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Evaluation of geochemical data by two different XRF spectrometers in sediments from the Santiaguillo Basin (state of Durango, Mexico)

Jesús David Quiroz-Jiménez and Priyadarsi D. Roy*

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Resumen

La espectrometría de Fluorescencia de Rayos-X (FRX) es uno de los métodos más usados para medir datos geoquímicos y aplica técnicas tanto destructivas como no destructivas en la preparación de muestras. Se evaluó la calidad de 12 diferentes elementos usados comúnmente en la investigación de sedimentos calculando la precisión y exactitud de las concentraciones obtenidas en un espectrómetro portátil de FRX por dispersión en energía (Niton FXL 950), a partir de un método de preparación de muestras no destructivo. Los diferentes modos de medida del espectrómetro y la matriz de las muestras no influyeron significativamente en la precisión, aunque ambos influenciaron los valores de la exactitud. La concentración elemental en sedimentos con textura y matriz variables, colectados de la cuenca de Santiaguillo (estado de Durango, México), medida en el FRX por dispersión en energía mostraron correlaciones positivas altas ($r \geq 0.8$) con respecto a las concentraciones obtenidas en un espectrómetro convencional FRX para Si, Ti, Al, Fe, Mn, Ca, K, Sr y Zr. Las correlaciones bajas para Mg, P, y Rb ($r \leq 0.7$) fueron causadas probablemente por la matriz. Los resultados de elementos con correlaciones altas se corrigieron usando las ecuaciones de regresión lineal entre los datos de ambos espectrómetros.

Palabras clave: fluorescencia de rayos-X, comparación, precisión, exactitud, sedimento, matriz.

Abstract

XRF spectrometry is one of the commonly used methods to measure geochemical data and it involves both destructive and non-destructive sample preparation techniques. We evaluated data quality of 12 different regularly used elements in sedimentary research by calculating precision and accuracy of data generated in a portable ED-XRF (Niton FXL 950) and by applying a non-destructive sample preparation method. Different measuring modes of the spectrometer and matrix of the samples did not significantly influence the precision but influenced the accuracy. In sediments of variable matrix collected from the Santiaguillo Basin (state of Durango, Mexico), concentrations of Si, Ti, Al, Fe, Mn, Ca, K, Sr and Zr measured using the portable ED-XRF showed strong and positive correlations ($r \geq 0.8$) with concentrations obtained using a conventional WD-XRF. The matrix-effect possibly caused lower correlations ($r \leq 0.7$) for Mg, P and Rb. Concentrations of elements with stronger correlations were corrected using the linear regression equations generated after comparing results of both spectrometers.

Key words: X-ray fluorescence, comparison, precision, accuracy, sediment, matrix.

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Introduction

In multi-disciplinary research ambits, chemical composition of the geological deposits (e.g. sediment, soil and rock) has a variety of applications. Abundance and mobility of different inorganic elements provide useful information about the source of volcanic products (Roy *et al.*, 2012; Arce *et al.*, 2015), identification of provenance and tectonic settings of sedimentary basin (McLennan, 1989; Armstrong-Altrin *et al.*, 2012; Nagarajan *et al.*, 2014; Zhang *et al.*, 2016), estimation of diagenesis, chemical weathering, physico-chemical conditions of depositional environment (Nesbitt and Young, 1982; Cox *et al.*, 1995; Fedo *et al.*, 1995; El-Shafeiy *et al.*, 2016), and evaluation of the degree of contamination by anthropogenic activity (Kibria *et al.*, 2016; Retama *et al.*, 2016). Inductively coupled plasma mass spectrometry (ICP-MS), scanning electron microscopy (SEM), atomic absorption spectrometry (AAS) and X-ray fluorescence (XRF) spectrometry are some of the techniques that are used to obtain geochemical data. However, XRF spectrometry is relatively economic and quicker compared to all other earlier mentioned techniques and hence it is one of the commonly used methods in geosciences.

