

Diálogo Andino
Revista de Historia, Geografía y Cultura Andina

Diálogo Andino - Revista de Historia,
Geografía y Cultura Andina

ISSN: 0716-2278

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Diálogo Andino - Revista de Historia, Geografía y Cultura Andina, núm. 41, junio, 2013,
pp. 101-132

Universidad de Tarapacá
Arica, Chile

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ANCIENT EL NIÑO EVENTS, HUMAN ADAPTATION, AND ECOLOGICAL TRANSFORMATIONS: EARLY FORMATIVE PERIOD (2400-1450 B.C.) OCCUPATIONS IN SOUTHERN COASTAL ECUADOR

ANTIGUOS EVENTOS DE EL NIÑO, ADAPTACIÓN HUMANA Y TRANSFORMACIONES ECOLÓGICAS: PERÍODO FORMATIVO TEMPRANO (2400-1450 AC) OCUPACIONES EN LA COSTA SUR DE ECUADOR

John E. Staller*

El Niño-Southern Oscillation is a warming of surface sea temperatures in the eastern Pacific Ocean. Such climatic and oceanographic perturbations have dramatic impacts upon human adaptation and sociocultural development. Evidence multidisciplinary Artificial a mound of dirt in the ceremonial center of Valdivia The Emerancia have documented the abandonment of the site in relation to the El Niño phenomenon. The initial site abandonment was in response to intense or mega event dated to 2150 BC, associated with the formation of the beach, singing fossil reoccupation C14 and dated ca 2200-1450 BC and final abandonment dated to 1450 B.C. Final abandonment is associated with an earthquake and a short-lived reoccupation. Data from excavation, regional settlement patterns and shellfish frequencies are presented to determine whether repeated and final site abandonment was related to El Niño. Results indicate widespread environmental degradation and geomorphological changes to the surrounding coastline were related to El Niño, and that it was clearly a factor to sociocultural development and adaptive responses. These data explore chronology, assess the intensities, and measure the effects of ancient El Niño events upon pre-Hispanic occupations this ceremonial center and pre-Hispanic occupations along the Arenillas River valley, El Oro Province, Ecuador.

Key words: El Niño, Andes, coastal geology, geomorphology, extinctions.

El Niño-Oscilación del Sur es un calentamiento de las temperaturas de la superficie del mar en el océano Pacífico oriental. Tales perturbaciones climáticas y oceanográficas tienen un impacto dramático sobre la adaptación humana y el desarrollo sociocultural. Evidencias multidisciplinarias de un montículo de tierra artificial en el centro ceremonial La Emerancia de Valdivia han documentado que el abandono inicial del sitio fue en respuesta a un intenso megaevento fechado ca 2150 a.C., asociado con la formación de la playa, canto fósil, C14 y reocupación de fecha ca 2200-1450 a.C. y con abandono final fechado ca 1450 a.C. El abandono final está asociado con un terremoto y una reocupación de corta duración. Los datos de la excavación, los patrones de asentamiento regionales y las frecuencias de los mariscos se presentan para determinar si repetida y último abandono del sitio se relaciona con El Niño. Los resultados indican que la degradación generalizada del medio ambiente y de los cambios geomorfológicos en la costa de los alrededores estaban relacionados con El Niño, y que era claramente un factor de desarrollo sociocultural y las respuestas de adaptación. Estos datos exploran cronología, evaluar las intensidades, y medir los efectos de los antiguos eventos de El Niño sobre las ocupaciones prehispánicas de este centro ceremonial y ocupaciones prehispánicas a lo largo del valle del río Arenillas, provincia de El Oro, Ecuador.

Palabras claves: El Niño, los Andes, geología costera, geomorfología, extinciones.

Introduction

El Niño/Southern Oscillation is related to changes in oceanic currents and trade winds, characterized by a general warming of surface sea temperatures along the Eastern Pacific, and a lessening or reversing of northeast trade winds, bringing warm humid air over coastal South America (Philander 1998:108; Fagan 1999:119-138). These climatic

and oceanographic alterations create a reduction of upwelling cold waters along the west coast of South America. These climatic changes result in dramatic perturbations to maritime and terrestrial flora and fauna and, consequently, human adaptation. El Niño phenomenon in its current manifestations has its origins around c. 5800 years ago (Sandweiss 1986, 2003; Sandweiss *et al.* 2009; Andrus *et al.* 2008). When El Niño is particularly extreme or intense, as

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the 1983/84, or 1997/98 events, they are referred to as Mega El Niño, which are broadly differentiated by intensity and duration, or a combination of both (Moseley 1987:7; Moy *et al.* 2002). The following study presents multidisciplinary evidence consisting of both regional survey and excavations, as well as radiocarbon and AMS dates, geomorphological evidence and statistical analysis of minimum number of individuals (MNI) of marine shell excavated from natural stratigraphic layers as well as arbitrary increments, to determine the approximate antiquity and duration of El Niño events. These data address their possible relationship to widespread changes in human adaptation and the natural landscape and geomorphology between c. 2200 and 1450 B.C.

These multiple lines of evidence are from archaeological research in southern El Oro Province, Ecuador (Figure 1). Excavation and regional survey uncovered indications of regional perturbations, involving the ecology, geology, geomorphology, and regional settlement patterning related to El Niño/Southern Oscillation (Staller 1994:131-153, 211-232, 335-347)¹. Shell counts of minimum number of individuals (MNI) provide a basis for assessing the times of occurrence, intensity, and duration, as well as how such climatic events effected human adaptation (see also Staller 1992/93, 1996, 2000, 2001b). There is evidence for a general trend of increased frequency of El Niño events that in some instances changed the climate and coastal habitats permanently. These El Niño-induced alterations required major adaptive changes and short-term increased dependence upon certain seasonally specific resources and suggest the long-term cultural response to such ecological and geomorphological transformations favored flexibility or increased diet breadth rather than specialization and/or dependence upon particular resources (Staller 1994, 2001a, 2001b; see Binford 1989, 2001).

Ocean Currents, Trade Winds, Sea Surface Temperatures and Climate

Oceanographic research has indicated that the El Niño phenomenon is related in part to changes in oceanic trade winds and the maintenance of a westward rising water gradient across the eastern Pacific (Wrtki 1979, 1982; Wrtki *et al.* 1976; Philander 1989; McPhaden and Picaut 1990). During a normal dry season, the trade winds create an upwelling or an upward flow of cold deep water along the eastern equatorial Pacific. During an El Niño, trade winds weaken or reverse and cold sea surface temperatures (SST) are replaced by warm surface water along the eastern Pacific (McPhaden 1999:950-951; Moseley *et al.* 1981). Warmer SST is initially apparent on offshore areas in the Santa Elena Peninsula (Rasmussen 1985:168; McPhaden 1999:figure 3; McPhaden and Picaut 1990).

Cold SSTs are ordinarily accompanied by a strong, cool breeze that slightly lowers the air temperatures for most of the year producing moisture in the form of a dense fog or “*garúa*” (Thayer and Barber 1984:6). Cloud forests in the lower slopes of the Andes trap and recycle huge quantities of moisture in the form of clouds moving in off the ocean, even during the driest months (Parker and Carr 1992:17). The moisture generated by cloud forests maintains a high water table and year-round flow of the coastal streams, which have obvious implications for long-term human adaptation and population density. In coastal Peru, the *garúa* nourishes lush patches of vegetation, or “fog meadows”, called *lomas* formations, which are greenest between April and December (Lanning 1965:68). Such islands of vegetation in a sea of desert sand were important to early human adaptation and ecological diversity, and their geographic distribution is greatly dependent upon El Niño cycles (Lanning 1965:70-72). Although *lomas* have yet to be identified in coastal Ecuador, Edwin Ferdon (1950:61) mentions seeing, “the existence of isolated continually moist, hill top forest cover” near Manta in northern Manabí Province. If *lomas* existed further to the south of Manta, they would certainly have been modified by habitations or agricultural practices.

One of the primary factors affecting climate and the onset of El Niño is a change in the prevailing trade winds and resultant SST (McPhaden 1999:950; Vuille *et al.* 2000). Trade winds along the eastern Pacific blow from the southwest to the east

¹ Excavations at La Emerenciana and regional survey and excavations along the lower Arenillas River were under the direction of the author over a period of twenty-one months. Archaeological research was under the auspices of the Museo Antropológico in Guayaquil, Ecuador and Department of Anthropology, Southern Methodist University, Dallas, Texas. The project was fully funded by a Fulbright research scholarship in 1988 and research grant in 1989, under the auspices of the Fulbright Association and the Institute for International Education (IIE) in Washington DC (Staller 1994).

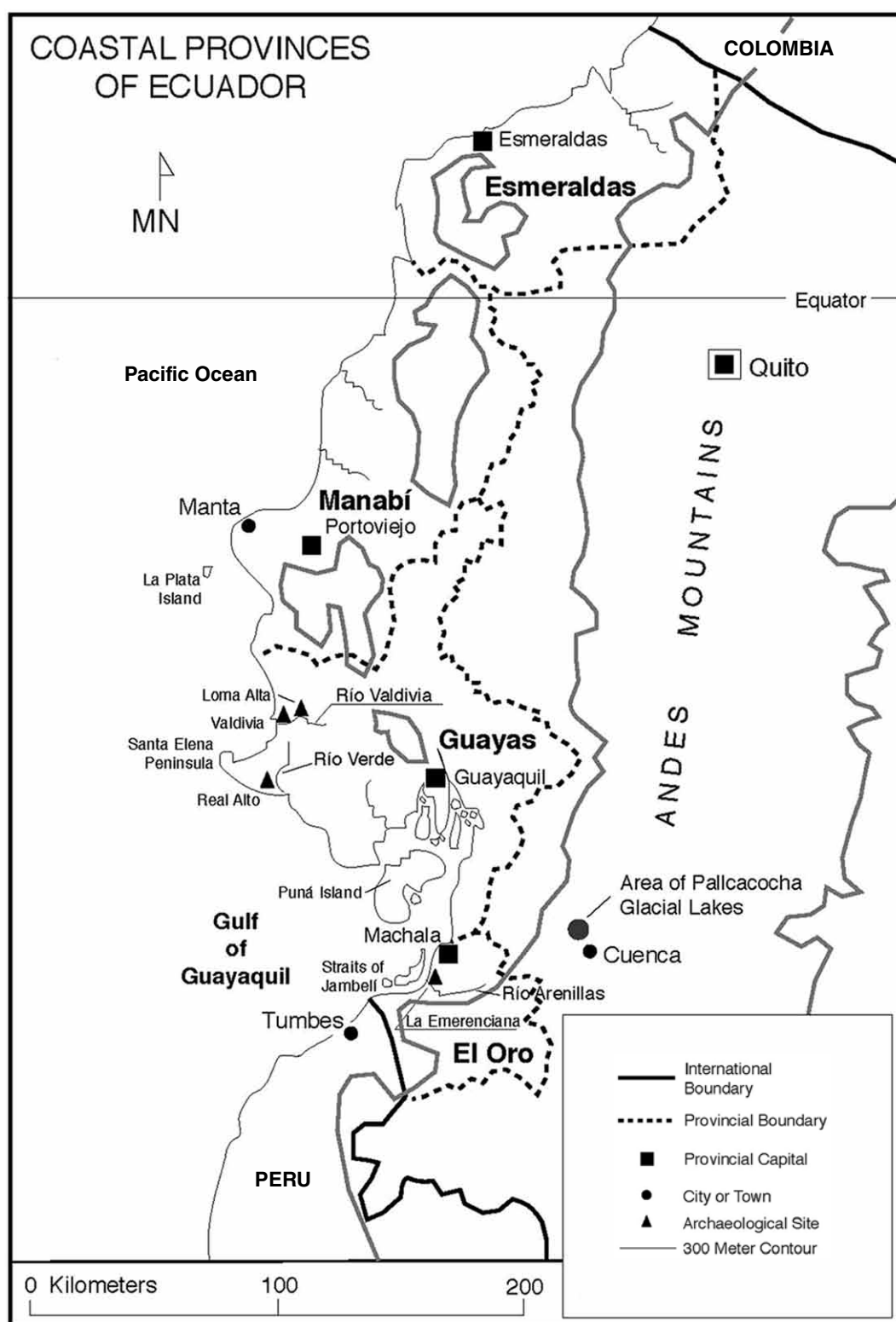


Figure 1. Coastal provinces of Ecuador showing the Gulf of Guayaquil, Straits of Jambell and Valdivia sites, modern cities, as well as towns mentioned in the text.

during the dry season, and during El Niño weaken or falter due to decreased variations in land and sea temperatures (Ferdon 1950:37; Wrtki *et al.* 1976; McPhaden 1999). When trade winds subside, the ocean responds with dramatic temperature increases, setting altered currents into motion (Thayer and Barber 1984:4; Vuille *et al.* 2000). A rise in sea level sets the stage for high tides that can produce large-scale erosion (Thayer and Barber 1984:4; Cane 1983:1190-1191; Moreano 1984; Wrtki 1982).

Increases in precipitation can produce fluvial erosion, causing significant increases in sediment load in the stream channels resulting in large-scale modification of the coastline and sudden and dramatic increases in mean water level for periods of several weeks or even months (Arntz 1986; Craig and Shimada 1986; Wells 1987; McPhaden 1999). In low energy tidal-dominated settings such conditions can produce flash floods and increased river discharges that may radically alter the landscape (Thayer and Barber 1984:4; Cane 1983:1190-1191; Caviedes 1975:Figures 2, 4; Wells 2001:150; Keefer *et al.* 2003; Sandweiss *et al.* 2007).

