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DRASTIC and GOD vulnerability maps of the Cabril River Basin, Portugal

DRASTIC e GOD: mapas de vulnerabilidade da bacia do rio Cabril, Portugal

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Abstract

The main purpose of this study is to compare two methodologies, DRASTIC and GOD, applied to the basin of the Cabril River in determining groundwater vulnerability to pollution. The justification for the research work in this area relates to the fact that we are facing a basin adjoining the town of Vila Real (Portugal), with a propensity to expand into this area, and where there are many springs, both near the base of the basin and in hillside areas.

As a result of the application, the GOD and DRASTIC maps were obtained. To better compare these final vulnerability maps, the conversion of the ranges of DRASTIC indices to a normal-ized scale from 0.0 to 1.0 value was performed, staying in the same rating scale as the GOD method. It was observed, through the GOD method, that the area of the basin has mostly low or negligible vulnerability, related to the high depth of the groundwater level and also to the low degree of confinement. Moreover, the results of the DRASTIC method indicate that most of the basin presents medium to low vulnerability, especially in areas of higher altitude. However, both methods make high vulnerability correspond to areas covered by alluvium.

It may be concluded from this study that the production of results using different methods of analysis, although with similar objectives, may differ in certain locations. Thus, the outlook for the future is that the vulnerability mapping techniques ought to include physical methods, improving the credibility of these studies.

Keywords: *aquifer; vulnerability, DRASTIC, GOD, GIS*

Resumo

O principal objetivo desse estudo é a comparação de duas metodologias, DRASTIC e GOD, aplicadas à bacia hidrográfica do rio Cabril, na determinação da vulnerabilidade das águas subterrâneas à poluição. A justificativa do trabalho, nessa região, prende-se ao facto de estarmos perante uma bacia confinante com a cidade de Vila Real (Portugal), com propensão para se expandir para essa área e onde existem muitas nascentes, tanto junto à base da bacia, como nas zonas de meia encosta.

Como resultado da aplicação, obtiveram-se os mapas GOD e DRASTIC. Para melhor comparar esses mapas finais de vulnerabilidade, converteram-se os intervalos dos índices DRASTIC para uma escala normalizada de 0,0 a 1,0 valores, ficando em igual escala de classificação do método GOD. Constatou-se, através do método GOD, que a área da bacia apresenta, na sua maior parte, vulnerabilidade baixa ou desprezível, relacionada com a elevada profundidade do nível de água subterrânea e, também, com o baixo grau de confinamento. Por outro lado, os resultados do método DRASTIC indicam que grande parte da bacia apresenta vulnerabilidade média a baixa, principalmente em zonas de maior altitude. No entanto, ambos os métodos

apontam para a existência de vulnerabilidade elevadas nas zonas cobertas por aluviões.

O que se pode concluir do presente estudo é que a produção de resultados por métodos de análise distintos, embora com objetivos semelhantes, pode ser diferente em determinados locais. Desta forma o que se pode diagnosticar no futuro é que as técnicas de mapeamento da vulnerabilidade deverão contemplar métodos físicos, tornando estes estudos mais credíveis.

Palavras Chave: *Aquífero, vulnerabilidade, DRASTIC, GOD, GIS*

1. Introduction

The world population began to worry about the groundwater resources only about two decades ago. Those who depended on these resources for drinking water consumption did nothing to maintain the quality of water, much less evaluated the risk of contamination (Foster 1987, Foster *et al.* 2006). With the introduction of the term “groundwater vulnerability to pollution” numerous definitions, methodologies and skills about it followed, and it resulted in two findings: intrinsic vulnerability, which results from the characteristics of the context (aquifer type and coverage, permeability, depth, recharge, among others), without considering the impact of pollutant substances, and specific vulnerability that gives superiority to the type and amount of polluting agents (Auge 2004).

There are several methodologies to determine aquifers' vulnerability to pollution. However, the methods used in this study were the DRASTIC and the GOD. The DRASTIC method (Aller *et al.* 1985) seeks to systematize the determination of the potential of the pollutant to reach the saturated zone and, for such, it considers 7 factors: Depth to the water table; Net Recharge; Aquifer material; Soil type; Topography; Impact of the unsaturated zone; and Hydraulic Conductivity. In turn, the GOD method determines the intrinsic vulnerability; hence, it does not consider the type of pollutant. It is based on the designation of indices between 0 and 1 with the variables, Ground water occurrence; Overall aquifer class; and Depth to groundwater table (Foster 1987, Foster *et al.* 2006). The DRASTIC model has been applied in the study of aquifers' vulnerability to pollution, and references may be made to cases such as