In XRF spectrometry, the primary X-rays interact with a sample and generate secondary X-rays (fluorescence) depending upon chemical composition of the geological material. Intensity of this fluorescence is used to determine concentration of different elements present in the sample (Lozano-Santacruz and Bernal, 2005). Presently, XRF spectrometers can measure wavelength or energy of the reflected X-rays and hence they can be grouped as wavelength dispersive XRF (WD-XRF) and energy dispersive XRF (ED-XRF). The conventional or traditional XRF spectrometers (generally WD-XRF) can measure concentrations of a large number of major and minor elements in fused discs and trace elements in pressed pellets. They do not allow recovery of the analyzed samples. However, the portable XRF spectrometers (generally ED-XRF) measure concentration of a limited number of elements in relatively less time. They require smaller amounts of sample and apply simpler preparation techniques. Hence, the samples can be completely recovered for other geochemical and mineralogical analyses.

Over the last couple of decades, several studies were focused on stratigraphic variation of elemental concentrations in sedimentary deposits in order to reconstruct the paleo-

environmental conditions of a location or region (Cheshire *et al.*, 2005; Metcalfe *et al.*, 2010; Roy *et al.*, 2013; Lozano-García *et al.*, 2015; Noble *et al.*, 2016) and also to evaluate the anthropogenic influences in different environments (Ruiz-Fernández *et al.*, 2007; Ontiveros-Cuadras *et al.*, 2014). The temporal variations in elemental contents provide information about varying precipitation/evaporation ratio, strength of aeolian activity, and rate of contaminant inflow. All such research is carried out with sediments recovered through cores and hence the sample size is relatively small. Thus, it is required to reuse the same sample for multiple chemical analyses preferably with application of non-destructive methodologies.

In the present study, we evaluate the data quality of twelve different commonly used elements (Si, Ti, Al, Fe, Mn, Mg, Ca, K, P, Rb, Sr and Zr) in sedimentary research by calculating precision and accuracy of the data generated using a portable ED-XRF (Niton FXL 950). Elemental concentrations were measured in 36 samples of a core collected from the Santiaguillo Basin (state of Durango, Mexico) by two different XRF spectrometers (portable and conventional) that apply two different methodologies (non-destructive and destructive). The main objective is to estimate the difference between data obtained in the portable ED-XRF and conventional WD-XRF using samples of variable matrix (i.e. grain size, texture, carbonate and organic matter content) from the Santiaguillo Basin as a case study.

Study site and sample

The Santiaguillo Basin (24° 44' N, 104° 48' W, 1960 m asl) is located at foothills of the Sierra Madre Occidental Mountains in the state of Durango. It is present at a distance of ~800 km northwest of Mexico City (Figure 1). Catchment of this basin mainly comprises igneous rocks. Both ignimbrite and rhyolite outcrops are present in the surrounding ~2500-3000 m high mountains (Nieto-Samaniego *et al.*, 2012). Outcrops of andesite and basalt are exposed to the east and west of the basin, respectively. Additionally, the eastern catchment also hosts smaller outcrops of metamorphic and sedimentary rocks (Munguía-Rojas *et al.*, 1998, not seen in Figure 1). Erosion of all the above-mentioned rocks provides sediments to the basin. A total of 36 different samples were selected from a 10.26 m long previously studied sediment core from this basin. Different grain size (i.e. silty-clay, silt, silty-sand, fine-sand, fine-medium-

