by *Alchorneae*, *Cecropia*, *Myrtaceae* and *Moraceae*, of local plants on the lake margins such

at  $\sim 820$  BP and 380 BP, and more significant frequencies ( $\sim 4\%$  and  $\sim 6\%$ , respectively) of *Weinmannia*.

An increase in the frequency of spores up to 25% is observed at  $\sim 350$  BP. One hypothesis could be associated with the oxidation and degradation of more fragile pollen types, resulting in a relative enrichment of more resistant pollens, such as spores (Colinvaux et al. 1999). However, oxidation signs or bad preservation were not observed in the pollen content. The second hypothesis is associated with the decrease of the water level, probably related with a warmer and/or drier climate and with consequent wider exposure of the lake margins. Fern spores and Cyperaceae, usually found in lake margins, simultaneously increased their frequencies. Pollens were recorded between  $\sim 400$  and 320 BP (Fig. 4). A third hypothesis is also possible: an increase in humidity and erosion process, with the consequent input of fern spores and Cyperaceae, usually found in terrestrial environments, into the lake basin. Geochemical interpretation presented by Oliveira et al. 2009 supports the last hypothesis. According to these authors, the composition of the Lagoa Grande sediments can be described in terms of a dominant geogenic component, association of mica, kaolinite and goethite with elements of low mobility during weathering (Ba, Cs, Rb, Sr, Th, U, Zr, Hf, Ta, Cr, Sc, Zr, Hf and Ta). This component may have been brought by erosion and removal of deeper soils from the catchment area, where these elements are found at the highest abundance and largely derived from the B-horizon of the surrounding soil profiles. In addition, there is a minor biogenic component and related to the O-horizon (shallow layers) of these soils, which is characterized by the association of chlorite with the transition metals (Fe, Zn and Cu). This probably represents the input of organic lake material eroded from the uppermost organic horizons of the catchment soils. In the O-horizon of these soils, many metals are found forming stable organo-metallic compounds. This is particularly the case for Zn and Cu, which are among the most essential elements for plant growth and is accumulated in the humus layer (Bertoni et al. 2000). This component is relatively more



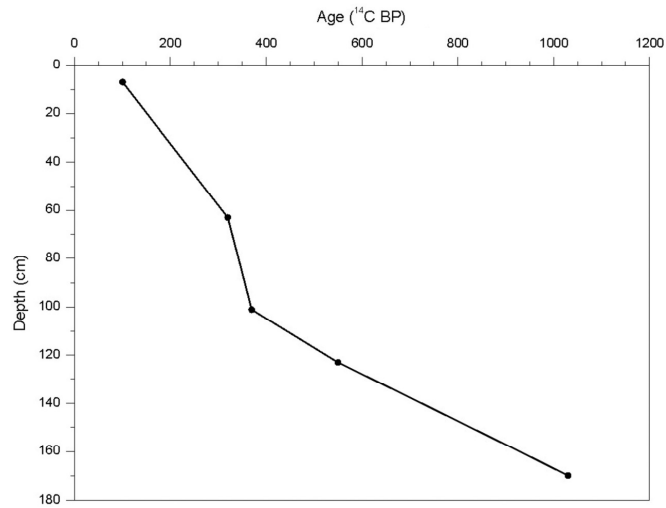


Fig. 3 – Age in relation to sediment depth.

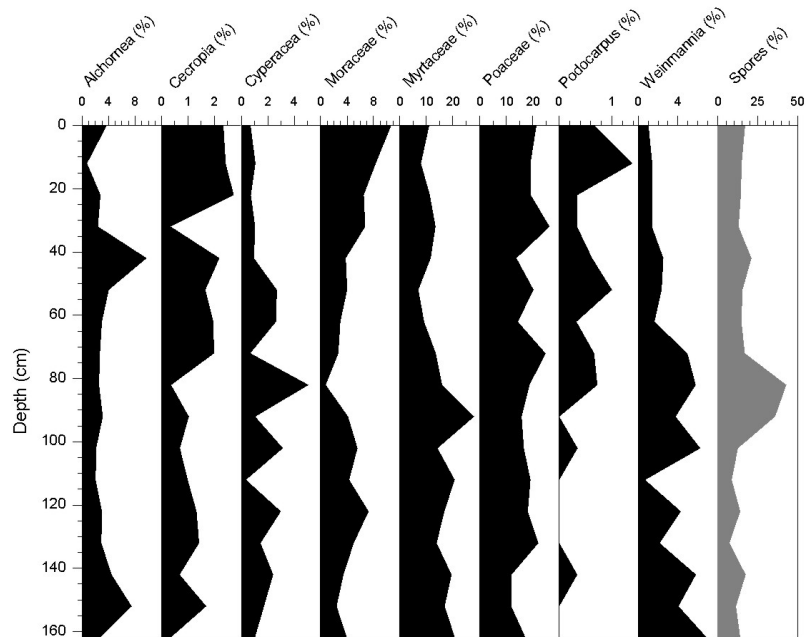
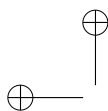