El Niño related sea level rise occurs during high “spring tides” (*marejadas maretazos*), and accompanying waves often result in mass destruction along the littoral (Arntz 1986:7). Under certain conditions, surge waves can completely engulf the back barrier and severely erode island barriers and shallow reefs. As the storm passes or the tide ebbs, the increasing height differential between these merging bodies of water causes sudden and massive seaward discharges called “storm surge ebb-residual flow,” which create new channels and widen old breaches (Carter 1988:235; Wells 2001:152-153, 161). Such discharges are critical in fashioning the barrier coast and determining patterns of sediment dispersal across the adjacent shelf (Hayes 1975, 1979; Swift 1975, 1976; Swift *et al.* 1985). Archaeological and stratigraphic evidence at La Emerenciana indicates final site abandonment at around 1450 B.C. was related to such dynamic climatic conditions during a major El Niño event (Staller 1994).

Areas along the eastern equatorial Pacific between 1° to 3° South Latitude usually experience an incremental increase in precipitation during El Niño and decreases in duration and intensity north and south of 5° South Latitude (McPhaden 1999:950-951; Rodbell *et al.* 1999:519; Vuille *et al.* 2000:2520; Moy *et al.* 2002:163). The intensity and duration

of El Niño cycles are irregular, with mega events occurring roughly two to four times in a century (Moseley 1987:7). The 1983/84 Mega-El Niño was followed by another Mega event during 1997/98, but these were distinct in their effects and duration. El Niños vary in intensity and duration and occur in two to seven year cycles (Kerr 1998). The physics of the phenomenon are complex and represent a worldwide weather perturbation (Sandweiss 2003; Sandweiss *et al.* 1996, 2009).

The climate in coastal El Oro is classified as “semi-arid,” distinguished by a marked annual variation of wet and dry seasons (Ferdon 1950:52, Figure 17; Momsen 1968). The nine-month dry season lasts from May to January (Tables 1-2). However, orographic rainfall occurs in higher elevations in the sierra and highland streams drain into the Jambelí Estuary, maintaining a constant supply of fresh water to the lowlands (Ferdon 1950:43; Murray *et al.* 1975:346). Average annual precipitation is insufficient to sustain a year-round agricultural economy (Staller 1994; Tykot and Staller 2002). Contemporary agriculture is generally by floodwater farming and small-scale pot irrigation. The coastal savanna between the Arenillas and Buenavista Rivers is low relief topographically and does not lend itself to irrigation (Staller 1994). Regional survey indicates that pre-Hispanic irrigation canals or raised fields are completely absent. The driest subregion is directly adjacent to the Peruvian frontier, the coastal savanna or *Pampas de Cayanca* (Figure 2).

Average annual precipitation is highly variable, more than 90% occurring between May and July, the rainy season (Ferdon 1950:figure 17; Momsen 1968:92, 96; Wernstedt 1972:121; Cañadas-Cruz 1983:25; Staller 2001a:table III). Regional variability in rainfall between the seashore and further inland is evident by a 30% reduction in average annual rates between Puerto Bolivar and Machala, an overland distance of less than 10 km, and a 75% reduction between Puerto Bolivar and Zorritos at the Peruvian frontier (Table 1). Climatic conditions closely approximate other regions of the Gulf of Guayaquil such as the southeastern portion of Puná Island and inland regions of the Santa Elena Peninsula (Momsen 1968:96, table 1; Troll 1968). Annual temperatures average 24.7°C and range from 27° to 35°C in the northern part of the Gulf near Guayaquil and 22.7° to 27.1°C in southern coastal El Oro (Wolf 1892:389; Svenson 1946:405; Wernstedt 1972:121; Staller 2001a:table IV). A 4°C

Table 1. Average annual rainfall (mm), coastal El Oro Province, Ecuador.

Weather Station	Months											
	J.	F.	M.	A.	M.	J.	J.	A.	S.	O.	N.	D.
Machala	109.9	156.9	177	105.9	46.9	19.0	17.0	20.0	16.0	18.0	9.9	11.9
P. Bolivar	66.0	134.8	160	87.9	11.9	13.9	9.9	8.8	9.9	13.9	7.1	10.9
Zorritos	18.0	56.1	33	17.0	0.0	1.0	1.0	1.0	0.0	0.0	1.0	1.0

Average Annual Precipitation: Machala (708.9) P. Bolivar (505.9) Zorritos (129.0) (after Wernstedt 1972).

Table 2. Average annual temperatures (°C) coastal El Oro Province, Ecuador

Weather Station	Months											
	J.	F.	M.	A.	M.	J.	J.	A.	S.	O.	N.	D.
Machala	26.0	26.2	26.5	26.6	26.0	24.2	23.3	22.7	23.0	23.2	23.7	25.2
P. Bolivar	25.8	26.5	26.6	26.5	25.5	23.8	23.2	22.7	23.0	23.2	23.7	25.0
Zorritos	26.6	26.8	27.1	26.1	25.6	24.2	23.0	22.7	22.7	23.2	23.3	25.1

Annual Average Temperatures: Machala (24.7°) P. Bolivar (24.7°) Zorritos (24.7°) (after Wernstedt 1972).

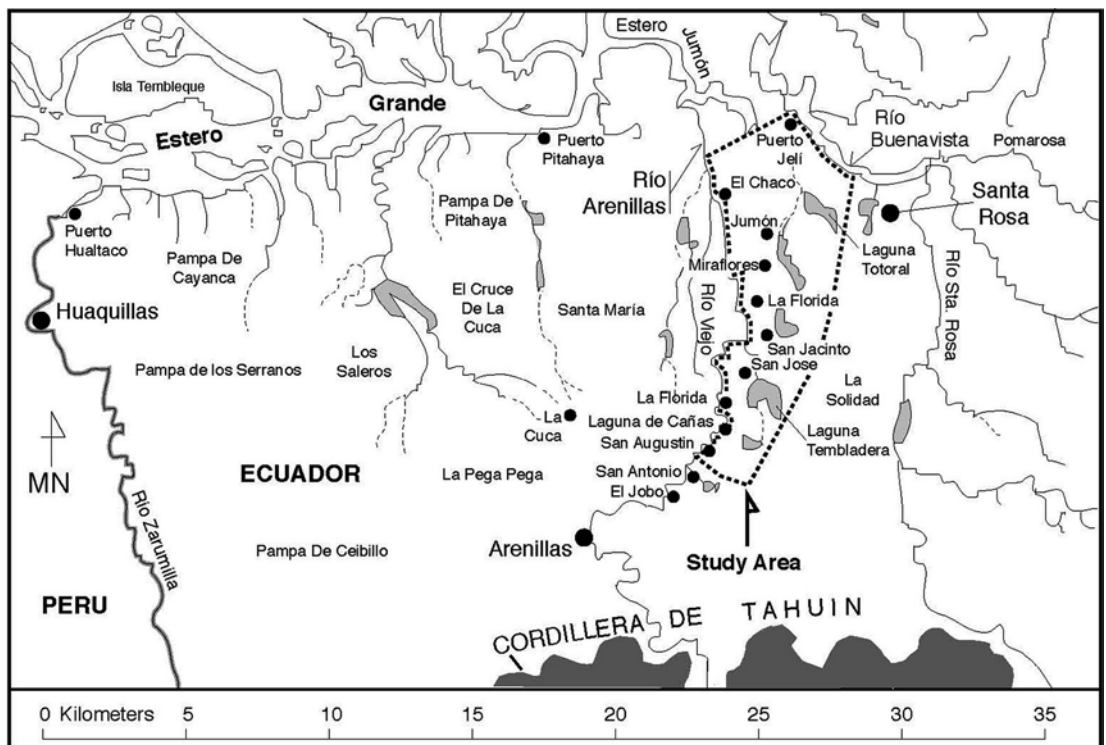


Figure 2. Southern coastal El Oro Province. Landscape features are indicated and regional names and study area are highlighted. Areas over 100 masl along the Cordillera de Tahuín are color indicated.

annual variation is a function of diurnal changes in cloud cover and the retreat of the Peru Current, while the 8°C annual variation to the north reflect higher levels of precipitation, greater humidity, and solar radiation (Ferdon 1950:51-52, 55, 56, figure 24; Momsen 1968:96-98; Murray *et al.* 1975:345; Cane 1983) (Table 2).

Environmental Setting: Coastal El Oro Province, Ecuador

Southern coastal Ecuador is situated in the Straits of Jambelí along the southern portion of the Gulf of Guayaquil and represents a barrier island estuary. Meandering streams interlaced with fresh and salt-water lagoons and oxbow ponds, are dispersed throughout the landscape – habitats particularly sensitive to climatic perturbations. Southern coastal El Oro Province encompasses the area from the Jubones River, north of Machala, to the Peruvian border, an east to west distance of about 45 km. The Cordillera de Tahuín is to the southwest and these foothills range between 100 to 200 masl and are distinct geologically from the Andes (Feininger 1980). The north/south transitional gradient represents a distance of about 65 km and ranges from wet subtropical forests in the Guayas Basin, Ecuador, to the extremely dry desert coast of northernmost Peru (Staller 1994, 2000). The estuary represents an ecotone, or transitional environmental zone and the local flora and fauna are characterized by a high incidence of endemism (Gentry 1977, 1986, 1988; Dobsen and Gentry 1991; Parker and Carr 1992). The region is renowned in the biological sciences for its endemic orchid species (*Orchidaceae* spp.) as well as hummingbirds e.g., (*Amazilia* spp.) (Ridgely and Greenfield 2001:356-359). Unlike other regions of the Ecuadorian coast, coastal El Oro is ecologically distinct due to the close proximity of the Andes, which begin their ascent only 15 km from the active shoreline² (see Figure 1).

The dry tropical forests consist of xerophytic thorn brush, dense clusters of mesquite or *trupillo* (*Prosopis pallida*) trees and *algarrobo* (*Hymenaea* spp. L.), as well as various species of columnar cactus (*Cereus* spp.) (Acosta-Solis 1959, 1970). A tall ceibo (*Ceiba trichistandra* Bakh) tree at the north end of the site is the most prominent natural feature

(Staller 1994:plate 7). Local villagers related that ceibos represent natural landmarks because they can be seen from great distances – the last remnants of what was, up to about eighty years ago, a biologically diverse old-growth forest (Staller 1994:189, 199-201; Staller 2001b).

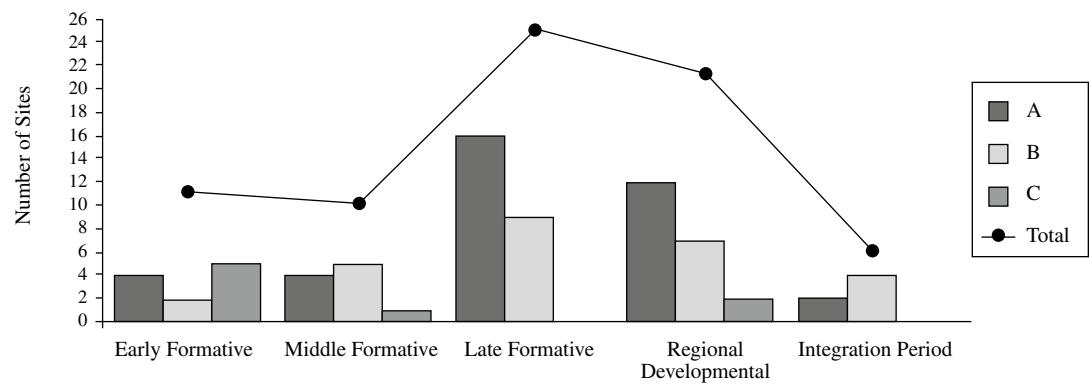
The area around La Emerenciana is referred to as “*los algarrobos*”, after the dense stands of *algarrobo* (*Prosopis* spp. L.) trees along this portion of the Buenavista River (Staller 1994, 2010; Tykot and Staller 2002). La Emerenciana is named after a small port once located 200 meters to the northeast along the Buenavista River (Staller 1994). The port was abandoned in 1964-65 after tectonic uplift made the river too shallow, indicating that, like northernmost Peru, this area is affected by tectonic uplift and/or subsidence (Staller 1994).

The dry tropical forests consist of five primary environmental zones: 1) mangrove forests situated on the offshore islands and foreshore lagoons and beaches; 2) Fresh and saltwater marshes and swamps beside the primary river channels, ox-bow lagoons and seasonal ponds; 3) dry tropical forests on the coastal plain or savanna; 4) a banded salt flat, or *saltiral*, distinguished by a dramatic reduction or a complete absence of surface vegetation, forming a natural barrier between mangrove (*mangal*) plant communities and the dry tropical vegetation of the lowland savanna (*pampas*); and 5) piedmont forests concentrated in the higher elevations of the Tahuín Cordillera and foothills of the Andes (see Figure 2; Table 3). The floral and faunal ecology are seasonally varied, highly diverse and ecologically complex, thus extremely fragile and susceptible to the effects of El Niño induced climatic variation (Wolf 1892:Map II; Cañadas-Cruz 1983; Caviedes 1975, 1984; Chapman 1976, 1977; Carter 1988).

Environmental and climatic research on a series of glacial lake deposits from the nearby highland Cuenca Valley, Ecuador produced extraordinarily detailed data on long-term El Niño events, chronology, tectonic activity, environmental change and the Holocene climate (Rodbell *et al.* 1999, 2002; Vuille *et al.* 2000; Moy *et al.* 2002; Hansen *et al.* 2003; Andrus *et al.* 2008). These lake core and archaeological data provide independent lines of evidence from the surrounding highlands and regions to the south that document possible effects of ancient El Niño related phenomenon. They are also evidence of El Niño chronology that precisely correlates with the multiple lines of data presented here.

² These topographic conditions more closely approximate regions to the south, the Peruvian and Chilean coast.

Table 3. Distribution of Prehispanic sites by association to environmental zones in the regional settlement survey.