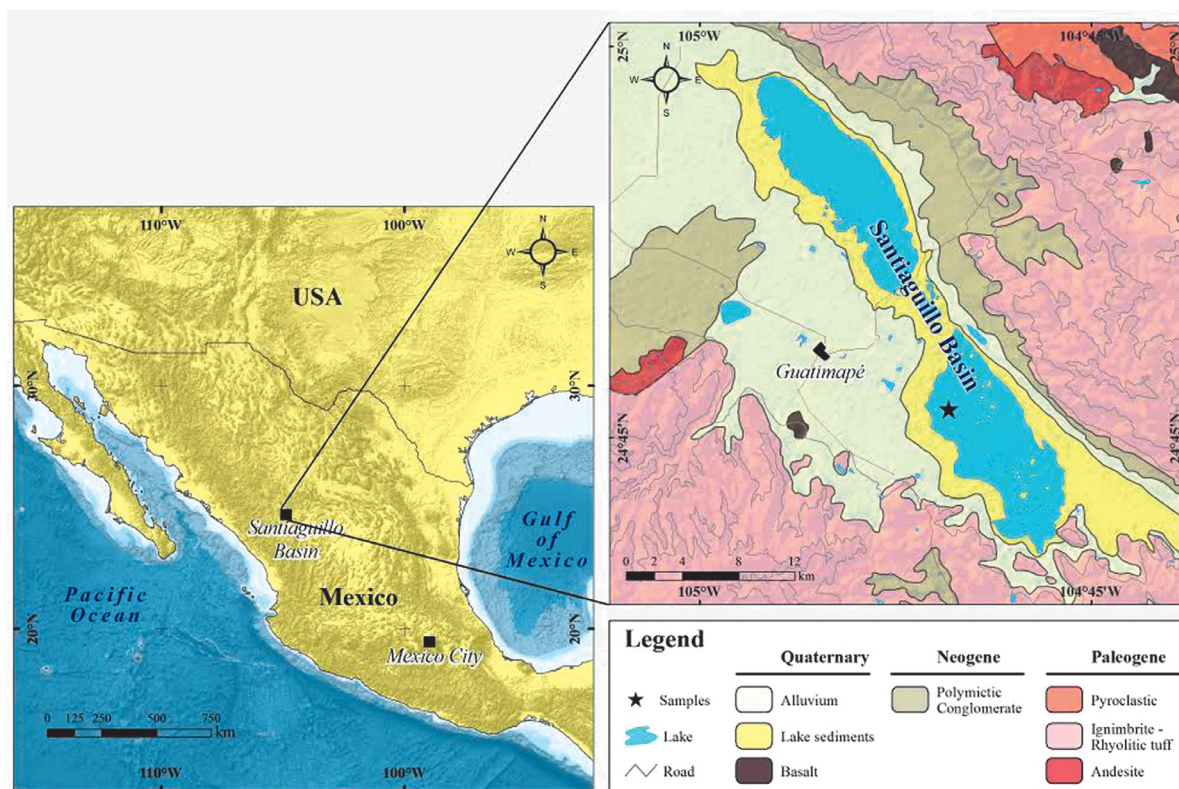


Figure 1. Map showing location of the Santiaguillo Basin and geology of the catchment.

sand and medium-coarse-sand), 0.3-2.0% of organic carbon (OC) and 2.3- 45.5% of calcite (CaCO_3) characterize the variable matrix of the samples (Table 1, Quiroz-Jiménez *et al.*, 2017)

Methodology

Samples were oven dried at 50°C, homogenized and grounded with an agate pestle till a texture similar to clay was achieved. Concentrations of 12 different elements (Si, Ti, Al, Fe, Mn, Mg, Ca, K, P, Rb, Sr and Zr) were measured in two different XRF spectrometers (i.e. portable Niton FXL 950 and conventional Siemens SRS 3000) following different sample preparation techniques. Three different geological reference materials (SRM-2702, SdAR-M2 and SRM-2709a) were measured only in the portable Niton FXL 950 XRF spectrometer.

Conventional Siemens SRS 3000 XRF

SRS 3000 is a WD-XRF with X-ray tube made of Rh. Analysis of major and minor elements like Si, Ti, Al, Fe, Mn, Mg, Ca, K and P was carried out using a destructive sample preparation method. It involved mixing 0.8 g dry and powder sediments with 7.2 g of (1:1) LiBO_2 and $\text{Li}_2\text{B}_4\text{O}_7$ mixture and this mixture was fused

in a Pt/Au (95:5) alloy crisol as per Lozano-Santacruz *et al.* (1995). Concentrations of trace elements like Rb, Sr and Zr were measured after preparing pressed pellets from the mixture of 4 g of powder sediment and binder following the instructions of Lozano-Santacruz *et al.* (1995) and Verma *et al.* (1996). The loss on ignition was estimated by a gravimetric method after heating 1 g of sample up to 950 °C in porcelain crucible for 1 hour. The measured major elements have precision of <5% and trace elements have precision of <10% (Lozano and Bernal, 2005).