the Pampeano aquifer, located north of the province of Buenos Aires (Reynoso *et al.* 2005), or the Oued Laou aquifer, located in the northeast of Morocco and considered one of the most important in this region (Salhi *et al.* 2007). This latter aquifer has been threatened by numerous sources of contamination over the past years, and the situation is increasingly aggravating, in the absence of any plan of environmental protection. In this context, Salhi *et al.* (2007) carried out a study on the development of vulnerability maps of the aquifer, which seeks to function as a useful tool for administrators of water resources when making decisions. As for the application of the GOD method in the study of vulnerability to groundwater pollution, reference is made to the work carried out by Barboza *et al.* (2007) in Ponta de Fruta (Vila Velha, Espírito Santo, Brazil), a region with poor sanitation conditions.

There are cases when different methodologies are applied to a given region, comparing the results. In 2003, Gogu *et al.* (2003) tested five methods in assessing the intrinsic vulnerability of a karst aquifer located in the region of Condruz (Belgium). The methods adopted were: EPIK, GOD, “German Method”, ISIS and DRASTIC. After comparing the results, it was concluded that the GOD method, the “German Method” and the ISIS method ascribe more than half of the area under study a high vulnerability. Moreover, DRASTIC and EPIK methods indicate that most of the area under study has moderate vulnerability. Similar research studies have been carried out in other hydrographical basins, in order to assess the vulnerability of groundwater to pollution caused by anthropogenic pres-

ures, such as the works by Mendoza and Barmen (2006), which used the DRASTIC and GOD methods in the basin of Artiguas River, Nicaragua. Other studies made comparisons between methods of analysis of the aquifers' vulnerability to pollution, such as the ones by Santos & Pereira (2011) and Tavares *et al.* (2009) in Brazil; Vias *et al.* (2005) and Martínez-Bastida *et al.* (2010) in Spain; Draoui *et al.* (2008) in north-eastern Morocco; Roberta *et al.* (1998) in Italy; Ibe *et al.* (2001) in Nigeria; Pacheco and Sanches Fernandes (2013) in the River Sordo basin, Portugal; etc., being also evaluated, in these cases, the risk of pollution.

Basically, the calculation made by these methodologies, and its mapping, are based on a GIS (Geographical Information System), as illustrated by the work of Daniela (1999).

The pressing need for a country planning policy based on scientific methodologies and criteria has to be effective. The decision support systems, as a tool to assist spatial planning, have received some attention in the area of aquifers' vulnerability to pollution, and specifically through the DRASTIC model, as occurred in Hungary (Leone *et al.* 2009). Furthermore, in the United States of America, particularly in Colorado, the contamination of aquifers with nitrates has been the subject of studies and publications, being used as necessary and effective methodologies to facilitate policy decisions concerning the protection of groundwater (Cepelch *et al.* 2004). The present research study fits into this same context, its objective being to lay a foundation for the sustainable management of water resources in the region of Vila Real.

Study area

Cabril River flows from west to east and is located west of the city of Vila Real, in northern Portugal.

It flows into the Corgo River, at the southern end of town (Figure 1), about 1.5 km from the city centre. Cabril

River is about 10.5km long and has a basin area of about 63.8 km² (Borges 2006). The altitudes in its basin vary

between 304 m and 1325 m, with slopes between 0 and 106%. It is, therefore, a relatively rugged topography, of the

type of medium altitude mountains. In areas of lower elevation, specifically in the areas surrounding Cabril River

and its tributary streams, the land is little rugged.

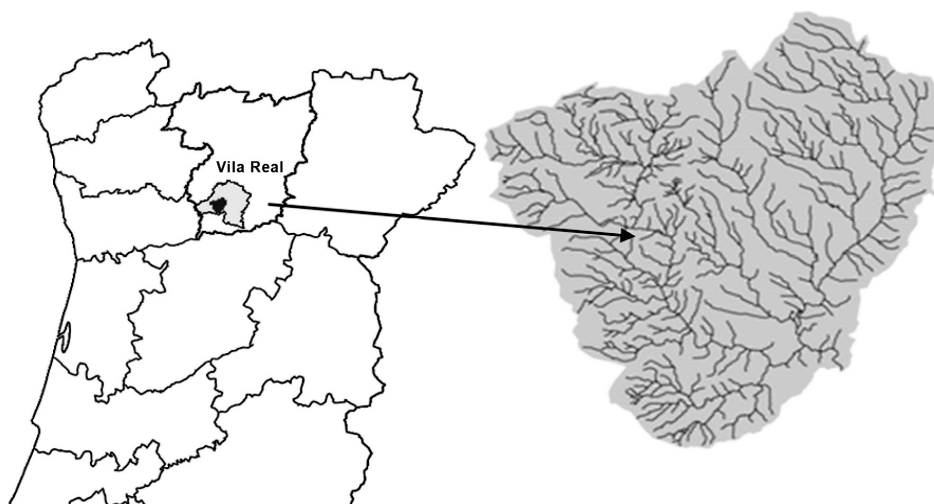


Figure 1
Location of the Cabril River hydrographical basin in the municipality of Vila Real.