Key:

- A = Sites located within the floodplain adjacent to wetland ponds and lagoons.
- B = Sites located on the lowland savanna on knoll tops adjacent to wetland ponds and lagoons.
- C = Sites located on the salt flats (saltiral) and fossil beach ridges adjacent to brackish ponds and the mangrove forest.

*Values are based upon cultural components from both single and multicomponent sites identified in regional settlement survey.



Figure 3. A vertical section of the fossil beach ridge associated with La Emerenciana as it appears along the recently modified Arenillas River channel. This beach ridge was clearly visible in regional survey and its cross section exposed after modification of the river channel by the Ecuadorian military as part of the Tahuín Dam Project (Facing SW).

Three fossil beach ridges were identified in systematic regional survey (Staller 1994, 2000). These geological features record accretional history that over the millennia spread fine white sand throughout the lowland savanna (Figure 3). Beach ridges in some cases represent proxy records of prehistoric El Niños³ (Richardson 1983). Other lines of evidence include flood deposits, and archaeological middens and soils.

Large-scale excavations at the Valdivia ceremonial center of La Emerenciana indicate repeated site abandonment as do sites identified in regional settlement patterns (Staller 1994). Geological and historical data further indicate this is a tectonically active region (Gleaser 1978; Barazangi and Isacks 1979; Feininger 1980, 1982). Although seismic activity cannot be completely ruled out as a factor to site abandonment, more recent climatic data indicate El Niño must also be taken into account. Climatic evidence suggests an increase in the frequency and duration of El Niño related phenomena along the eastern Pacific after 5800 cal B.P. with the modern and historical range established ~ 3000 cal B.P. (Sandweiss 2003; Sandweiss *et al.* 1996, 2001, 2007, 2009; Andrus *et al.* 2008). Such data appear to be important for our understanding of settlement patterning because during this time the region experienced an initial dramatic decline and subsequent increase in population density and site size during the Late Formative i.e., after c. 1000-800 B.C., involving a preference for alluvial soils and direct access to lowland lagoons and ponds (Staller 1994, 2000, 2010).

Pre-Hispanic Settlement Patterns along the Lower Arenillas River

Systematic settlement survey between the lower Arenillas and Buenavista River valleys recorded a total of fifty-two sites corresponding to the entire pre-Hispanic sequence and document the first evidence of Valdivia culture this far south along the coast (see Figure 4, Table 3). Survey methodology involved systematic 100% coverage of the terrain using both topographic maps and aerial photos obtained from the Ecuadorian military (see e.g., Willey 1953, 1962; Parsons 1974). Differences in

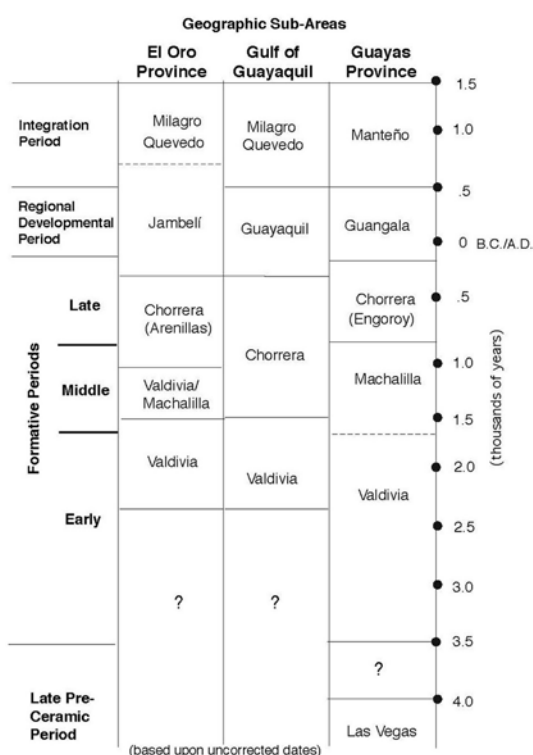


Figure 4. Cultural chronology of southern and southwestern coastal Ecuador.

settlement patterning have also been reported as indicators of climate and changes in human adaptation. Eleven late Valdivia sites were identified in two distinct patterns, along the most inland fossil beach ridge, and on knoll tops beside stream channels. Valdivia ceremonial centers apparent by the presence of two artificial earthen mounds were La Emerenciana and Jumón (Staller 1994, 2000). An earthen mound at La Emerenciana revealed four upright fully articulated bundle burials and one upright bundle burial was uncovered from one of the two earthen mounds at Jumón⁴ (Staller 1994,

³ The beach ridge under La Emerenciana is the furthest inland, thus the earliest of such geomorphological features (Staller 1994).

⁴ Burials at La Emerenciana included four adult females fully articulated bundle burials all associated with the stratum 5 floor 2 occupation and a sub-adult fully articulated associated with the final occupation of the site pertaining to stratum 6 (Staller 1994; Tykot and Staller 2002; Ubelaker and Jones 2002). The burial at the earthen mound in Jumón was also tightly flexed adult female (Staller 1994). These data are consistent with research at earthen mound at other Valdivia ceremonial centers such as Real Alto, Loma Alta, and San Isidro, supporting the prevailing assertion that they were first and foremost burial mounds (Lathrap *et al.* 1977; Norton 1982; Marcos 1988; Ziedler 1988; Staller 2001a).

2000, 2001a, 2001b). Middle Formative settlements had significant reductions in occupation density and site size, and clustered on knoll top locations beside stream channels with ready access to coastal lagoons, seasonal ponds, and alluvial soils. Late Formative sites were absent along the innermost beach ridge. It is precisely at the beginning of the Late Formative (3000 cal B.P.) that modern and historic El Niño frequency patterns are established.

Valdivia localities are situated along the foreshore and near-shore estuary, as well as riverine localities (Staller 2001a). Coastal sites listed under category “C”, are absent after the Early Formative and reappear during post-formative times – a period of over a millennium (Table 3). Increases in Late Formative settlements with high densities of occupation debris reflect a clear preference for knoll top localities beside stream channels with direct access coastal ponds and lagoons as well as terrestrial habitats and alluvial soils (Figure 5). A preference for riverine settings reflects a greater dependence upon domesticated plants and terrestrial animals and birds. The significant increase in site number and size during the Late Formative continues into the subsequent periods (Staller 2010:figures 4.22a-b). Such patterning suggests a corresponding shift in adaptation, possibly initiated by a greater overall frequency of El Niño.

Extensive lamination in Laguna Chorrera lake core deposits dated to between 3300 to 2600 cal. B.P. provide corroborating evidence of increased El Niño activity (Rodbell *et al.* 1999:figure 3; Moy *et al.* 2002:figure 1a-c). Increased El Niño frequency may explain the dramatic reduction in pre-Hispanic occupation during the Middle Formative, and why solely early diagnostics were identified, as well as the total abandonment of the foreshore between c. 1450-500 B.C. The destruction and/or burial of the barrier reef may have been related to such reoccurring El Niño related activity or to the Mega event associated with site abandonment in stratum 5 and provide a basis for understanding the extinction of oyster for over millennia. Oysters reappear in abundance during the Regional Developmental and Integration Period (see Figure 4, Table 3). Jambelí Phase occupations with deep occupation horizons and middens reappear in the foreshore. Shell middens dispersed throughout the *Pampas de Cayanca* near Huaquillas stand over 13 to 16 meters high and 150 to 200 meters at the base (Staller 1994:plate 1).

Aridity and climatic instability provided the greatest challenges to adaptation in southern coastal El Oro. Archaeological and stratigraphic evidence from settlement survey and excavation indicates the ancient economy was mixed and diversified, and included hunting, plant gathering,



Figure 5. Fisherman casting a net on the Laguna de Tembladera during the dry season. The heavy fog reflects the *garua* that normally permeates this region of the coast from early April until late December (Facing North).

agriculture, and aquatic resource exploitation as primary components (Staller 1994, 2000, 2001a, 2001b, 2010). Stable isotope signatures from the La Emerenciana skeletons indicate aquatic resources formed a major portion of the diet and, although maize was consumed, it played a minor role in the subsistence diet and primarily consumed as beer or *chicha* (Tykot and Staller 2002:tables 5-6; see also Lathrap *et al.* 1977). There is some reason to suspect that the climatic and environmental changes induced by El Niño, may have played a central role in fostering a greater dependence upon agriculture, although the overall response to variations in resource availability generally favored flexibility or increased diet breadth over specialization.

Pre-Hispanic Occupations at La Emerenciana

The Valdivia ceremonial center La Emerenciana covers some 12.72-hectares making it one of the largest Early Formative in coastal Ecuador (Lathrap *et al.* 1977; Norton 1982; Marcos 1988; Staller 2001a). It is situated on a fossil beach ridge at 2.5 masl about 3 km from the present shoreline (Staller 2001b:figure 7, 2003). Primary occupation spanned 700 to 650 years and corresponds to the final portion of the Early Formative Period c. 2200-1450 B.C. The Jelí Phase ceramic complex associated with Valdivia sites in this region is distinguished by the earliest directly and indirectly dated pedestal bowls, stirrup-spouts and single spout bottles, as well as, the earliest red on white banded pottery in

the Andes (Staller 1994:356-400, 2001b:figures 22-29). Various archaeologists have suggested the red on white-banded tradition in southern highland Ecuador and the northern highlands of Peru is representative of the earliest pottery cultures in this part of the Andean highlands and representative of a cultural horizon called Chaullabamba (Uhle 1920a-b, 1922a, 1928, 1929, 1938; Staller 2007; Collier and Murra 1943; Hocquenghem 1991; Hocquenghem *et al.* 1993; Moore *et al.* 2008). These regions of the Andean sierra relate to non-Quechua and Aymara speaking cultures involved in the early spread of ceramic technology, maize (*Zea mays* L.), *Strombus galeatus* conch and *Spondylus* spp. oyster shell to the adjacent highlands and south along the coast (Collier and Murra 1943; Paulsen 1974; Hocquenghem *et al.* 1993; Staller 2007). Initial appearance of red on white-banded diagnostics in coastal El Oro Province dates to as early as 2200 B.C., this based upon both AMS and conventional dates (Staller 2001b, 2007; Staller and Thompson 2002; Tykot and Staller 2002). Initial site abandonment and burial by a fossil beach ridge may relate to a Mega-El Niño dated to 2150 B.C. consistent with similar dates in regions of the Andes (see Rodbell *et al.* 1999:figure 3; Dillehay and Kolata 2004:table 1). Reoccupation and fossil beach ridge formation is ^{14}C dated to between c. 2200-1450 cal B.C. Another abandonment is followed by a brief final reoccupation dated to c. 1450 B.C. (Table 4). Geographic and cultural definitions of the various archaeological subregions have been previously

Table 4. La emerenciana radiocarbon chronology

^{14}C Laboratory No.	^{14}C B.P./	Calib.4.1.2 1- δ age range B.C.
a) SMU-2241	3361 B.P. \pm 246 years	1935-1323 cal. B.C.
b) SMU-2226	3400 B.P. \pm 220 years	1941-1428 cal. B.C.
c) Beta-125106	3720 B.P. \pm 40 years/ 3700 B.P. \pm 40 years	2137-1979 cal. B.C.
d) SMU-2225	3707 B.P. \pm 148 years	2288-2245 cal. B.C.
e) Beta-125107	3810 B.P. \pm 50 years/ 3860 B.P. \pm 50 B.P.	2240-2201 cal. B.C.
f) SMU-2563	3775 B.P. \pm 165 years	2459-1922 cal. B.C.

Note: All material dated is charcoal. BETA dates are AMS corrected for $^{13}\text{C}/^{14}\text{C}$ fractionation. SMU dates are standard assays. All dates are calibrated using Calib 4.1.2 (Stuiver *et al.*, 1998), with a minus 24-year Southern Hemisphere atmospheric sample adjustment and are reported here as a one-sigma range. Staller (1994:393-396, Figure 5) provides the contextual information on the SMU dates. Staller and Thompson (2002:45, Table 10) provide contextual information on the Beta dates. All are from Stratum 5 except SMU-2241, which is from Stratum 6.

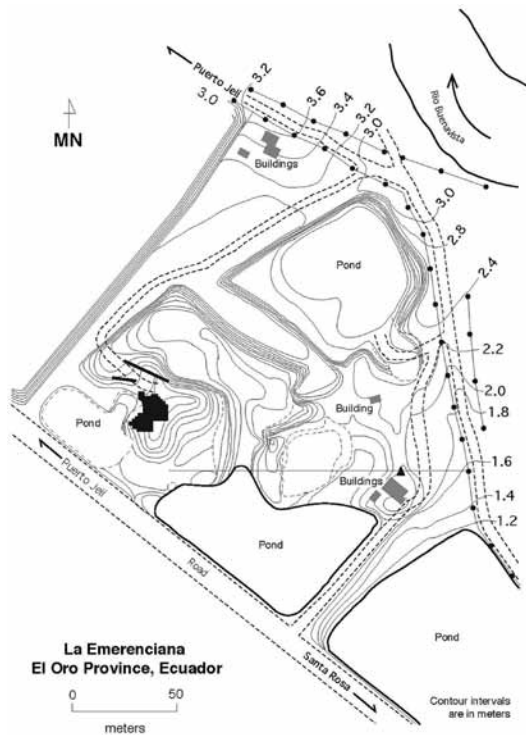


Figure 6a. The ceremonial center of La Emergenciana is situated on a fossil beach ridge, one of three such topographic features identified in regional survey in the *salitral* or salt flats. Two earthen mounds were identified and the larger NW mound was the focus of the excavation. Another earthen mound was identified under a modern habitation just south of the site datum. Parts of the surrounding coast were modified by shrimp pond construction.

discussed in detail (see Staller 2001a:197-201, figure 2).