Fig. 4 – Percentage of *Alchornea*, *Cecropia*, *Cyperaceae*, *Moraceae*, *Myrtaceae*, *Poaceae*, *Podocarpus*, *Weinmannia* and spores from Lagoa Grande.

ble cause for this is the variation in the erosion regime in the catchment area. At the onset of the sedimentation (1030 to 730 BP), the input of deeper layers of soil was reduced as a result of the relative catchment stability.

deeper layers of the soil profile were successively eroded. The inflow of those materials into the lake led to a relative depletion of the biogenic component of the sediments. This trend was enhanced between 490 to 360 BP.



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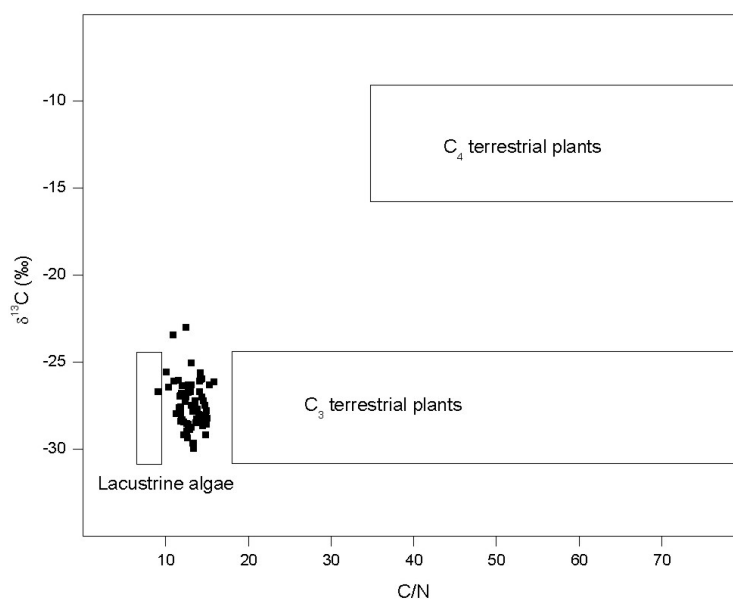


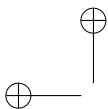
Fig. 5 –  $C/N \times \delta^{13}C$ .

ment area, except for brief episodes of enhanced erosion around 340 BP, 280 BP and 200 to 170 BP.

The changes in the erosion regime can be assumed to be controlled by climate fluctuations. A shift toward increased rainfall, particularly if it occurs in the summer months increasing the seasonality and enhancing erosion in the catchment area. This would increase the supply of geogenic material into the lake, diluting the material from the more organic topsoil.

A link can thus be established between the geochemical signature of the sediments and the climatic conditions at the time of their deposition. During the last millennium, in the Lagoa Grande region, the period inferred from the geochemical data from 730 to 360 BP (wetter), with the wettest conditions between 490 and 360 BP, overlaps approximately that one suggested by pollen and isotopic data (see the following discussion) as a warmer (inference from pollen) and wetter (inferences from both indicators) period of ~370 to 340 BP. Therefore, the hypothesis of an increase in both humidity and erosion process, with the consequent input of fern spores and Cyperaceae, will be considered. Concerning the elemental and isotope data, it was observed

tive of significant input of organic matter of  $C_3$  probably associated with the increase of spores of Cyperaceae and arboreal species as Myrtaceae and *Weinmannia*, during more humid and erosive period (Fig. 4). In the interval ~340 to 320 BP, it was observed the lowest values of TOC in the whole sediment profile, as the lowest grain *per* gram of sediment, which could be connected with a small input of the organic matter in the sedimentary basin, possibly associated with the humid and erosive period. At the same time, a significant increase in *Weinmannia* frequencies is observed (Fig. 4). A pollen record in Cambará do Sul (29° 50'06"04" Southern Brazil) shows that *Weinmannia* became a common taxa in the Araucaria forest between 1520 and 1770 AD (~430 BP and 180 BP) (Beltrame et al. 2004), suggesting a shift to warmer climatic conditions. These authors associated this interval with a cooler phase (~550 BP to 200 BP) called Little Ice Age (LIA), one of the most important variations of the last millennium (from the 14<sup>th</sup> to the 19<sup>th</sup> centuries). This is known from North Hemisphere records in the last millennium, described by Pederson et al. 2005, and in the last century by Fontana, 1976. Font 1988, Ramil-Rego et al.