In the far northeast, the geology of the basin is characterized mainly by metasedimentary units from the Ordovician period (Armorican quartzite). The southeast quadrant is covered by modern alluvium, valley bottom sand-clay deposits, and undifferentiated river terraces. In the northwest sector, the granite of Lamas de Olo emerges, as well as various units of the Schist-Greywacke Complex (Douro Group, Middle to Upper Cambrian). Emphasis is also put on

the presence of quartz, aplite, pegmatite and aplite-pegmatite veins, especially in the upper sector of the hydro-graphical basin. The remaining sectors are characterized by the emergence of various types of granite. The fracture of the basin is dominated by the following families: i) NNE-SSW to NE-SW and ii) ENE-WSW (Matos 1991). The schist umbric Lithosols and the granite umbric Lithosols are the dominant soil units in the basin under study, as well as soils derived from ter-

aces. The hydrograph-ical basin of Cabril River has a moderate extent which is fit for agriculture, pastures and forests. Regarding the use of the soil, a part of the basin is characterized by intensive irrigation, watered prairie and pinewood. The climatic characterization, according to data collected at the udometer stations closest to the basin, shows a rainfall weighted average close to 1236 mm and an average monthly temperature varying between 10°C and 15°C.

2. Methodology

DRASTIC Application

Parameter (D) represents the thickness of material that a pollutant will have to go through before reaching the

aquifer. That thickness was obtained by measuring the depth of water in wells (Table 1), converting afterwards those

values into vulnerability indices. The number of wells inventoried is presented in Figure 2(a).

Wells	Depth of Water (m)	Wells	Depth of Water (m)	Wells	Depth of Water (m)	Wells	Depth of Water (m)
1	5.40	9	2.60	17	8.00	25	4.50
2	1.30	10	2.30	18	2.00	26	2.10
3	10.50	11	0.30	19	8.70	27	0.80
4	5.70	12	7.30	20	18.00	28	1.30
5	8.40	13	6.90	21	10.60	29	4.80
6	2.60	14	7.40	22	3.50	30	3.50
7	4.80	15	5.70	23	2.60	31	1.90
8	3.90	16	4.40	24	7.40	32	7.20
						33	1.20

Table 1
The depth of water in wells

They are not evenly distributed, and the absence of wells in considerable land extents north and west of the basin may be observed. These land extents are located in areas where high altitudes predomi-

nate, where there are also high rainfalls, where water withdrawals for irrigation are essentially spring-type and not well-type. From the data collected (location of each well and its depth of hydrostatic

level - pNHE), the pNHE map for the basin and the corresponding D index map were obtained by interpolation (Figure 2(b)), using, for that, the ArcGIS software (ESRI 2007).

Parameter (R) represents the annual recharge of the aquifer. In general, the higher the re-charge, the greater the potential for the contamination of an aquifer. However, a high volume of recharge also causes dilution of pollutants, reducing their pollution effect. To access parameter R, a hydrological modelling

software called ArcSWAT was used (Winchell *et al.* 2008). In the context of that modelling, the recharge is calculated at the scale of the sub-basin by a balance of water and it is, therefore, necessary to have information on the daily rainfall over the basin. The two udometer stations with this type of data are located outside the

basin. The range of data adopted in the model-ing encompassed the 1959-1967 period. After outlining the sub-basins map (Figure 2(c)), determining the respective values of annual recharge, and applying the conversion table relating to this pa-rameter, the spatial distribution of R was obtained, as illustrated in Figure 2(d).