Excavations at the ceremonial center of La Emergenciana

La Emergenciana is situated with direct access to maritime and estuarine resources, as well as riverine soils, and is the largest site with earthen mounds identified in regional survey for any time period. Excavations were restricted to the summit of the northwest platform, one of two earthen mounds at the site (Figure 6a). The earthen mound is oval in shape 2.5 meters high, and measures 200 meters (N-S) by 150 meters (E-W) with two oval clay platforms on the summit⁵ (Staller 1994:209). Four

⁵ The north platform measures 5 (N-S) by 4 (E-W) meters, while the southeast platform measures 3.5 (N-S) by 2 (E-W)

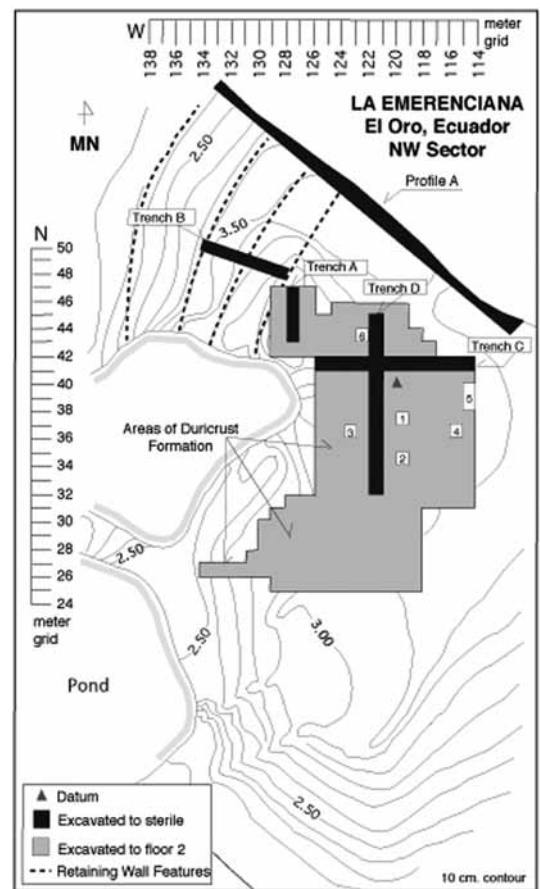


Figure 6b. Excavations on the NW platform mound showing designated Trenches A-D and Cuts 1-6 dug to sterile levels, and units dug to the surface of floor 2.

trenches (A-D) were dug to sterile levels, and 332m² of a buried prehistoric occupation surface (floor 2) were exposed. Five-meter square units (Cuts 1 to 4, 6), and a 1 by 2 meter pit (Cut 5) were excavated to sterile to record more specific data on stratigraphic variation (Figure 6b). A twenty-nine meter long vertical section (Profile A) was cleared in order to provide continuous stratigraphic information on this portion of the mound⁶ (Staller 2001b:Figure 13).

meters, and is 15 cm thick (Staller 1994; Tykot and Staller 2002). These clay platforms (*baharrique*) are held together with cane branches that were tied together with cord. Their impressions were still visible when they were exposed from the overlying sediments (Staller 1994).

⁶ The excavations at La Emergenciana were carried out in part because the site was partially destroyed by artificial ponds constructed on the western parts of the site by local shrimp farmers. The primary reason that the site was still partially intact was because the ponds would not hold water because



Figure 7. Profile A in cross section immediately after exposing the sediments. Trench B is in the forefront. Artificial step or terrace features were initially identified in the Trench B excavations and verified in Profile A (Facing SW).

Profile A and the Trench D excavations uncovered three descending retaining walls or stepped terraces in the west and northern parts of the earthen mound, and these modifications were verified in the Trench B excavation (Figures 6a, 7).

Excavations were primarily by natural stratigraphic layers and uncovered one hundred and thirty nine features, primarily consisting of architectural modifications associated with artificial mound construction, burials, four fully articulated upright tightly flexed burials, and various ritual offerings (Staller 1994, 2010; Staller and Thompson 2002:figure 7; Ubelaker and Jones 2002). Initial occupation is associated with stratum 3 and dated to ca. 2400 B.C., followed by beach ridge formation (stratum 4). Later occupations date to between 2000-1450 B.C. and are associated with stratum 5 and a brief reoccupation associated with stratum 6 (Figure 8, Table 4). Lithic debris was limited to only 20 artifacts, but included exotics; two obsidian flakes sourced to different outcrops in the Valley of Quito that represent the earliest dated obsidian in coastal Ecuador (Staller

1994; Asaro *et al.* 1994). Most artifacts consisted of sherds, and clear evidence of ceramic production⁷ (Staller 1994). Offerings consisting of ocher covered pebbles and a few chipped quartz flakes and some polishing stones⁸ (Staller 1994).

Differences in the color, texture and composition of the various strata allowed identification of the stratigraphic units and they were classified on the basis of grain size analysis (Tables 5a-b). Layered sequences of stratigraphic units were divided according to their artifact content. All trenches and units were excavated following the conformities and contours of the natural stratigraphic layers, and differentiated by the physical properties of the strata.

⁷ Ceramic manufacturing at the site was documented by a semi-subterranean kiln with waster sherds at its base identified in cross section in the southern portion of the site (Staller 1994).

⁸ Ocher covered pebbles and stones are commonly left as offerings at later Andean sites in both the highlands and coast, but have their earliest appearance at this and other Valdivia ceremonial centers. There was no evidence of domestic habitations in contrast to other Valdivia ceremonial centers such as Real Alta and Loma Alta (Marcos *et al.* 1976; Lathrap *et al.* 1977).

of the fine white sand on which the site stands (Staller 1994, 2001).

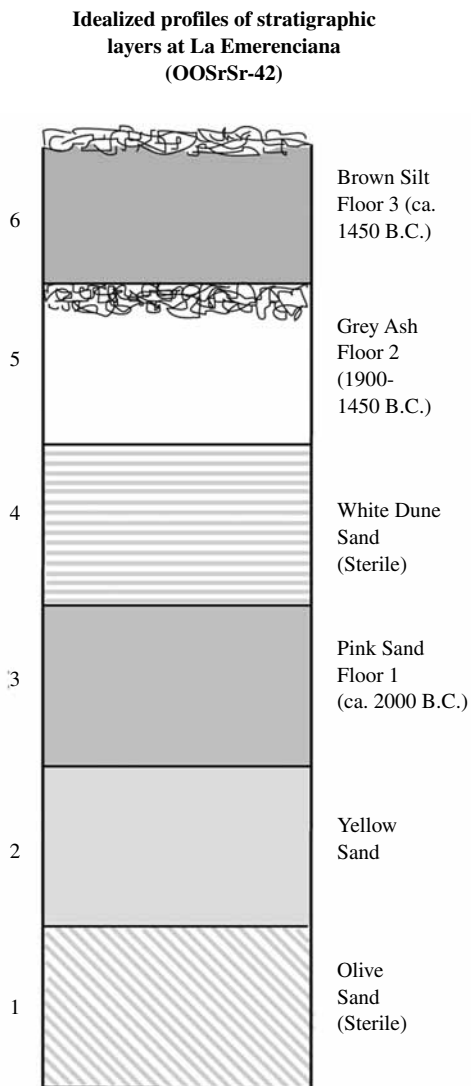


Figure 8. Idealized profile of the various stratigraphic layers 6-1 identified in excavations.

Artifacts established the stratigraphic sequence and permitted the recognition of reversed stratigraphy, as well as primary and secondary deposits (Joukowsky 1980:152; Stein 1990:516).

A concentration of marine shell extending over 40 cm in cross section was identified on the northern portion or seaward portion of the site in Trenches A, B, and northern portion of Profile A. Arbitrary increments of 20 cm were used, since the smallest natural unit of analysis (i.e., shell layer) was too large to detect subtle changes in the vertical distribution. On the summit of the earthen mound the shell layer has a maximum depth of only 5 to

10 cm over the surface to floor 2, and was absent on the southeastern portion of the mound in the Cut 4 and 5 (Figure 9). The homogeneous grey ashy loam (stratum 5, living floor 2) has been identified at Valdivia sites throughout coastal Ecuador (e.g. Meggers *et al.* 1965; Estrada *et al.* 1964; Lathrap *et al.* 1977; Marcos 1988) and also at later Jambelí Phase sites (Staller 1994; see also Currie 1985). Archaeologists proposed explanations for such soils, although geological or geomorphological evidence has been lacking. The surface of stratum 5 designated as living floor 2, represents a buried occupation surface (see Figure 9). Highest concentrations of cultural remains, primarily sherds and ancient shells, were in the uppermost 10 cm – apparently because the interface formed on a stable soil surface (Staller 1994).

Natural Stratigraphy at La Emerenciana

The A-Bw/Btn-Bk horizon sequence of soils is characteristic of well-drained semi-arid conditions (Table 6). The five lower strata were continuous across the site, and two stratigraphic layers (3 and 5) contained prehistoric artifacts pertaining to Valdivia culture. Stratum 6 was restricted to the northwest sector and associated with the final abandonment (Staller 2001a:table II). The following geomorphological, archaeological and shellfish counts are presented to test the hypothesis that repeated site abandonment and widespread geomorphological changes were induced by ancient El Niño events (Staller 1994:144-153).

The lowest layer, stratum 1 is olive beach sand with nodules of decomposing organic and gravel inclusions identified at 97 cm below datum (Figures 10-12). Inspection of a vertical section along the stream channel measuring 20m long and 5m deep suggests this layer reaches to at least three meters, with gravel lenses and shell lag deposits, remnants of foreshore deposits of fluvial origin throughout. Cross-bedding was noted below 2.5m, the result of fluvial processes – post-depositional weathering through the movement of groundwater, leaching and dissolution of carbonates altered the chemistry of the sediment (Woods 1977:249). Various shell lag deposits and gravel lenses indicate groundwater movement, evapotranspiration or hydrological and sea level changes over geological time (see Figure 12). The gravel lenses and clay shell lag deposits effect retention of organic matter,

Table 5a. Grain size analysis (Hydrometer).

STRATUM	WEIGHT (g)	GRAVEL (%)	SAND (%)	SILT (%)	CLAY (%)
6	50.00	0	88	8	4
6-5 interface	40.84	<0.5	90	7	2
5-4 interface	36.08	<0.5	83	11	6
4	50.00	0	90	4	6
3a (calcrete)	56.13	0	96	2	2
2	50.00	0	92	6	2
1	50.00	0	92	8	0
1d (shell lag)	50.00	0	94	2	4

Table 5b. Grain size analysis (Pipette)

STRATUM	WEIGHT (g)	SAND WEIGHT (g)	SAND (%)	SILT (%)	CLAY (%)
5b	27.14	23.73 (3.34)	87	6	6
5	22.92	19.15 (3.67)	84	7	9
1a shell lag	36.32	30.45 (5.73)	84	7	10

Note: Weight (g) refers to corrected total weight. Residuals or dry weight is given in parentheses beside the sand weight and is also in grams. Analysis and interpretation was carried out by the department of geography at the University of Wisconsin-Milwaukee.



Figure 9. Floor 2 excavations showing units, trenches and exposure of the paleosol. Note the depth of Stratum 6 in the western part of the excavated area (Facing North).

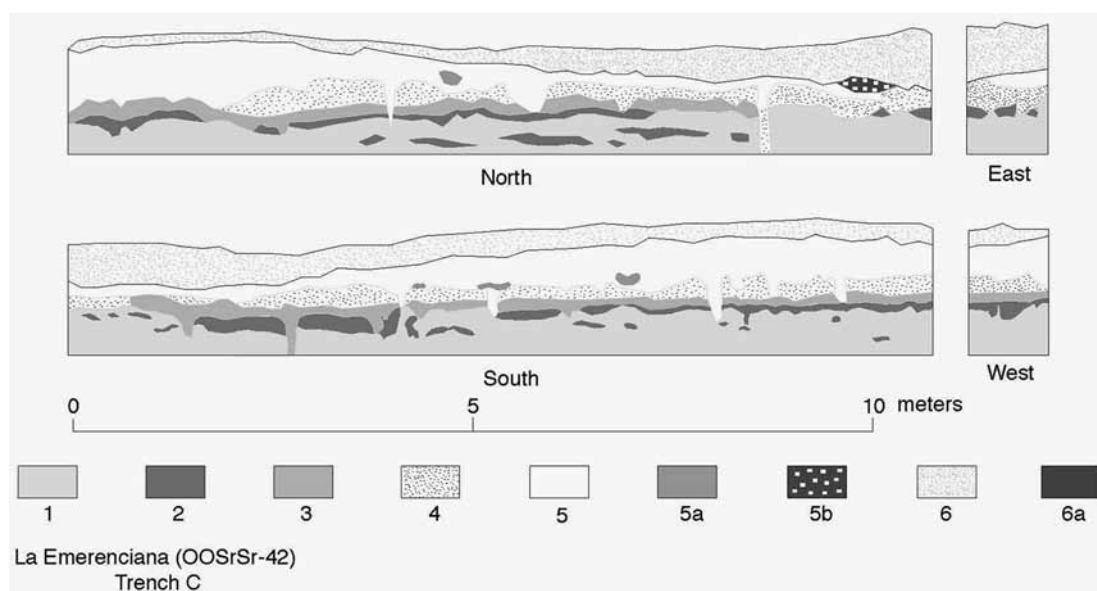


Figure 10. Vertical section of Trench C showing the stratigraphic layers. See Table 3 for key to stratigraphic layers and sublayers.