Portable Niton FXL 950 XRF

Niton FXL 950 is an ED-XRF with a 50 kV X-ray tube made of Ag and a Geometrically Optimized Large area Drift Detector (GOLDD). Analysis of elemental concentration was carried out using a non-destructive sample preparation method. Approximately ~5 g dry and powdered sediment was placed in a plastic capsule and it was covered using 4 μm thick polypropylene X-ray film. Samples were analyzed in three different modes (designed for soil sample and metal enriched mining samples) of the equipment and each mode has different filters (i.e. soil mode: main, high and

Table 1. Texture of sediments collected from the Santiaguillo Basin and abundances of organic carbon (OC) and calcite (CaCO₃).

sample	depth (cm)	texture	OC (%)	CaCO ₃ (%)
1	40	silt	0.6	12.5
2	56	silty-sand	0.9	9.3
3	66	silt	0.6	13.7
4	106	silty-sand	0.9	7.8
5	126	silt	2.0	18.1
6	136	silty-sand	0.8	12.1
7	204	silty-sand	0.9	9.2
8	244	silty-sand	0.6	9.2
9	274	medium-coarse-sand	0.9	3.2
10	304	silty-sand	0.7	6.8
11	314	silt	0.4	29.4
12	404	silt	0.6	12.5
13	414	silty-clay	1.0	7.7
14	444	fine-medium-sand	0.5	4.1
15	474	fine sand	0.8	11.3
16	494	silty-sand	0.9	6.0
17	504	fine sand	0.8	5.0
18	524	silty-sand	1.0	12.1
19	554	silt	1.0	4.2
20	582	silty-sand	0.9	13.2
21	599	fine sand	0.7	3.3
22	639	fine sand	0.9	2.3
23	659	silty-sand	1.0	17.6
24	679	silt	1.3	4.2
25	696	silty-clay	0.7	5.9
26	726	fine sand	0.3	45.5
27	836	silty-sand	1.2	23.1
28	856	silt	0.6	9.1
29	886	silty-sand	0.6	18.3
30	896	silty-sand	0.5	6.2
31	916	silty-sand	0.7	6.0
32	936	silty-sand	1.3	12.8
33	976	silty-sand	1.3	6.9
34	986	silty-sand	0.9	13.7
35	1006	silty-sand	0.9	5.4
36	1016	silty-sand	0.7	20.1

low range filters; mining Cu/Zn mode: main, high, low and light range filters; mining Ta/Hf: main, high, low and light range filters). We measured each sample for 60 seconds in each filter of the three different modes. Principal difference between the modes is sensibility to different range of elements. For example, the soil mode does not detect lighter elements like Si, Al, Mg and P. However, the mining modes (Cu/Zn and Ta/Hf) can measure all the 12 elements including the lighter elements.

Both the precision and accuracy of the measured elements were determined after measuring elemental concentrations in three different reference materials (SRM-2702, SdAR-M2 and SRM-2709a) in seven repetitions, in all the modes of Niton FXL 950. National Institute of Standards and Technology (NIST) created the reference material SRM-2702 by using inorganic marine sediments and it does not present concentrations of Si, Fe and Zr (NIST, 2016). The United States Geological

Survey (USGS) created SdAR-M2 using metal rich soil (IAG, 2015) and NIST designed SRM-2709a as a reference material for soil, sediment and other geological samples of similar matrix (NIST, 2009).

Results and Discussion

Precision (%RSD)

Precision was obtained from Relative Standard Deviation (%RSD) of each element measured in three different reference materials (SRM-2702, SdAR-M2 and SRM-2709a) in three different modes of Niton FXL 950. It was calculated using the formula $\%RSD = 100 \times (s/X)$, where s is the standard deviation and X is the mean concentration of each element in seven repetitions. Soil mode did not measure concentrations of Si, Al, Mg, and P, whereas both the mining modes (mining Cu/Zn and mining Ta/Hf) analyzed concentration of all the 12 different elements (Table 2). Independent of matrix of the reference materials, Mg (%RSD=12-35.2) and P (%RSD=2.4-4.3) were the less precise elements in both the mining modes. Ca (%RSD=0.3-1.4) showed variable values of precision in mining Cu/Zn mode and both Ca (%RSD=0.2-1.9) and Al (%RSD=1.0-2.3) had variable precision in mining Ta/Hf mode. All the other elements had precision of ≤ 1.1 in the three different modes of analyses.

Soil mode: Fe (%RSD \leq 0.3) and K (%RSD \leq 0.4) were the most precise elements, whereas Mn (%RSD=1.4-1.8) was comparatively less precise. All other elements (Ti, Ca, K, Rb, Sr, Zr) had %RSD \leq 1.0.