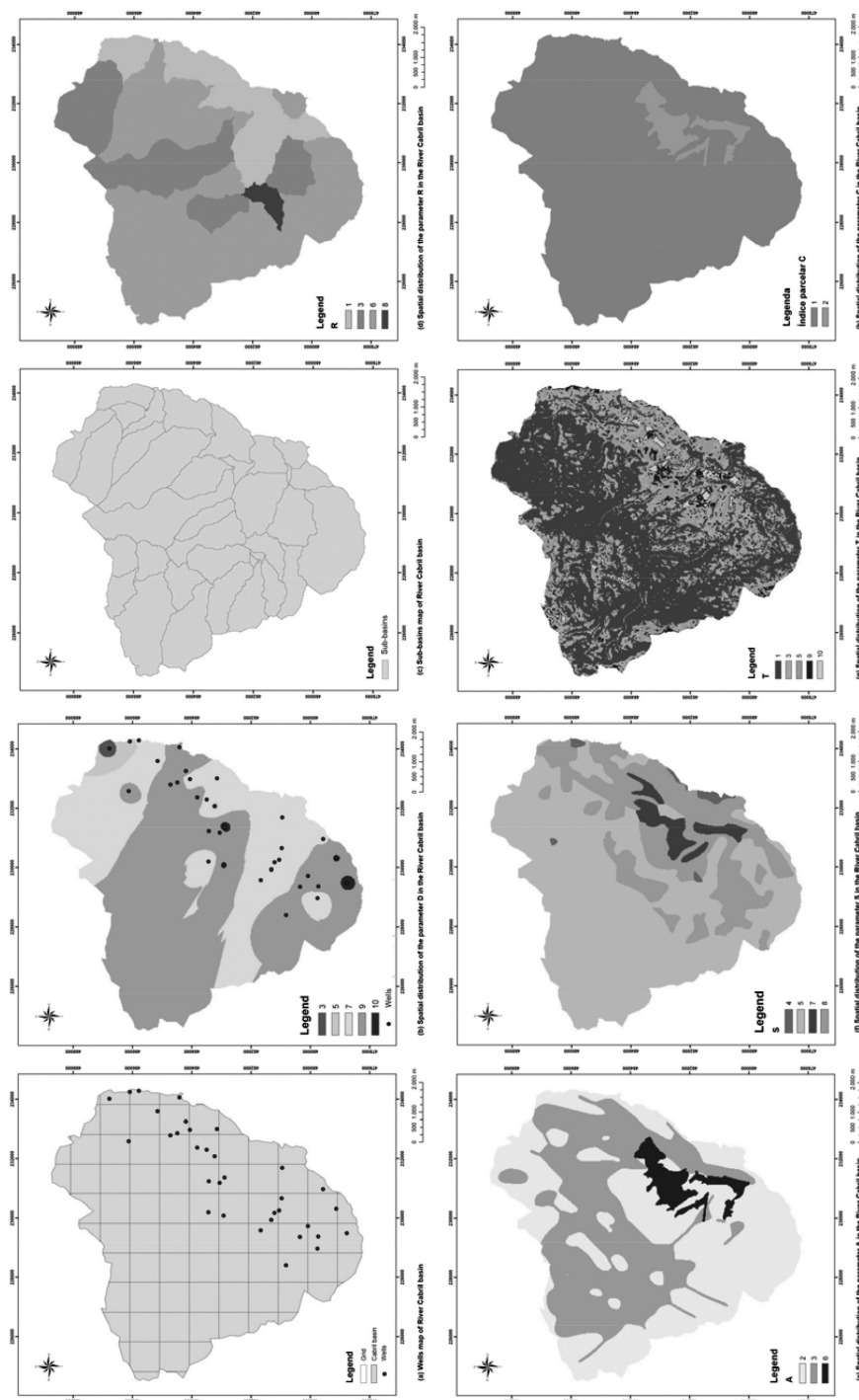


Figure 2
Spatial data of Cabril River Basin: wells locations (a), depth of hydrostatic level (b), sub-basins (c), annual recharge of the aquifer (d), material of the aquifer (e), soils (f), topography (g) and hydraulic conductivity (h).

The material of the aquifer (A) describes the mitigation capacity of the rocky soil (porous, fractured or karstified), consolidated or not, which serves as an aquifer. In general, the coarser-grained rocks or those with more fractures or

cavities have higher hydraulic conductivity and, there-fore, less mitigation capacity. Moreover, these characteristics of the aquifer material also influence the pathways used by pollutants, that is, their hydrodynamic dispersion. In the case under

study, the map of parameter A combines lithology and fracture. In that context, the values for this parameter were defined on the basis of the geological map and the fracture density map. The conversion of the lithological and structural data into

values of the A index was based on the work by Pacheco *et al.* (2004), carried out in a basin adjacent to the Cabril River basin (Sordo River Basin). Initially, the lithotypes were distinguished in Terraces / Alluvial and Granite / Paleozoic Metasediments. On the basis of the geological map of the basin, the first group was assigned the index 6. To the second group, considering the fracture densities calculated on the basis of geological structures marked on the geological chart (0 to 9.36 km/km²), the indices 2 (class 0 to 8 km/km² from Pacheco *et al.* 2004) and 3 (grade 8-16 km/km²) were assigned. Once the variation ranges of index A were established, the final map of the parameter illustrated in Figure 2(e) was drawn up.