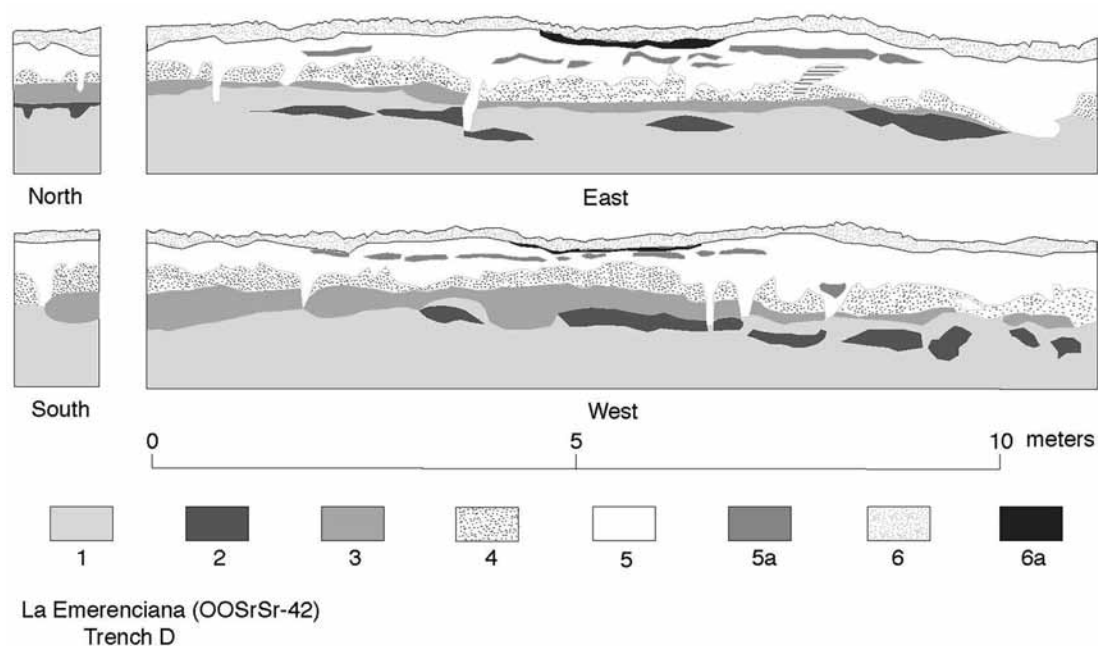


Figure 11. Vertical section of Trench D showing the various stratigraphic layers. See Table 3 for key to stratigraphic layers and sublayers.

explaining in part the preservation of faunal remains (Staller 1994). Macrobotanical remains were not preserved in any of the sediments, however plant microfossils were recovered from carbon residues in pottery sherds (Staller and Thompson 2000, 2002; Tykot and Staller 2002).

Stratum 2 is yellow sand the result of post-depositional weathering associated with groundwater movement (see Table 6). The deposition and decomposition of organic and inorganic compounds left concentrations of calcium compounds, carbon, phosphorous, and trace metals that apparently created

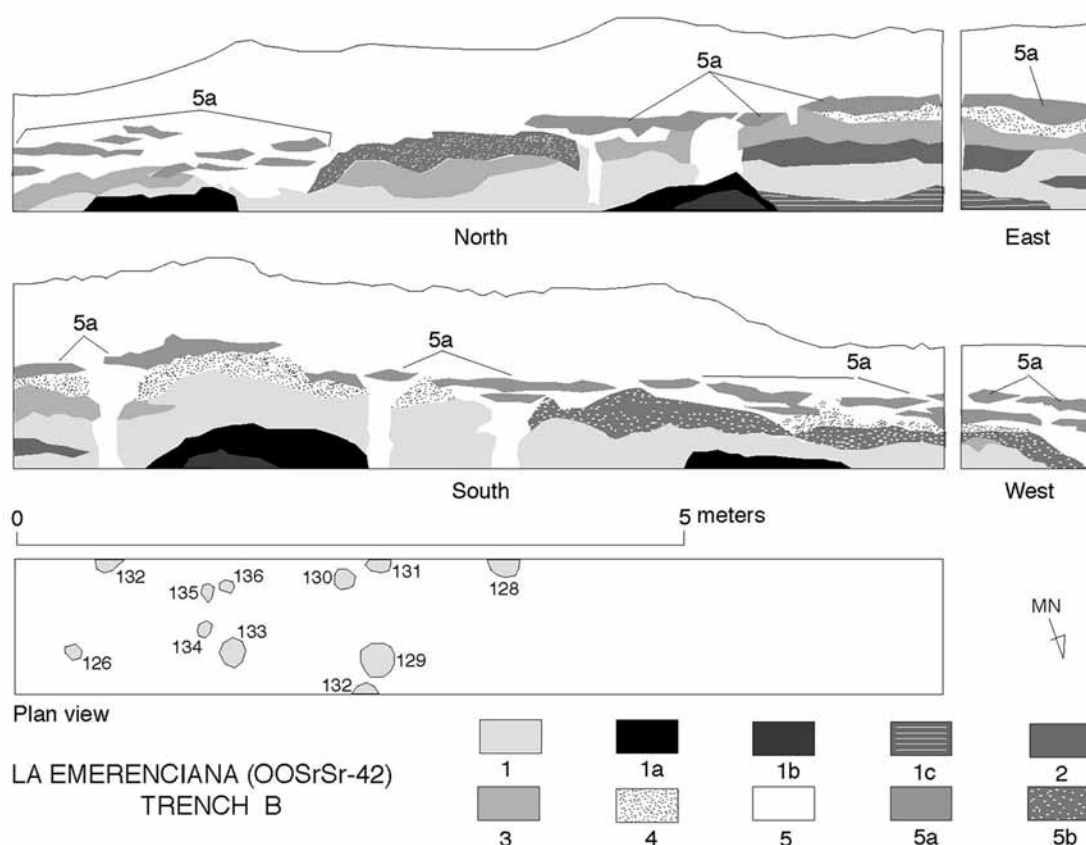


Figure 12. Vertical section and plan view of Trench B showing the various stratigraphic layers. Features are architectural, posts molds associated with the construction of retaining walls. See Table 3 for key to description of layers and sublayers.

chemical changes in the soil. Geological studies indicate this coloration may also be produced by the movement of iron oxides via water percolating through shell lag deposits (Woods 1977:298).

Stratum 3 is pink sand extending between approximately 94 and 124 cm below datum, containing archaeological features including pits, numerous post molds, and portions of foundations of habitation structures, indicating this was a prehistoric occupation surface designated living floor 1 (see Figures 10-11). Ceramic diagnostics correspond to the middle of the sequence and, despite the various features, there was a relative paucity of marine shells or artifacts, and a total absence of organic and faunal remains – probably related to leaching and post-depositional alterations in the lower layers. Leaching of phosphates and iron oxides from stratigraphic layers 5 and 3 was evidenced in the underlying layer and calcrete deposits by a yellow to light brown coloration (see Table 6). The

abundant faunal remains found with various features in stratum 5 were not preserved in this layer.

Stratum 4 is a sterile layer made up of fine white quartzitic dune sand with calcium carbonate inclusions. Some blending was noted with the overlying grey ashy loam and part of a fossil beach ridge extending to the Peruvian border and beyond (see Figure 3). Lake core deposits in the Andes Cordillera directly northeast provide indirect lines of evidence to strengthen the stratigraphic interpretation for repeated site abandonment. Beach ridge formation is often a result of El Niño related phenomena and may correspond to a Mega-El Niño recorded in the Laguna Pallcacocha lake cores ^{14}C dated to 4040 B.P. (Rodbell *et al.* 1999:81, figure 3; Moy *et al.* 2002:figure 1a-c). This chronology roughly coincides with an El Niño dated to ca. 2150 B.C. recorded at several sites in the Jequetepeque River valley to the south (Dillehay and Kolata 2004:4326, table 1). Beach ridge formation is

Table 6. Stratigraphic layers at La Emerenciana

Stratum	Depth	Horizon	Color	Description
6	0-55 cm	A	10YR 5/3 -10YR 5/4	Brown fine silty loam, loosely consolidated in the upper levels, denser in lower levels, with evidence of bioturbation (fluvial deposit).
5	15-93 cm	B	10YR 6/1 -10YR 5/1	Homogeneous grey ashy loam, loosely packed, very fine texture, fine quartz inclusions with the consistency of talc, and artifact and shell remains in the upper-most levels of the stratum (Living Floor 2) (ethnostratigraphic).
4	36-92 cm	C	10YR 8/3	White dune sand, finely textured very loosely consolidated, with calcium carbonate inclusions in the upper levels (eolian deposit).
3	78-145 cm	Bwn	7.5YR 6/4 -7.5YR 7/4	Pink quartz sand finely textured well consolidated, free of inclusions (Living Floor 1) (ethnostratigraphic).
2	64-134 cm	Bwk	2.5Y 8/6 -2.5Y 8/8	Yellow sand finely textured, loosely consolidated, with calcium carbonate small pebble inclusions (3 mm-1 cm) (eolian deposit).
1	97-cm	C	5Y 8/2 -5Y 8/4	Olive white sand, finely textured, moderately packed, with small (3 mm-2 cm) beach pebbles and calcium carbonate inclusions (fluvial deposit).
6a	5-28 cm	Ap	10YR 8/2 -10YR 8/3	Fine white ash with carbon inclusions. A substratum is a result of recent agricultural activity (ethnostratigraphic).
5a	6-72 cm	Bwt	10YR 8/1 -10YR 8/4,	Densely packed pale white clay fine textured, free of inclusions, hard and densely packed. Represents a prepared clay surface (ethnostratigraphic).
5b	65-80 cm	Bw	10YR 3/6 -10YR 4/2	Dark brown ashy loam finely textured, the result of post-depositional weathering and decomposition of Stratum 5 (ethnostratigraphic).
5c	57-74 cm	various	Animal burrow	
4a	37-45 cm	Bk1	10YR 7/4	Very pale brown, extremely hard calcrete conglomerate, calcrete sand with no sublayers identified in profile, some inclusions (post-depositional weathering).
3a	57-145 cm	Bk2	10YR 6/4	Light yellowish brown, extremely hard calcrete nodules high clay fraction (post-depositional weathering).
1a	57-68 cm	Bt	5Y 6/4	Shell lag deposit, hard yellow olive clay finely textured with shells inclusions throughout (eroded deposit).
1b	65-82 cm	Bg	2.5Y 6/8	Light olive yellow clay finely textured with organic nodules, high clay fraction, and shell inclusions on the bedding plane (eroded deposit).
1c	45-150 cm	Bt	10YR 8/3	Fine fraction white clay with extensive small to medium sized beach pebble inclusions (eroded deposit).
1d	57-145 cm	Bk3	10YR 6/4	Light yellowish brown, extremely hard calcrete conglomerate made up of calcified sand, no inclusions, but nodules and internodular fillings (post-depositional weathering).
1e	57-145 cm	Bk4	2.5YR 5/4	Reddish brown, extremely hard calcrete internodular, filling has a relatively high clay fraction (post-depositional weathering).

Note: All soil colors are classified using the Munsell Soil Color Chart 1975 Edition. Differences in color were sometimes noted within a particular stratum, and these designations were the most characteristic for the stratum as a whole. Depths are given as below datum, and indicated as minimum and maximum levels which of course varied in different areas of the excavations.

roughly contemporaneous with those recorded in the nearby lake cores and archaeological sites in northern coastal Peru supporting the hypothesis that it is a result of El Niño.

Stratum 5 extends between 10 cm to just under a meter in depth, with an average thickness of roughly 35 cm (see Table 6). Stratum 5 constitutes the existing land surface over 75% of the total area of the site and represents the primary occupation floor. Architectural features associated with the northwest earthen mound measure 75m north-south by 47m east-west and approximately 1.5m high (Staller 1994:319, 2001a). Few cultural remains were found in the lower and middle levels, indicating it was kept meticulously clean – a cultural pattern consistent with ceremonial centers in other regions of the Andes. Nevertheless, the mound was covered with oyster shell and sherds during the stratum 5 abandonment and later stratum 6 reoccupation, perhaps intentional to protect the mound from rising floodwater (Staller 1994).

Evidence of disconformity in the lower interface of stratum 5 may be a result of depositional events (see Figures 10-11). Most archaeological features were in the upper 10 cm and either ritual or architectural including two complete oval or elliptically shaped daub platforms, four burials, numerous clay-lined and sealed offering pits, retaining walls, and post impressions (Staller 1994, 2001b; Ubelaker and Bubniak Jones 2002). Ceramic diagnostics and ^{14}C dates suggest occupation corresponding to the final portion of the culture sequence, ca., 2200-1450 B.C.⁹ (see Table 4).

A very lightly to strongly carbonate massive duricrust formation was identified on living floor 2, stratum 5 and under stratum 6 on the SW side of the platform (see Figure 6b). The 5 to 10 cm thick calcified quartzitic sand crust is made up of soft and friable silicified sands with no apparent nodular development indicating this represents a silcrete the result of superficial diagenesis (Goudie 1973:8). Silcrete formations are related to seasonal fluctuations in climate resulting in the chemical alteration of sediments through hydration (*ibid.*:112). They form near stream channels in areas with a high water table – characteristics that describe the coastal setting and geomorphology (*ibid.*:17, 36). Silcrete

crusts may be eolian, as the constant abrasion of fine-grained quartz sands lead to concentrations of fine siliceous dust, which is susceptible to solution and alteration by fog dew (*garúa*) or changes in groundwater (*ibid.*:140-141).

This formation suggests the earthen mound and surrounding area underwent climatic fluctuations that included high wind velocities and a rise in ground water levels – factors consistent with the onset of an El Niño which first effected the adjacent highlands in the middle of what is normally the dry season i.e., July or August. Presumably, after stratum 5 abandonment increased precipitation may have raised the ground water immediately before inundation by flooding and/or associated seismic activity, and site burial through tidal asymmetry. Environmental reconstruction suggests the area to the north was the side facing the prevailing winds and the main lagoon. During periods of intense precipitation, elevated lunar tides may have left sea salts in solution on what was the foreshore and salt flats. Post-depositional weathering or evapotranspiration later crystallized these marine salts because the evaporation rate is greater and the ground water richer in dissolved sea salts resulting in duricrust formation, an interpretation consistent with such climatic and geomorphological processes.