Mining Cu/Zn mode: Si and Fe (%RSD \leq 0.4) were the most precise elements and Mg was the least precise element (%RSD $>$ 12). Compared to Ti, Al, Mn, Ca, K, Rb, Sr and Zr (%RSD \leq 1.4), P (%RSD $>$ 2.9) exhibited less precision.

Mining Ta/Hf mode: In general, precisions calculated in this mode were comparable to precision calculated for mining Cu/Zn mode. Fe (%RSD=0.2) was the most precise element and Mg (%RSD $>$ 15.3) was the less precise. P showed a precision of 2.4 and all other elements had precision of $<$ 1.9. Al showed relatively lower precision in this mode (%RSD=1.0-2.3) compared to the mining Cu/Zn mode (%RSD=1.0-1.2).

Accuracy (%ACC)

Accuracy of elements in three different modes of Niton FXL 950 was obtained by calculating $\%ACC = (X - X_{std} / X_{std}) \times 100$, where X is the mean concentration of an element measured in seven repetitions of a geological reference material and X_{std} is the certified concentration of that element. Table 3 presents certified

Table 2. Precision of elements calculated for three different modes of Niton FXL 950. It was estimated from Relative Standard Deviation (%RSD) of elemental concentrations measured in three different reference materials.

element	% RSD = $100 \times (\sigma/X)$								
	Soil mode (n=7)			Mining Cu/Zn mode (n=7)			Mining Ta/Hf mode (n=7)		
	SRM-2702	SdAR-M2	SRM-2709a	SRM-2702	SdAR-M2	SRM-2709a	SRM-2702	SdAR-M2	SRM-2709a
Si				0.4	0.3	0.2	0.7	0.5	0.2
Ti	0.3	0.5	0.4	0.6	0.9	0.5	0.5	0.8	0.6
Al				1.0	1.2	1.1	2.2	2.3	1.0
Fe	0.3	0.3	0.2	0.2	0.4	0.3	0.2	0.2	0.2
Mn	1.8	1.7	1.4	0.7	0.6	1.0	1.1	0.9	1.2
Mg				21.8	32.1	12.0	15.3	35.2	16.6
Ca	0.7	0.6	0.3	1.4	1.4	0.3	1.9	1.1	0.2
K	0.4	0.4	0.3	0.6	0.3	0.4	0.8	0.4	0.5
P				2.9	3.0	3.9	2.7	4.3	2.4
Rb	0.6	0.9	0.7	0.6	0.4	0.8	1.0	0.9	1.0
Sr	1.0	0.8	0.6	1.0	0.5	0.5	1.0	0.7	0.6
Zr	0.5	0.4	0.8	0.4	0.8	0.8	0.8	0.7	1.0

values of elemental concentration in 3 different reference materials and accuracy of 12 different elements calculated for three different modes. The positive values indicate overestimation of an element with respect to the certified value and negative values represent underestimation of an element compared to the certified value.

Depending upon the geological reference materials, each element showed different values of accuracy even though the analysis was carried out in the same mode. Mg (%ACC up to 252.5) and P (%ACC up to 264.9) showed the highest error in both the mining modes and soil mode did analyze both the elements. Similarly, both the mining modes reported the exact concentration of Mn as reported in the literature (%ACC=0) for the reference material SdAR-M2.

Soil mode: Ti (%ACC=1.8-24.4), K (%ACC=1.7-15.7), Ca (%ACC=6.7-24.2), Rb (%ACC=4.3-15.5) and Sr (%ACC=1.5-15.4) showed variable accuracy. Depending upon the geological reference material, both Ti (%ACC=24.4; SdAR-M2) and Mn (%ACC=24.5; SRM-2702) were the least accurate elements. Ti and K were the more accurate elements (%ACC<10) in reference material SRM-2702, Rb, Sr and Zr were accurately measured in SdAR-M2 and Ti, Ca, Sr and Zr were the more accurate elements measured in SRM-2709a.