The soils (parameter S) control the concentrations of pollutants in the water, and hence their adverse effects, through their permeability characteristics and capacity to produce colloidal complexes which determine their retention power. According to Pacheco *et al.* (2004), and based on textural characteristics of soils in the hydrographical basin of Cabril River, a correlation between soil typologies and parameter S was established. The soil units present in the basin are the lithosols (area: 43 km²), sometimes sandy (0.4 km²), the terrace anthrosols (17.1 km²), fluvio-

soils (1.6 km²) and cambisols (1.1 km²). Conversion into vulnerability indices of parameter S resulted in values of 5, 8, 8, 7 and 7, respectively (Figure 2(f)). Urban areas have been assigned the index 4.

The parameter topography (T) assesses the influence of the relief of the area under study on the fate of pollutants spilled over planet Earth. This parameter sets the probability of a pollutant to flow freely or remain at the surface long enough to infiltrate. The hydraulic gradient, the direction of flow, and the development of soils are related to the slope. Usually, higher slopes involve greater hydraulic gradients, and naturally higher velocities of groundwater flow. On the basis of the Terrain Digital Model of the Cabril River Basin, its slope map was drawn up, using, for that, tools from ArcGIS. After the production of this map, slopes were converted into indices T_i, according to the DRASTIC conversion table (Figure 2(g)).

The unsaturated zone (parameter I) is the horizon underlying the soil and above the ground-water table. The type of material that this area is composed of influences the time of contact of the environment with the pollutant, boosting the development of biodegradation, neutralization, mechanical filtration, chemical reaction, volatilization and

dispersion processes. In the unsaturated zone, fracture is particularly important, since it promotes the vertical percolation of the pollutant to the aquifer. Due to lack of data on the unsaturated zone of the Cabril River Basin, a typical index of 4 was assumed for the entire basin, corresponding to metamorphic/igneous rocks, as the basin is composed mostly of this type of rocks.

The hydraulic conductivity parameter (C) describes the ease with which the water passes through the pores, cracks, cavities or layering of the aquifer, being conditioned by the quantity, volume and connectivity of the voids in the middle. For the basin under study there are no data on this parameter. For the basin which is contiguous of Sordo River, Pacheco and Van der Weijden (2007) determined values of $3.1 \pm 2.9 \times 10^{-6}$ m/s (0.27 ± 0.25 m/d), for the granites and metasediment rocks. These values are much lower than the minimum value of class C = 1, which is 4.6 m/day. For that reason, based on the geological mapping of the area, the following values were assigned: index 2 to alluvium; index 1 to granites and paleozoic metasediments (Figure 2(h)).

In accordance with the model, the DRASTIC vulnerability index was calculated by the expression (1):

$$\text{DRASTIC} = D_p \times D_i + R_p \times R_i + A_p \times A_i + S_p \times S_i + T_p \times T_i + I_p \times I_i + C_p \times C_i \quad (1)$$

where, i represents the index assigned to the parameter in question and p the weight assigned to that same parameter.

The weight of each parameter reflects its importance relative to the remaining parameters. It ranges from 1 to 5

(quantification of the DRASTIC index to general situations of vulnerability), resulting in (2):

$$\text{DRASTIC} = 5 \times D_i + 4 \times R_i + 3 \times A_i + 2 \times S_i + 1 \times T_i + 5 \times I_i + 3 \times C_i \quad (2)$$

GOD Application

The determination of the GOD vulnerability index involves identifying the degree of confinement of the aquifer groundwater, quantified by a value in the range from 0.0 to 1.0, the lithological characterization of the horizon located

above the saturated zone of the aquifer, categorized by a value in the range from 0.4 to 1.0, and, finally, the estimation of the depth of the groundwater table, converted into values in the range from 0.6 to 1.0. The index is the product of the previous

values, that is, GOD = Groundwater occurrence (G), Overall lithology of aquifer or aquitard (O) and Depth to groundwater table (D) (Foster *et al.* 2006). GOD values are matched with a qualitative scale of vulnerability, as presented in Table 2.