The uppermost layer, stratum 6, is dark brown silt loam ranging between 10 cm to 58 cm thick, and represents a flood deposit, possibly combined with seismic activity (Keefer and Moseley 2004; Keefer *et al.* 2003). Recent Mega-El Niño events indicate this region is highly susceptible to denudation and inundation and that coastal flooding is related almost exclusively to such events. Moreover, research on El Niño related effects associated with seismic activity indicate that when such phenomena occur together, they can produce landslides, increased runoff and sediment transport during post earthquake rainstorms (Keeler and Moseley 2004:10878, 10881). Ceramic diagnostics and ^{14}C dates from stratum 5 suggest stratum 6 was deposited at ~ 1450 B.C. (see Table 4). Stratum 6 has the appearance of alluvium and is loosely consolidated and clearly distinct in color and texture from the underlying strata. Environmental reconstruction suggests the area of the site faced the main estuary lagoon and artifacts from stratum 6 were from deposits disturbed by bioturbation through plant root action or recent human activity. A horizontal feature interface representing a plow zone, noted as thin lenses of burned ash (Sublayer 6a), extend

⁹ Like other earthen mounds at Valdivia ceremonial centers, the NW platform was also a burial mound (see e.g., Lathrap *et al.* 1975; Marcos 1988; Marcos *et al.* 1976)

to 15 cm in depth in several areas (see Figure 11). Anecdotal evidence indicates this was the result of recent small-scale slash and burn maize agriculture. The only area where such cultivation was carried out was where stratum 6 represented the existing land surface.

Evidence from the nearby Laguna Pallcacocha lake cores record an El Niño event date to c.1500 B.C. which, given the apparent high levels of sediment load, may have been accompanied by or occurred subsequent to seismic activity (Jordan *et al.* 1983:figures 1 and 2; Rodbell *et al.* 1999:figure 3; Moy *et al.* 2002:164, figure 1; Hansen *et al.* 2003:102). Tectonic activity associated with the Atacazo Volcano in the Jama Valley, Manabí Province date to 1680 to 1550 B.C (Isaacson 1994:135). Subsequently, Hansen *et al.* (2003:102) report a 1 cm thick tephra layer from this lake core related to volcanic activity dated to 2500 cal. B.P and another just above it to 2200 cal. B.P. (Rodbell *et al.* 2002:figure 3). Such volcanic activity did not have a dramatic effect in this region as no visible evidence of tephra deposits were identified in regional survey, with pre-Hispanic deposits readily discernible on the existing land surface and in cross section on exposed profiles. The central and northern highlands of Ecuador ranging between 5° North to 3.5° South Latitude are the regions with the most volcanic activity. Ceramic diagnostics and ¹⁴C dates from stratum 5 suggest a possible correlation between the 2500 cal. BP eruption and other events of the same approximate age, and site abandonment.

These internally consistent lines of evidence infer stratum 6 was deposited as a result of a flood associated with tidal asymmetry, possibly combined with and/or closely following seismic activity. The stratum is anomalous in the sense that it is restricted to only a portion of the site and largely absent in the area to the west, where the existing surface consists of the fine white sand similar to stratum 4. Research on seismic activity in association with El Niño events in other Andean regions has shown that anomalous flood deposits may be a product of seismic landscape shattering and El Niño related floods and when these occur in combination, they result in particularly large runoff and sediment transport (Keefer and Moseley 2004:10882).

In addition, adaptive shifts in response to bio-turbation were also noted in settlement patterns. Perhaps the most dramatic evidence is the total disappearance of oyster (*Ostrea columbiensis* L.)

from all Middle and Late Formative sites identified in regional survey. Barrier reefs upon which such oyster beds existed were probably buried by excessive runoff of fluvial sediments and/or uplift and destroyed (Nummedal 1983). Seismic activity would have created extraordinarily high sediment loads in the coastal stream channels¹⁰. Uplift would have exposed at least some of the beds to the elements and to human and terrestrial predators, as this is a shallow estuarine environment. Moreover, stratum 6 represents the only transgressive unit in an otherwise progradational stratigraphic sequence (Staller 1994:table 13).

The permeability of eolian sediments affected the geomorphology of the various strata identified in the excavations. The coarser grain sediments that make up stratum 6 are more permeable than the finer grained sands and ashy loam below. Consequently, flow of rainwater percolating through the lower stratigraphic deposits is more restricted (Holliday 1990:536). Differences in the permeability of the uppermost layer and the grey ashy loam below it are reflected by gaps, or separations of as much as 5 cm, related to the buildup of rain water at the buried stratum 6-5 interface. The poorly drained sediments that make up the lower layers (4-1) is further evident by the presence of decomposed or hydrated organic matter throughout.

In summary, the stratigraphy is characterized by well-defined interfaces with clear differences in color, texture and composition among the various layers. The brown silt and ashy loam, strata 6 and 5 respectively, essentially preserving the overall integrity and stability of the underlying strata. The moist and fine grain sands quickly lost their overall structure when exposed to the natural elements, making the delineation archaeological features extending into the lower sand layers (strata 4 to 1). Stratigraphic evidence at La Emerenciana and dates from nearby lake cores support the interpretation of repeated site abandonment due to climatic upheaval and/or geomorphological changes consistent with El Niño. Layers 4 to 1 represent seaward components of a barrier coastline, shoreface, foreshore, backshore, and dune facies respectively. With

¹⁰ Oyster shells are totally absent from Late Formative sites. Detailed research of several partially exposed shell middens standing over twelve meters high near Huaquillas indicate a total absence of oyster shell over a thousand year time period (Staller 2000).

the exception of uppermost layer stratum 6, the stratigraphic sequence, is consistent with progradational coastal development. Uppermost stratum 6 represents a transgression consistent with a low energy estuary with deposition due to flooding, tidal wave asymmetry, and/or a dramatic influx of fluvial sediments, an indicator of El Niño and also an indication such climatic events may have been associated with seismic activity (Arntz 1986:7, 18-19; Keefer *et al.* 2003:figures 2, 3; Keefer and Moseley 2004).

Shell Frequencies: Analysis of Minimum Number of Individuals (%MNI)

Minimum number of individual (MNI) shellfish counts from the midden on the northern portion of the mound provides additional lines of internally consistent evidence with the stratigraphic and geomorphological data. A shell midden on the northern edge indicate oyster was the most intensively exploited species this on the basis of %MNI frequencies from Profile A and Trench A and B (Figures 13-14; Tables 7a-d, 8a-d). Multidisciplinary evidence indicates this is a single depositional event, perhaps to protect the retaining walls surrounding the north or seaward side of the earthen mound. In order to minimize sample bias shell species were collected from randomly selected excavation units. Oysters were most intensely exploited in stratum 5, followed by Pointed Venus Clam from coastal lagoons and mudflats, and *Protothaca* Clam buried in nearby mudflats or coastal bays (Olsson 1961:297-298; Keen 1971:190, 193). Such distinct species provide indirect indications of adaptive shifts, and geomorphological or environmental changes related to El Niño. Moreover, a 40% reduction in the size of oyster specimens in stratum 6 indicates the barrier reef was affected by geo-climatic changes¹¹.

Six hundred oysters were randomly selected from a predetermined number of arbitrary selected excavation units from stratum 5 and 6, 300 from stratum 6 and another 300 from stratum 5. Oysters from stratum 5 averaged between 90 and 125 mm long, while those in stratum 6, 45 to 70mm. The 40% size reduction was visibly evident in excavation and on floor 2 surface remains. Constricted growth rings were evident on the majority of

bivalves, particularly *Anadara*, or *pata de mula* (*A. grandis* Broderip & Sowerby) and simile ark or *concha prieta* (*A. tuberculosa* Sowerby) providing direct evidence of climatic perturbation supporting the hypothesis that strata 5 and 6 abandonment were a result of climatic change related to El Niño events. During such climatic perturbations alluvial and fluvial sediments can accumulate very rapidly (Hayes 1975; Carter 1988). Current understanding of the phenomenon indicates *all* flooding in this region is a result of El Niño, regardless of intensity and/or duration. Ecotones or transitional environmental zones are prone to draught, with short growing cycles, and nine-month dry season making them highly susceptible to El Niño related perturbations. A portion of the mound was buried by sediment with abandonment of stratum 5, reflecting a major El Niño event associated with seismic activity. However, earlier and later evidence of El Niño provide no evidence whatsoever of coastal transgression, but rather large-scale continuous progradation (Staller 1994).

Evidence of prehistoric El Niños may also be derived from variation in the shell species frequencies from archaeological sediments. Shell frequencies were determined by raw counts according to species. When total counts were uneven, the extra bivalve was counted as a complete individual. Once categorized and quantified by species totals were broken down to percentages of minimum numbers of individuals (%MNI). These frequencies were segregated by arbitrary increments with Trench B and on the basis of natural stratigraphic layers elsewhere (Figures 13, 14, Tables 7a-d, 8a-d). Differences in mollusk species in archaeological sediments have been reported as evidence of modifications to aquatic habitats as a result of cultural preference and/or natural factors such as climate (Stein 1990, 1992a, 1992b). Economic shifts and/or climatic change was inferred: 1) a cultural bias for particular food species and, 2) that various taxa inhabited the area during site occupation (Tables 7a-d). Potential factors complicating the statistical reliability, such as sea level changes, geomorphological, and environmental change (all side effects of El Niño) were considered at the outset of this analysis.

A random sample of shells was selected from the surface, identified according to species and assessed by relative frequencies. Only five species were used for comparison, since the low frequencies of some specimens made them statistically insignificant. The

¹¹ Length of shell specimens refers to the measurement of the dorsal to the ventral surface edge.

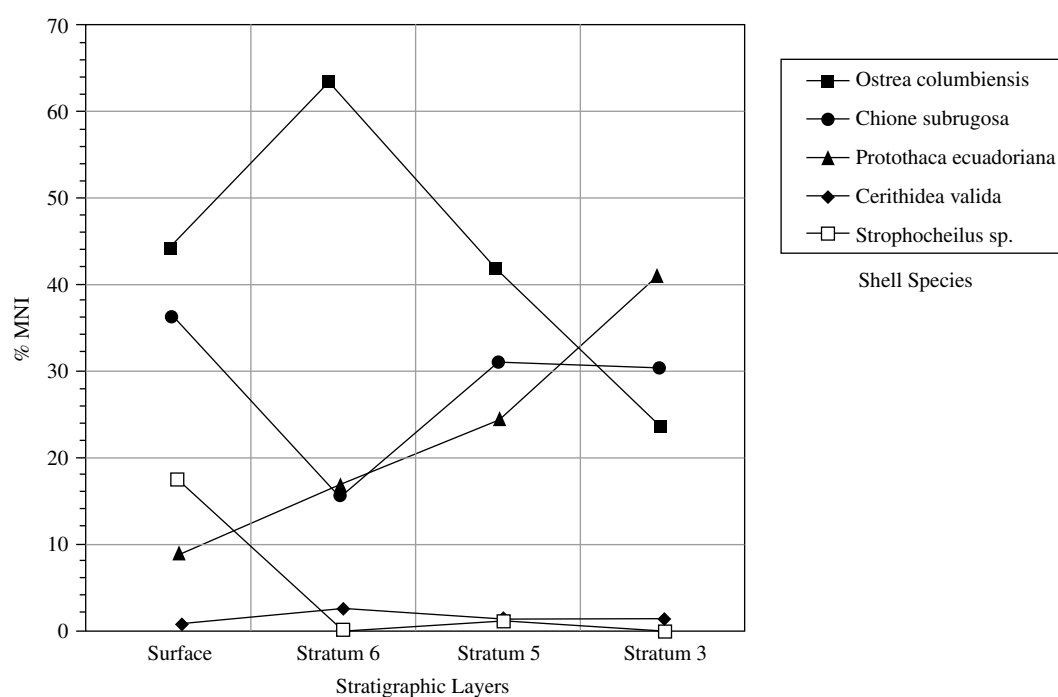


Figure 13. Percentage frequencies by minimum number of individuals (%MNI) by stratigraphic layer.