to the highest spore frequency and TOC) for the interval 370-340 BP. This can be connected with the LIA period, which has been poorly recorded in South America. Lichenometry dating of glaciers in the Cordilleras in Bolivia, ( $16^{\circ}17'S$ ,  $68^{\circ}06'W$ ; Rabatel et al. 2005) established a chronology for the LIA in that area, with a maximum occurrence in the second half of the 17<sup>th</sup> century ( $\sim 300$  BP) and finishing between 1870 and 1910 AD ( $\sim 80$  to 40 BP). These authors concluded that the second part of the LIA could have been dry in the Andes, because tropical glaciers retreat significantly when precipitation decreases.

Many authors had identified the LIA in the tropical Andes, especially in the Cordilleras of Peru (Clapperton 1983, Rodbell 1992, Solomina et al. 2007, 2008), Argentina (Iriondo and García 1993, Piovano et al. 2002), Uruguai (Del Puerto et al. 2006), Equator (Hastenrath 1981) and Bolivia (Thompson et al. 1986, Rabatel et al. 2005). The archives for Brazil record the LIA period associated to warm climatic inferences in South region (Behling et al. 2004).

In Southern Peru, a dry climate seems to have prevailed in the 1720-1880 AD ( $\sim 230$  to 70 BP) period (Thompson et al. 1986, Rabatel et al. 2005). In Argentina, close to Buenos Aires, a cool episode was observed (Iriondo and García 1993) and associated with the LIA period, in which an aridity episode was reported. In the same country, Laguna Mar Chiquita also shows dry conditions during the LIA period (Piovano et al. 2002).

These drier events could be correlated with the Maunder Minimum (1645-1715 AD;  $\sim 305$  to 235 BP), described as a climax of the LIA, when the earth gradually warms until the warming is further accelerated by anthropogenic greenhouse gases (Mann et al. 1999, Shindell et al. 2001, Youshimori et al. 2005), mainly considering an error (precision) of  $\pm 50$  years for each radiocarbon dating in the 400-320 BP, inferred warmer and wetter climate period in this work.

#### CONCLUSIONS

Environmental conditions were relatively stable with the maintenance of the forest during the last 1000 years

the period. However, the bulk and isotopic analyses of sediment core indicate significant variations, probably associated with changes in the Lagoa Grande water column, eventually linked with a wet and warm period (from  $\sim 370$  BP to  $\sim 340$  BP), that was associated with the Little Ice Age (LIA) and Maunder Minimum.

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#### RESUMO

Este estudo apresenta dados paleoambientais baseados em análises de pólen, da composição elementar e isotópica da matéria orgânica (TOC, N,  $\delta^{13}C$  e  $\delta^{15}N$ ) e datação  $^{14}C$ , de um testemunho sedimentar lacustre de 170 cm. A amostragem foi realizada na Lagoa Grande localizada no Parque Estadual Turístico do Alto Ribeira – PETAR, Vale do Ribeira, sul do estado de São Paulo, sudeste do Brasil. As variações nas frequências relativas de pólen arbóreo (em porcentagem) apresentaram-se entre 40 e 80% ao longo de todo o testemunho. Os valores de  $\delta^{13}C$  variaram de  $-23\text{‰}$  à  $-30\text{‰}$ , indicando a contribuição de plantas  $C_3$  (terrestres) e algas na matéria orgânica sedimentar. Os resultados de  $\delta^{15}N$  apresentaram valores entre 3 e  $4,5\text{‰}$ , também sugerindo uma mistura de algas e plantas  $C_3$ . A datação  $^{14}C$  indica idade moderna para as camadas superficiais do testemunho e uma idade de até 1030 anos AP para a camada mais profunda do mesmo. Um provável clima mais úmido para o período de  $\sim 370$  anos AP até  $\sim 340$  anos AP foi inferido a partir dos resultados apresentados, correspondendo a uma parte do período do *Little Ice Age* (LIA – Pequena Idade



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