0.0	0.1	0.2	0.3	0.4	0.5	0.6	0.7	0.8	0.9	1.0
Negligible	Low	Medium	High	Extreme						

As there is no specific information, the degree of confinement of the aquifer in the hydrographical basin of Cabril River (parameter G) was established as follows: with no confinement in areas where soil cover is composed of terraces

and alluvium; with semi-confinement or confinement in the remaining sectors, depending on the density of rocks fracture. In the latter case, the smaller the fracture, the greater the aquifer's degree of confinement. To determine pa-

rameter G, in a first stage, two support maps were produced, using ArcGIS: (a) a G1 map, in which only the existence of free aquifers was considered. In this map, the area of terraces and alluvium was represented by the value 1 (free

Table 2

GOD values and corresponding classes of vulnerability. Adapted from Vogel (2008).

aquifer) and the remaining areas by the value zero (with the meaning of un-rated); (b) a G2 map, in which the existence of confined aquifers and semi-confined was considered, in line with density of rocks fracture. Three density ranges were arbitrarily considered. The limits of those ranges were defined on

the basis of the frequency of densities, each range representing one third of those frequencies: $[< 1.25 \text{ km/km}^2]$, $[1.25\text{--}3.45 \text{ km/km}^2]$, $[> 3.45 \text{ km/km}^2]$. Each range was assigned an index representing confined or semi-confined aquifers, respectively 0.2, 0.3, and 0.4. To draw the map of parameter G (Fig-

ure 3(a)) at each point of the basin the maximum value of G1 and G2 maps was selected. This option allowed that the zone of terraces and alluvium was invariably identified by the index 1 and that the remaining areas were identified by a value between 0.2 and 0.4, depending on the fracture density.

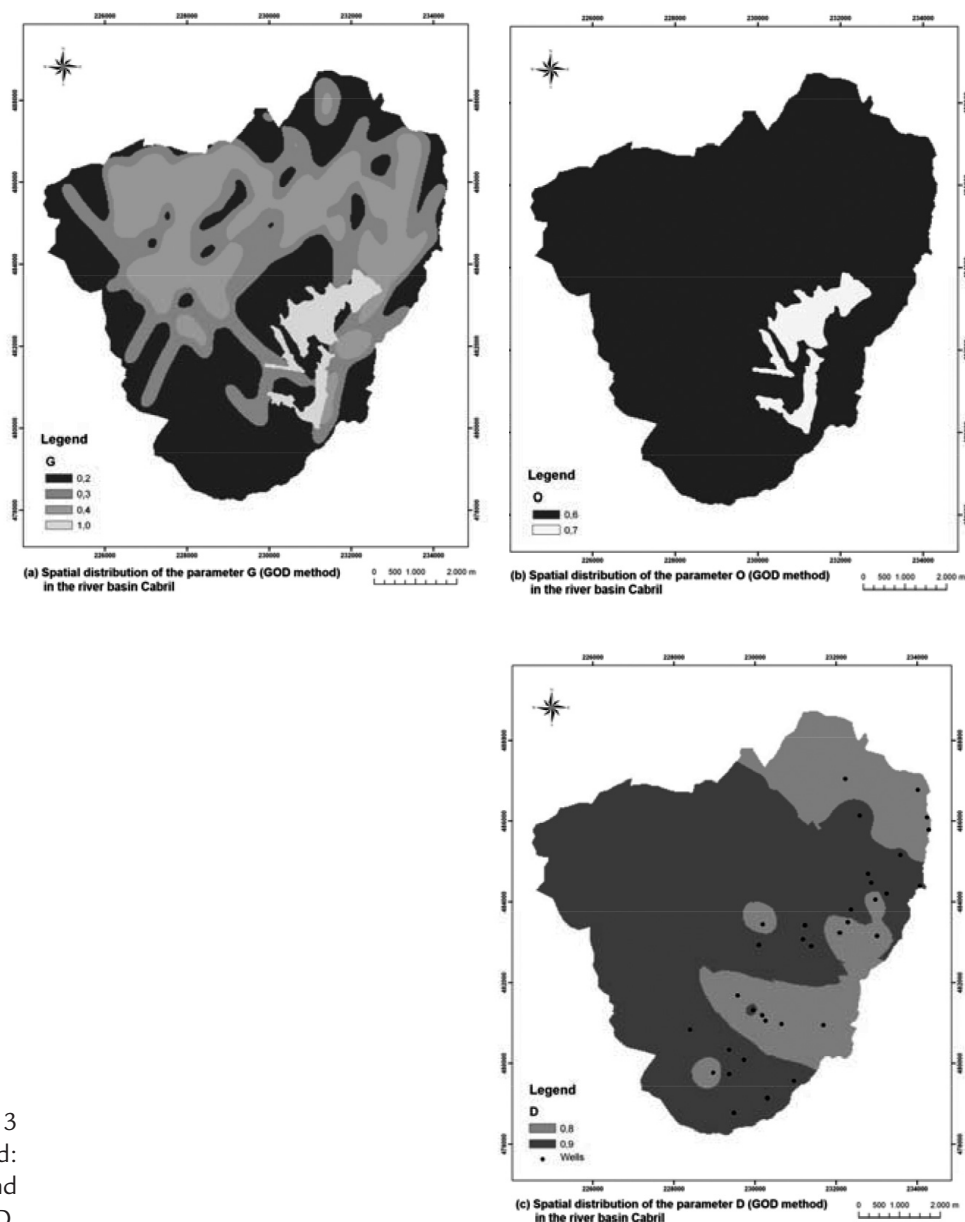


Figure 3
Spatial distributions of GOD method:
parameter G (a), parameter O (b) and
parameter D.