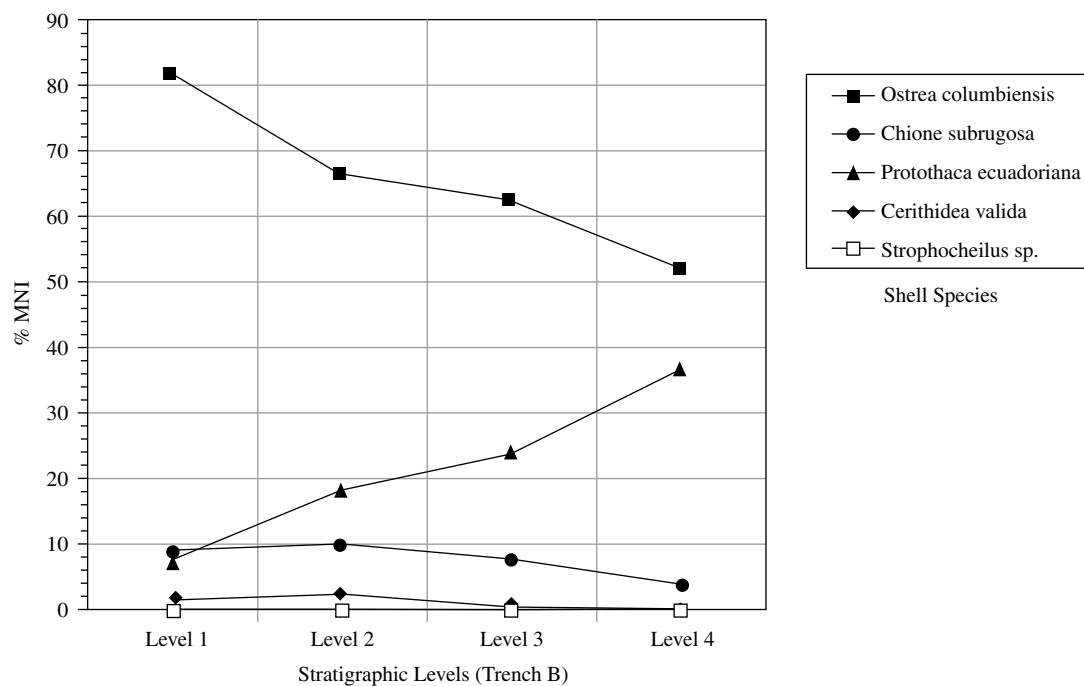


Figure 14. Percentage frequencies by minimum number of individuals (%MNI) by levels or arbitrary increments (Trench B).

raw counts indicated nutritional bias for Oysters (*Ostrea columbiensis* L.), pointed Venus clams (*Chione subrugosa* Wood), and Protothaca clams (*Protothaca ecuadoriana* Rafinesque) (Tables 7a-d). In more inland settlements, Simile Ark (*Anadara*

Table 7A. Shell frequencies: Surface

Species	Total N°	MNI.	Total (%)	MNI. (%)
<i>Ostrea columbiensis</i>	575	288	44.50	44.17
<i>Chione subrugosa</i>	471	236	36.45	36.19
<i>Protothaca ecuadoriana</i>	111	56	8.59	8.58
<i>Strophoecilus</i> sp.	121	121	9.37	16.99
<i>Cerithidea valida</i>	5	5	0.39	0.76
<i>Anadara similis</i>	4	2	0.30	0.30
<i>Anadara grandis</i>	3	2	0.24	0.30
<i>Polinices</i> sp.	2	2	0.15	0.30
Total	1292	712	99.99	99.99

Table 7b. Shell frequencies: Stratum 6

Species	Total N°	MNI.	Total (%)	MNI. (%)
<i>Ostrea columbiensis</i>	8991	4496	64.86	63.34
<i>Chione subrugosa</i>	2196	1098	15.84	15.46
<i>Protothaca ecuadoriana</i>	2367	1184	17.07	16.68
<i>Cerithidea valida</i>	174	174	1.25	2.45
<i>Strophoecilus</i> sp.	4	4	0.02	0.05
<i>Anadara tuberculosa</i>	41	21	0.29	0.29
<i>Anadara similis</i>	77	39	0.55	0.55
<i>Anadara grandis</i>	48	24	0.34	0.33
<i>Anadara obesa</i>	1	1	0.007	0.01
<i>Polinices</i> sp.	45	45	0.32	0.63
<i>Mactra augusta</i>	7	4	0.05	0.05
<i>Dosinia dunkeri</i>	5	3	0.03	0.04
<i>Cirripides</i> sp.	2	2	0.01	0.02
<i>Olivia callosa</i>	1	1	0.007	0.01
<i>Cerithium</i> sp.	1	1	0.007	0.01
<i>Cardita magastrophia</i>	1	1	0.007	0.01
Total	13861	7098	100.65	99.92

Table 7c. Shell frequencies: Stratum 5, Floor 2

Species (Local Name)	Popular Name	Total No.	Total (%)	MNI.(%)
<i>Ostrea columbiensis</i> (osti6n)	Mangrove Oyster	7906	3953	42.01
<i>Chione subrugosa</i> (almeja bajera)	Pointed Venus	5866	2933	31.17
<i>Protothaca</i> sp.(almeja (llorona)	Protothaca Clam	4442	2221	24.33
<i>Cerithidea valida</i> (churo)	Valide Horn	133	133	1.38
<i>Strophoecilus</i> sp. (caracol de monte)	Land Snail	100	100	1.06
<i>Anadara similis</i> (concha prieta)	Simile Ark	130	66	0.70
<i>Anadara grandis</i> (pata de mula)	Grandis Ark	44	22	0.23
<i>Anadara tuberculosa</i> (concha prieta)	Tuberculosa Ark	40	20	0.21
<i>Anadara emarginata</i>	Emarginate Ark	1	1	0.01
<i>Polinices</i> sp.	Moon Shell	22	22	0.23
<i>Dosinia dunkeri</i> (almeja disco)	Disk Dosinia	15	8	0.08
<i>Mactra augusta</i>	Mactra Clam	12	6	0.06
<i>Crepidula marginalis</i>	Slipper Shell	5	5	0.05
<i>Cerithium adustum</i>	Horn Shell	4	4	0.04
<i>Strombus galeatus</i>	Winged Conch	3	3	0.03
<i>Litorina</i> sp. (caracol mangle)	Periwinkle	2	2	0.02
<i>Neritina latissima</i> (caracol multicolor)	Virgin Nerite	2	2	0.02
<i>Olivia callosa</i>	Pacific White Venus	3	2	0.02
<i>Tagelus</i> sp. (nabajas)	Jackknife Clam	2	1	0.01
<i>Polymesoda inflata</i>	Inflata Marsh Clam	1	1	0.01
<i>Nassarius</i> sp.	Dog Whelk	1	1	0.01
<i>Mytella strigata</i> (mejill6n)	Strigata Mussel	1	1	0.01
<i>Olivella</i> sp.(carcol playero)	Dwarf Olive	1	1	0.01
<i>Cirripides</i> sp. (cirr6podo)	Barnacle	1	1	0.01
<i>Spondylus princeps</i> (mullu)	Thorny Oyster	1	1	0.01
Total		18736	9409	99.99

tuberculosa Sowerby), constituted the predominant species during the later formative sites, but were rare in the excavations suggesting interaction. Local fishermen related that when packed wet in

salt water, Simile Ark may be carried for several days without spoiling. Their abundance at formative sites indirectly reflects the importance of the mangrove forest and maritime resources. Their

Table 7d. Shell frequencies: Stratum 3

Species	Total N°	MNI.	Total (%)	MNI. (%)
<i>Ostrea columbiensis</i>	816	408	24.02	23.69
<i>Chione subrugosa</i>	1054	527	31.03	30.60
<i>Protothaca ecuadoriana</i>	1420	710	41.81	41.23
<i>Cerithidea valida</i>	25	25	0.73	1.45
<i>Anadara tuberculosa</i>	41	22	1.20	1.27
<i>Anadara similis</i>	4	2	0.11	0.11
<i>Anadara grandis</i>	6	3	0.17	0.17
<i>Anadara obesa</i>	2	1	0.05	0.05
<i>Anadara labiosa</i>	2	1	0.05	0.05
<i>Polinices</i> sp.	17	17	0.50	1.22
<i>Macra Augusta</i>	4	2	0.11	0.11
<i>Tagelus irregularis</i>	1	1	0.02	0.05
<i>Nassarius</i> sp.	1	1	0.02	0.05
<i>Crepidula marginalis</i>	1	1	0.02	0.05
<i>Oliva callosa</i>	2	1	0.05	0.05
Total	3396	1722	99.89	99.15

absence at La Emerenciana surprising because *Anadara* impressed cambered jars are emblematic of particular diagnostics in the Jelí Phase pottery complex (Figure 15a)¹².

Land snails (*Strophocelium* spp. Pilsbry) called “churo” are adapted to trees or the surfaces of mud flats beyond the limits of the high tide. They are a marginal food source and only exploited during periods of resource scarcity. Valide Horn Conch (*Cerithidea valida* Swainson) are buried in the brackish mud flats at high tide level or entirely out of water on reeds and twigs, and only consumed when more favorable species are unavailable. Both are indicators of environmental stress and anecdotal evidence a multipurpose species, and sometimes ground into lime for chewing coca leaves, or soaking maize kernels. These specimens were restricted to the uppermost 10 cm of stratum 5, thus potential climatic and environmental indicators when the region was undergoing geo-climatic alteration during abandonment (Staller 1994).

Results indicate distinct frequencies relative to the stratigraphic layers. Stratum 3, floor 1 data

indicate *Protothaca* Clam from the mudflats of coastal bays have the highest frequencies, followed by *Venus* Clam from the lagoons and mud flats, and finally by oysters attached to barrier reefs and mangrove prop roots. The environment appears to have been on or near a bay, with mudflats and direct access to a barrier reef. A 19% decline in *Protothaca* clam frequencies in stratum 5 may reflect gradual environmental changes from a bay into a barrier estuary (see Figure 13). Sand dunes and a fossil beach ridge 500 meters north of La Emerenciana are result of an El Niño associated with the deposition of stratum 6. Solely the uppermost layer includes red on white-banded diagnostics (Figure 15b). Reduction in these clams may also be related to an increase in white sandy dune sediments into the coastal bay due to the formation of dune ridges. Increase in sandy sediments in the intertidal zone could have buried the clams. Stability in Pointed Venus Clam frequencies suggest the lagoons scattered through both sides of the coastal streams, and particularly those behind the barrier reef and mangrove islands, continued to be exploited.

Relative changes in the %MNI frequencies and grain size analysis of stratum 6 support a hypothesis of rapid geomorphological change. Smaller oysters and constricted growth rings on most bivalves found on the surface are consistent with longer and higher levels of precipitation, excessive drainage runoff, and swift or sudden introduction of large volumes of fresh water and sandy sediment. El Niño can rapidly bury large areas of the coast in silt and mud, suppressing tidal currents and creating tidal wave asymmetry, which significantly increases the sediment load (Carter 1988; Cane 1983). Slight increases in Valide Horn frequencies on the surface of stratum 6 suggest less favored food species were exploited supporting geo-climatic changes. Grain size, stratigraphic and ceramic evidence suggest the brown silt (stratum 6) may represent a flood deposit that covered over the primary occupation layer in the northwest part of La Emerenciana. This depositional event occurred rapidly, inundating the site, perhaps explaining final site abandonment.

In conclusion, ceramic diagnostics in and on stratum 6 also infer final occupation was temporally restricted to a period of less than fifty years (Staller 1994:137-142). Effects of El Niño-related activity are evident by the total extinction of oysters at Middle and Late Formative sites. Beach ridge formation is also related to El Niño events, which

¹² The only other late Valdivia locality with *Anadara* impressed diagnostics known is the ceremonial center of San Lorenzo del Mate in the Guayas Basin. San Lorenzo is contemporaneous with La Emerenciana and has an eight-meter high earthen mound, the earliest mound of this scale along the Ecuadorian coast. *Anadara* impressed the diagnostics were identified in 2003 in the course of surface survey with my colleague and friend the late Ecuadorian archaeologist Lcdo. Felipe Cruz Mancilla.



Figure 15a. Jeli Phase cambered jars rims with *Anadara* impressed motifs. Such cambered jar fragments were found at all eleven Valdivia sites identified in regional survey and constituted the most commonly identified diagnostics in the excavations at La Emerenciana. The low numbers of *Anadara* spp. shell in the excavations suggest they may have been traded inland. Replication experiments have demonstrated that the impression could only have been made by the species *Anadara tuberculosa* Sowerby.

can rapidly introduce volumes of sandy sediments into an estuary, burying and changing the configuration of ponds and lagoons located along the coast and further inland. A 17% decline in Pointed Venus frequencies may reflect changes in lagoonal and mud flat habitats, perhaps flooding or periodic inundation of coastal lagoons by fresh water incursion. Continued decreases of *Protothaca* Clam may be due to continued silting of an ancient arcuate bay once located due west of the site. One of the most striking changes is an increase in land snails in surface samples, suggesting resource scarcity during final site abandonment (see Figure 13). Once the coastal waters receded this species would have been available in the trees and surrounding mud flats at the high tide line. Significant increases in Pointed Venus suggest this low wave energy estuary and the coastal lagoons quickly returned to normal. Decline in frequencies of *Protothaca* Clam infers continued



Figure 15b. Jeli Phase red on white-banded diagnostics. The red-on-white banded stylistic tradition has generally been associated with the early pottery complexes of southern highland Ecuador and northern highland and coastal Peru. Its earliest appearance pertains to late Valdivia culture in coastal El Oro Province.

environmental change from a coastal bay protected by a barrier reef, to a barrier island estuary. A 20% decline in oysters on the surface also mirrors such environmental changes. Extinction of oysters for a period of over 1200 years after the final stratum 6 abandonment from this region is another indicator of the geological scale of destruction to the estuary and the barrier reef brought on by increases in El Niño related activity (Staller 1994).

Analysis of Shell Remains from Arbitrary Levels (Trench B)

Since the smallest natural unit of analysis was too large to detect changes in the vertical distribution of shell species or artifacts, seashells from Trench B were excavated in arbitrary 15 cm increments, thus some mixing was presumed (see Figure 14). These raw data are not directly comparable to those from natural layers. Shell frequencies in the midden corresponded exclusively to strata 5 and 6. Trench B %MNI oysters frequencies indicate a 30% increase in the uppermost levels and a 20% decrease in *Protothaca* Clam, consistent with shell frequencies for strata 5 and 6. Pointed Venus Clam slightly increases in level 2, but then decreases in uppermost level 1 to less than 10% of the total,

while frequencies from natural stratigraphic layers average about 25% in stratum 5. Such differences may be related to sample size (see Tables 8a-d). Pointed Venus clams would be more intensely exploited during periods of resource scarcity. The frequencies further indicate lagoonal environments remained stable, although they were a relatively minor resource (see Figures 13 and 14).