Mining Cu/Zn mode: Both Mg (%ACC=87.9-252.5) and P (%ACC=51.9-138.8) were the least accurate elements. Ti (%ACC=0.7-30.2), Mn (%ACC=0-25.2), Ca (%ACC=5.1-17.9) and Sr (%ACC=2.3-17.5) showed variable accuracy. Si, Fe, Mn, Ca, K, Sr and Zr were accurately measured in SdAR-M2 and Ti, Fe, Mn, Ca, K and Sr were the more accurate elements measured in SRM-2709a. All these elements exhibited %ACC<10 and none of the elements in reference material SRM-2702 showed %ACC<10.

Mining Ta/Hf mode: Both Mg (%ACC=112.4-264.9) and P (%ACC=50.8-138.1) were the least accurate elements. Ti (%ACC=0.9-28.8), Al (%ACC=8.7-28.1), Mn (%ACC=0-24.7) and Sr (%ACC=0.1-16.6) showed variable accuracy. Al was the most accurate element (%ACC < 10) measured in reference material SRM-2702, Si, Fe, Mn, Ca, K, and Sr were measured with less error in SdAR-M2 and Ti, Fe, Mn, Ca, K and Sr were the more accurate elements measured in SRM-2709a.

Comparison of portable FRX vs. conventional XRF

Elemental concentrations of sediments collected from the Santiaguillo Basin were measured as a case study and the results obtained both in the portable as well as conventional XRF were compared in order to

Table 3. Accuracy of elements calculated for three different modes of Niton FXL 950. Mean of the measured values and certified value of three different geological reference materials are used for the accuracy (%ACC) calculation.

element	certified value			% ACC = $100 \times (X - X_{std})/X_{std}$								
	concentration (%)			soil mode (n=7)			mining Cu/Zn mode (n=7)			mining Ta/Hf mode (n=7)		
	SRM-2702	SdAR-M2	SRM-2709a	SRM-2702	SdAR-M2	SRM-2709a	SRM-2702	SdAR-M2	SRM-2709a	SRM-2702	SdAR-M2	SRM-2709a
Si		35.3	30.3					-5.7	-14.3		-6.6	-14.7
Ti	0.9	0.2	0.3	1.8	-24.4	-5.7	30.2	-13.3	0.7	28.9	-13.1	0.9
Al	8.4	6.0	7.4				10.9	31.2	16.2	8.7	28.1	14.8
Fe		1.8	3.4		-21.6	-11.6		-8.8	-2.8		-8.6	-3.0
Mn	0.2	0.1	0.1	24.5	-18.7	-11.1	25.2	0.0	-3.6	24.7	0.0	-2.3
Mg	1.0	0.3	1.5				87.9	252.5	149.6	112.4	264.9	159.0
Ca	0.3	0.6	1.9	-24.2	-21.4	-6.7	17.9	-5.1	7.0	19.3	-4.1	6.7
K	2.1	4.3	2.1	-1.7	-15.7	-14.4	34.5	-1.1	-4.4	33.3	-0.8	-4.5
P	0.2	0.0	0.1				51.9	138.8	78.9	50.8	138.1	79.4
Rb	0.013	0.015	0.001	-13.1	-4.3	-15.5	-19.9	-11.9	-21.5	-18.5	-11.1	-21.5
Sr	0.012	0.014	0.024	-15.4	1.5	-6.6	-17.5	-2.3	-7.9	-16.6	-0.1	-7.8
Zr		0.027	0.020		-3.5	-7.1		-9.4	-16.4		-14.2	-17.7

evaluate the matrix effect (Figures 2 and 3). Samples were powdered and homogenized in order to remove the texture influence. However, different abundances of organic carbon (OC: 0.3-2.0%) and calcite (CaCO_3 : 2.3-45.5%) suggest matrix of the sediment samples remained variable even after grinding and homogenizing the bulk samples.

Evaluation of the data obtained with Niton FXL 950 XRF indicates precision of <5 for all the elements except for Mg (%RSD >12) (Table 2). We did not observe significant influences of either the measuring modes of spectrometer or matrix of the sample on precision. Independent of the mode or matrix, the precision remained homogeneous. For example, Ti had precision of 0.3-0.6 and Ca of 0.7-1.9 in reference material SRM-2702 in all the three different modes. Similarly, Ti (0.3-0.5) and Ca (0.3-0.7) exhibited homogenous precision for all the geological reference materials with different matrix in soil mode. However, the measuring modes as well as matrix influenced the accuracy. Ti had an accuracy of 1.8-30.2 in the reference material SRM-2702 in all the three different modes. Concentration of Ca was underestimated in one mode and overestimated in another mode for the same reference material (Table 3).