As to parameter (O) (Figure 3(b)), it may be observed that the area of the basin only comprises two lithological groups: the alluvium and the igneous formations, to which the values 0.7 and

0.6, respectively, correspond.

Finally, parameter D (Figure 3(c)), evaluated through the values of the depth of the ground-water table under the wells inventoried, is represented

by only two ranges: depths below 5 m and depths between 5 and 20 m. The figures for each range are 0.9 and 0.8, respectively.

3. Results and discussion

Regarding the results and mapping of the DRASTIC methodology, and with respect to parameter D (Figure 2(b)), it may be observed that the area of the ba-

sin is occupied mainly by partial indices 7 and 9. These indexes match with the depth ranges between 1.5 m and 4.6, and between 4.6 m and 9.1 m, respec-

tively. The highest values of the depth of the hydrostatic level, corresponding to the partial indices 3 and 5, are located on the spots of higher elevation.

The lowest values are found in alluvial regions, except for two wells located at the southern end of the basin. Analysing the map of parameter R (Figure 2(d)), it may be observed that most of the area studied presents, for this parameter, the average index of vulnerability. This includes recharge values in the range from 102-178 mm/year, corresponding to the partial index 6. On the other hand, only one of the sub-basins, bounded by the ArcSWAT method, is represented by the partial index 8 (178-254 mm/year). The area corresponding to partial index 1 (< 51 mm/year) is located along the main course of Cabril River. The map of parameter A (Figure 2(e)) reflects the choices of vulnerability made concerning the granitic areas, to which to the lowest values of the indices correspond, and concerning the areas of alluvium classified with higher indexes. Also arising

from the basis options, index A values are influenced by the higher or lower fracture density. The map of parameter S (Figure 2(f)) reveals higher vulnerability indices in low elevation areas of the basin under study, where soil units with considerable thickness and free sandy textures were developed. The zones with less vulnerability are those of higher elevation, where soils present reduced thickness of free sandy textures. For the parameter T map (Figure 2(g)), it was concluded that the areas of greatest vulnerability index are the Cabril Valley and tributary streams, dominated by small slopes. The remaining areas of the basin have indices of reduced vulnerability, much of the basin being classified with the lowest vulnerability index. Given that a typical index of 4 has been assigned for the entire basin, the map of parameter I shows

no spatial variation. A similar situation may be associated to parameter C, since as a typical index 2 was adopted for the alluvium and an index 1 for the remaining area of the basin, reflecting the low hydraulic conductivities estimated for the granites and metasediments.

The evaluation of the DRASTIC index is presented in Figure 4. It is observed that, in general, the Cabril River hydrographical basin is characterized by reduced indices (class 110-119). The areas of the basin that present higher vulnerability with indexes from the 120-139 and 140-159 classes, are those areas where sediments and alluvium are found, as well as the areas with higher fracture. The occupation of areas by the classes of lower vulnerability is also to be noted. This is the region composed of terraces and some high elevation areas from the NE quadrant of the basin.

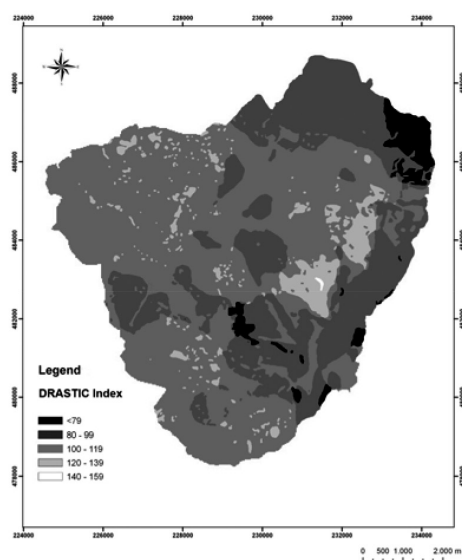


Figure 4
DRASTIC map for the Cabril River Basin.