Trench B shell frequencies are reliable indicators of environmental change, as their overall distribution is similar to %MNI from strata 6 and 5 (see Tables 7b-c). Congruence in frequencies for stratum 6 and those from level 1 may indicate that the uppermost level was associated with the final occupation. Oyster counts from both stratigraphic layers and arbitrary levels indicate they were the most highly exploited and even more intensely exploited during stratum 5 abandonment. Results indicate shell in the upper levels of Trench B represent a single continuous deposit, the product of the final occupation of stratum 5. The uppermost 15 cm may be associated with the brief stratum 6 occupation. However, stratum 6 was not visible in cross section, due to the large quantities of bivalves and sherds (Staller 1994:345). Artificial retaining walls and descending terraces were recorded in the Trench B excavation. The construction of artificial terraces to demarcate sacred space is common in the Andes particularly early *huacas* in the highlands (Hocquenghem *et al.* 1993). Various lines of evidence indicate two major El Niño events associated with site abandonment dated to c. 2150 B.C. and another dated to c. 1450 B.C. resulting in final abandonment. Auxiliary lines of evidence from lake core data from the Cuenca valley and coastal Peru also support the timing and interpretations repeated abandonment associated with ancient El Niño events. Other supporting lines of climatological and geomorphological evidence from contemporary El Niño events also support the scale of geological and geomorphological modifications presented herein. Environmental evidence including modern analogues of devastation wrought by El Niño upon this region and provides a basis for clearer understanding of its impact upon prehistoric populations.

Recent El Niño Related Effects in Southern Coastal Ecuador

The 1982/83 El Niño was the most devastating and destructive ever recorded in a two hundred

year historic record (Rasmussen 1985; Cane 1983; Wrtki 1982; Philander 1983; Caviedes 1984; Arntz 1986; Wells 1987). Precipitation levels along the Gulf of Guayaquil totaled 3650mm, where the average is 1000mm (Thayer and Barber 1984:4, 10; Arntz 1986:4). Mega-El Niños usually result in intense rainfall accompanied by flooding well beyond the rainy season. During the 1982/83 event this coastal region experienced a 20 to 40 cm sea level rise and remained at those levels for months, followed by another sea level rise several months later (Arntz 1986:7). Between December 1982 and June 1983, seawater levels rose approximately two meters burying most of the mangrove trees and shrubs in salt, sand, and fine mud (*ibid.*:22). Even more severe destruction befell this region during the 1997/98 event (McPhaden 1999). The 1997/98 Mega-El Niño was comparable in intensity, yet its affects were different. Its onset was already apparent in late June/early July of 1997 rather than late December. Interactions with local fishermen along the northern shore of the Santa Elena Peninsula indicated fish capture included essentially unknown or never before seen warm water fish species, and sea birds, particularly pelicans, were dying off in great numbers¹³. McPhaden (1999:950) points out the SST associated with the 1997/98 event developed at such an alarming rate from June to December, that new monthly records were set along the eastern Pacific – this based upon records going back to the middle of the 19th century. At the height of the Mega-El Niño in December, the SST anomalies averaged nearly 4°C, the largest variation ever recorded. The SST remained anomalously high at 29°C in the Eastern Pacific until mid-May 1998, when trade winds abruptly returned and cold subsurface waters upwell to bring on the conditions for the La Niña to follow (McPhaden 1999:951; Vuille *et al.* 2000). The study area was submerged under two meters of water for a period of several weeks creating exceptional hydrological conditions that carried more sediments into the barrier island estuary in a

¹³ While in the Santa Elena Peninsula in late June 1997 fishermen indicated to me that tuna had moved away from the shoreline into deeper waters, and fish species foreign to this region appeared, some of which they had never seen before. During El Niños warm water species from the north migrate south into this region. Tuna swim to deeper, cooler waters, and the shoreline was scattered with the carcasses of a number of bird species, turtles and fish, an indicator that this was a mega event.

Table 8a. Shell frequencies: Trench B Stratum 5, Level 1

Species	Total N°	MNI.	Total (%)	MNI. (%)
<i>Ostrea columbiensis</i>	4453	2227	81.78	80.71
<i>Chione subrugosa</i>	460	230	8.44	8.33
<i>Protothaca ecuadoriana</i>	410	205	7.52	7.43
<i>Cerithidea valida</i>	51	51	0.93	1.84
<i>Strophoecilus</i> sp.	1	1	0.01	0.03
<i>Anadara tuberculosa</i>	21	11	0.38	0.39
<i>Anadara similis</i>	11	6	0.20	0.21
<i>Anadara grandis</i>	20	10	0.36	0.36
<i>Polinices</i> sp.	17	17	0.31	0.61
<i>Macra augusta</i>	1	1	0.01	0.03
Total	5445	2759	99.94	99.94

Catalogue No. [4679] [4682] [4683] [4688] [4692] [4693].

Table 8b. Shell frequencies: Trench B Stratum 5, Level 2

Species	Total N°	MNI.	Total (%)	MNI. (%)
<i>Ostrea columbiensis</i>	2821	1411	68.20	66.71
<i>Chione subrugosa</i>	411	206	9.93	9.73
<i>Protothaca ecuadoriana</i>	758	379	18.32	18.00
<i>Cerithidea valida</i>	65	65	1.57	3.07
<i>Strophoecilus</i> sp.	1	1	0.02	0.04
<i>Strombus galeatus</i>	4	4	0.08	0.18
<i>Anadara tuberculosa</i>	25	13	0.60	0.61
<i>Anadara similis</i>	22	11	0.53	0.52
<i>Anadara grandis</i>	8	4	0.19	0.19
<i>Anadara obesa</i>	2	2	0.04	0.09
<i>Polinices</i> sp.	20	20	0.48	0.94
<i>Macra augusta</i>	1	1	0.02	0.04
<i>Nassarius</i> sp.	1	1	0.02	0.04
<i>Crepidula marginalis</i>	1	1	0.02	0.04
Total	4140	2119	99.94	99.93

Catalogue No. [4680] [4684] [4690] [4694] [982] [983] [1023] [1024].

Table 8c. Shell frequencies: Trench B Stratum 5, Level 3

Species	Total N°	MNI.	Total (%)	MNI. (%)
<i>Ostrea columbiensis</i>	596	298	64.57	63.13
<i>Chione subrugosa</i>	76	38	8.23	8.05
<i>Protothaca ecuadoriana</i>	225	113	24.37	23.94
<i>Cerithidea valida</i>	5	5	0.54	1.07
<i>Anadara tuberculosa</i>	2	1	0.21	0.21
<i>Anadara similis</i>	2	1	0.21	0.21
<i>Anadara grandis</i>	4	2	0.43	0.43
<i>Polinices</i> sp.	9	9	0.97	1.91
<i>Macra augusta</i>	2	1	0.21	0.21
<i>Cardita magastroph</i>	1	1	0.10	0.21
<i>Tellina ecuadoriana</i>	1	1	0.10	0.21
Total	923	472	99.94	99.98

Catalogue No. [4681] [4685] [4691] [4695].

Table 8d. Shell frequencies: Trench B Stratum 5, Level 4

Species	Total N°	MNI.	Total (%)	MNI. (%)
<i>Ostrea columbiensis</i>	645	323	53.57	53.03
<i>Chione subrugosa</i>	57	29	4.73	4.76
<i>Protothaca ecuadoriana</i>	453	227	37.62	37.27
<i>Anadara tuberculosa</i>	11	6	0.91	0.98
<i>Anadara similis</i>	18	9	1.49	1.47
<i>Anadara grandis</i>	11	6	0.91	0.98
<i>Anadara obesa</i>	2	1	0.16	0.16
<i>Polinices</i> sp.	4	4	0.33	0.65
<i>Macra augusta</i>	1	1	0.08	0.16
<i>Crepidula marginalis</i>	1	1	0.08	0.16
<i>Tellina ecuadoriana</i>	1	1	0.08	0.16
Total	1204	609	99.96	99.78

Catalogue No. [4713] [4714]

few days than would normally be deposited over a period of several months or even years of average inflow (Wells 2001:150). Conditions similar to these no doubt resulted in the deposition of uppermost stratum 6 at La Emerenciana, and the partial burial of the earthen mound.

Archaeological studies of ancient El Niños suggest they have a dramatic effect on the long-term cultural development of societies living in these coastal regions and those to the south (Moseley 1983, 1987; Moseley *et al.* 1981; Moore 1991; Keefer *et al.* 2003; Sandweiss 2003; Sandweiss *et al.* 2007, 2009). These data provide direct evidence that El Niño must be taken into account to interpretations of long-term cultural adaptation to both coastal and highland Andean regions (Keefer *et al.* 2003). Coastal stability and rate of geomorphic change are profoundly affected by El Niño cycles on a geological scale (Murphy 1939; Caviades 1984; Gill and Rasmussen 1983).

The distinctive side effects of such climatic events include unpredictable downturns of aquatic resources. Mega-El Niños can be both beneficial and detrimental to a fishing economy. Since mollusk and fish are displaced into shallower, more accessible waters, there is an initial *increase* in “catches” of some species (Arntz 1986:18-20). Growth ring studies on shellfish and aquatic populations indicate an initial rapid growth followed by constricted growth rings on bivalves, and higher than normal densities of scallops and other invertebrates (Arntz 1986; Rollins *et al.* 1986, 1987). Intertidal and subtidal mussel beds and kelp forests are usually destroyed, resulting in widespread mortality that may sometimes

take years to recover (Svenson 1946; Barber and Chavez 1983; Rasmussen and Wallace 1983; Arntz 1986:Figure 8). Such devastation is evident throughout the food chain (Holmgren *et al.* 2001).

El Niño can modify pH and algae levels, creating red tide as wetland lagoons and ponds turn to a maroon red color. After a period of several days, large numbers of fish and related aquatic fauna perish due primarily to lack of oxygen. Such sudden widespread mortality tempered a strong reliance upon such resources (Staller 1994). With the increased frequency of events during Middle and Late Formative times, the adaptive response favored a short-term dependence upon agricultural food crops. Modern analogues suggest that such events were central to scheduling and other adaptive adjustments to resource scarcity (Binford 1968, 1989, 2001; Flannery 1986). El Oro farmers refer to El Niño events as *inviernos buenos* (good winters) because after the rains subside, the hillsides are lush, green, and dense with vegetation, crops may be cultivated year-round, creating dramatically beneficial short-term consequences for agricultural specialization (Barber and Chavez 1983; Rasmussen 1985; Moseley 1983, 1987). The ecologically diverse environment of this region, appears to select for short-term survival strategies, e.g., increased specialization, and scheduling distinct collective strategies of resource exploitation, to take advantage of changing environmental and climatic conditions (Flannery 1973, 1986). The ancient adaptation was by necessity always flexible ranging from year round maritime resource exploitation to periods of resource dependence upon food crops enough to withstand periods of scarcity or environmental stress brought by seemingly random periods of drought (Staller 2001a:213, figure 5). During the Middle Formative period c. 1450-1000 B.C. there were dramatic reductions in site size, population density as well as a total absence of ceremonial centers¹⁴.

Late Formative period settlement patterns suggest a greater dependence upon terrestrial plants, and a preference for floodplain settings however the coast is reoccupied marine resources intensively exploited for the rest of the pre-Hispanic sequence (Staller 1994, 2000). A total absence of prehistoric settlements along the shoreline during the Middle and Late Formative period c. 1400-300 B.C. is suggested to be related to ecological transformations and geomorphological changes brought on by El Niño (see Table 3).

Concluding Remarks

In conclusion, the shell frequencies, stratigraphic, geomorphological, and geological evidence, as well as the formative settlement patterns, all point to the existence of ancient El Niño events in southern coastal Ecuador in the Early Formative. Direct and indirect evidence suggest repeated site abandonment at La Emerenciana was related such climatic events, and that human adaptation was permanently modified by their increased frequency over time. Initial site abandonment was in response to a Mega-El Niño event dated to ca. 2150 B.C. associated with fossil beach ridge formation and reoccupation ¹⁴C dated 2200 to 1450 B.C., and final abandonment dated to ca. 1450 B.C. Final abandonment is associated with an earthquake and a short-lived reoccupation. These internally consistent supporting lines of evidence for a Mega-El Niño resulting in initial abandonment at La Emerenciana and fossil beach ridge formation tentatively dated to 4040 B.P. (~4400 cal. B.P.) (Rodbell *et al.* 1999:figure 3). El Niño events associated with final site abandonment was also intense and destructive, possibly associated with, or closely followed by seismic activity resulting in permanent abandonment of the ceremonial center at ~1450 B.C. Changes in settlement patterning at this time involved a shift to knoll top settlement localities situated along the coastal streams. Perhaps the most dramatic geological evidence of this later event is further fossil beach ridge formations and the subsequent extinction of oysters, presumably an important source of protein, from this region for over a millennium.

¹⁴ Formative period ceremonial centers in coastal Ecuador are distinguished by earthen mounds of varying sizes (Staller 2000; Lunniss 2008). The earliest oval and circular earthen mounds pertain to Valdivia culture and have also been identified at sites of Real Alto, San Isidro, San Lorenzo del Mate (Staller 2001a; see also Lathrap *et al.* 1977; Marcos 1988; Marcos *et al.* 1976). Later ceremonial mounds and elite earthen platforms are of increasing scale and more geometric in form. Earthen platforms standing thirty meters high and sixty meters at the base pertaining to the Milagro Quevedo (AD 1100-1531) culture were identified in regional survey

north of Santa Rosa along the Buernavista River (Staller 1994).

Acknowledgments

I wish to express my sincerest gratitude to the Fulbright Association in Washington D. C. for funding this research through a research scholarship and a dissertation improvement grant. I am deeply indebted to the late Olaf Holm, director of the Museo Antropológico del Banco Central del Ecuador in Guayaquil, and the Department of Anthropology of Southern Methodist University for their sponsorship and generous and enthusiastic support. Sincerest thanks to Alfred Siemens

(University of British Columbia) for sharing his insights on the coastal hydrology and geography. Special thanks and sincere appreciation to Sr. Hector Washco of Arenillas, for providing the initial scientific identification of the shellfish species, and access to many important Ecuadorian botanical and biological publications on the local ecology. All stratigraphic and archaeological interpretations of the various layers and archaeological context are the responsibility the author as is the presentation and interpretation of these data.

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