The map of parameter G, obtained by applying the GOD methodology to the Cabril River Basin, is presented in Figure 3(a). The most vulnerable areas are the places where the soils are composed of alluvium and terraces, and the aquifer is assumed to be free. At these locations, the value of G is 1. According to the fracture density, the zones of granites and metasediments are assumed to be areas of semi-confined to confined aquifers, with values of G ranging between 0.2 and 0.4. The map of parameter O (Figure 3(b)) highlights the alluvium spots, with $O = 0.7$,

and the spots of igneous formations with $O = 0.6$. Finally, parameter D (Figure 3(c)) classifies the entire basin as high vulnerability, dividing it into sectors where $D = 0.9$ in which the depths of the groundwater table are less than 5 meters, and sectors where $D = 0.8$ corresponding to the areas where the depth of the groundwater table varies between 5 and 20 meters. It should be remembered that the wells inventoried are poorly distributed, as well as the absence of wells in large areas of higher altitude zones, situations that prevented an accurate assessment of parameter D.

The GOD vulnerability map is presented in Figure 5. It may be observed that only the Negligible, Low and High classes are present in the Cabril River Basin. For the most part, the basin is represented by Low vulnerability (0.1 - 0.3), being also covered by some areas where the vulnerability is negligible (0.0 - 0.1). These areas correspond to the locations where the depth of wells is higher (5 to 20m) and also to the sectors where the degree of confinement is low. In turn, areas that include free aquifers present high vulnerability (0.5 - 0.7).

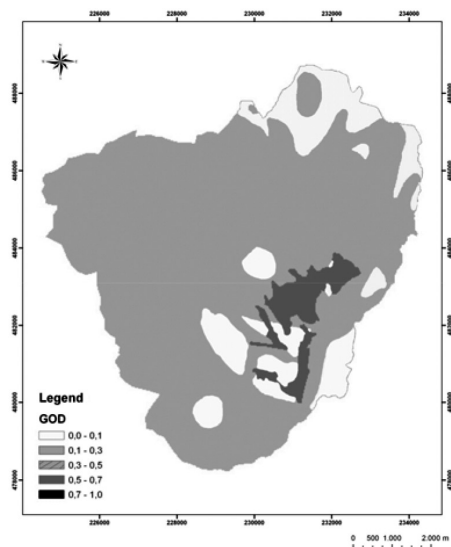


Figure 5
GOD map for the Cabril river basin.

The DRASTIC map was normalized to values from 0 to 1 (Figure 6), corresponding to the ranges of vulnerability of the GOD method, thereby

allowing the direct comparison of the two maps. It is verified, by applying the GOD method, that the area of the basin presents mostly low vulnerability, and

there is negligible vulnerability related to increased depth of groundwater level and also with low degree of confinement.

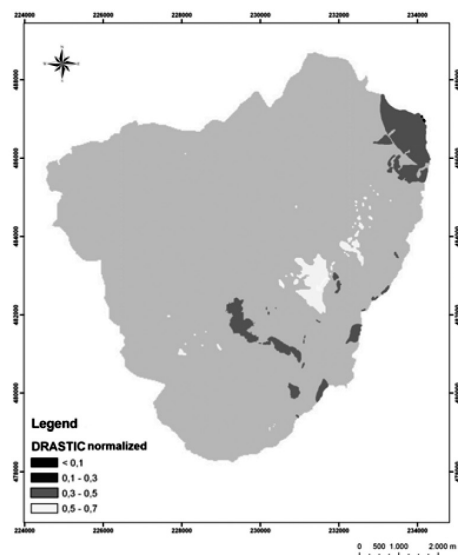


Figure 6
DRASTIC map normalized for the Cabril river basin.

The application of the DRASTIC method reveals that most of the basin presents medium vulnerability, and there

is low vulnerability located in an area of higher altitude. Both methods make high vulnerability correspond to areas covered

by alluvium, although by the DRASTIC method those vulnerabilities only cover a small portion of the alluvium.

4. Conclusion

The present study allowed determining which areas of the Cabril River Basin are more vulnerable to pollution, and therefore, depending on the vulnerability of the area under study, estimating the activities that may or may not be carried out at the surface, in order to protect groundwater, since we are dealing with an adjacent site to the capital of the Portuguese

northeast, with a tendency to grow in the area of the basin under study. It should be noted that it is one of the few Portuguese (inner) cities with growth rates. These studies are critical in areas where surface water is scarce; hence, the only reserve water for use of the resident population is the groundwater, which happens in this basin. The region has been very vulnerable in the

areas covered by alluvium, regardless of the method applied. This area ought to be well managed politically, with special care in terms of usage and protection of the aquifers in the region. The rest of the basin was considered by the DRASTIC method as of medium vulnerability and by the GOD method as of reduced vulnerability.

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