Accuracies of Ti (0.7-30.2) and P (51.9-138.8) remained heterogeneous for the different geological reference materials with distinct matrices in the mining Cu/Zn mode.

Except for Mg, P and Rb, all other elements exhibited positive and stronger correlation coefficients between concentrations measured using both the XRF spectrometers. Si, Ti, Al, Fe, Mn, Ca, K and Sr ($r \geq 0.9$) had relatively higher correlations compared to Zr ($r=0.8$) (Figures 2 and 3). Correlation coefficients of Mg, P, and Rb remained ≤ 0.7 . In sediments from the Santiaguillo Basin, concentrations of Mg (0.29-1.87 %), P (0.01-0.04 %) and Rb (18-188 ppm) were relatively lower compared to many other elements (e.g. Si and Ca). Lower abundances of Mg, P and Rb might have caused the lower correlations. However, some elements with lower concentrations showed higher correlations. For example, concentrations of Ti and Ca were 0.1% and 0.4% in some samples and both of them exhibited higher correlations. Similarly, Mn (0.01-0.02%) exhibited better correlation compared to P (0.01-0.04%) even though both had comparable concentrations. Thus, we rule out the influence of abundance on different correlation coefficients exhibited by the elements. The non-linearity of X-ray

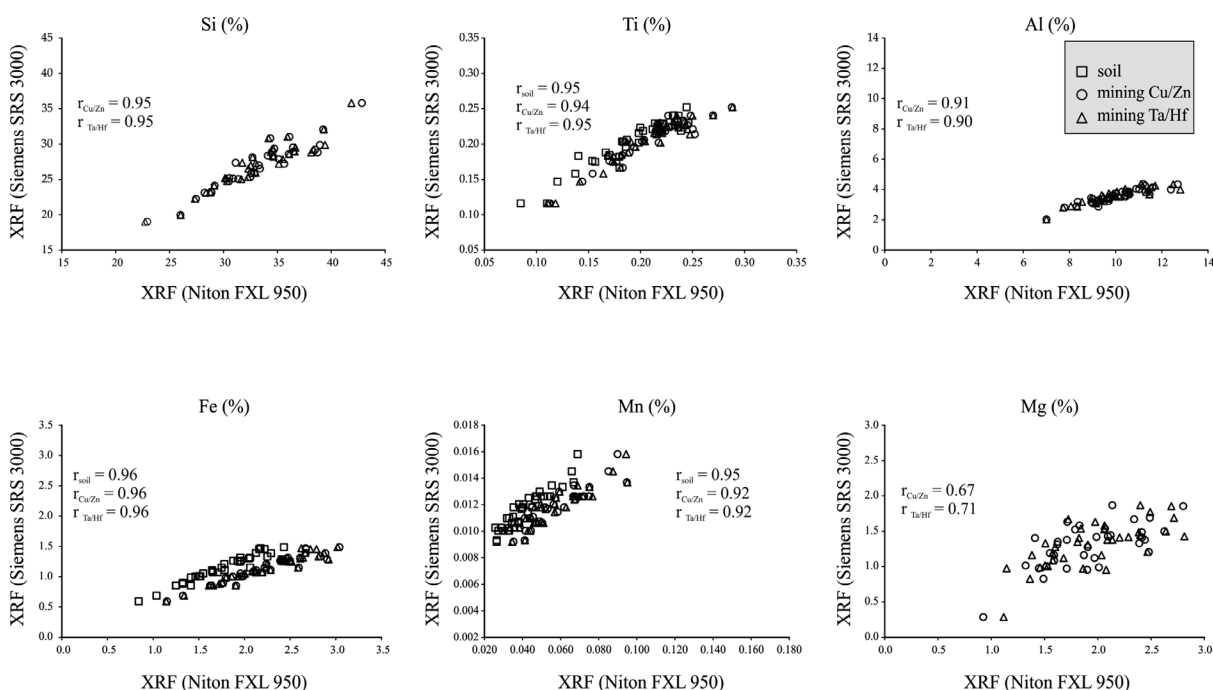


Figure 2. Comparison of concentrations of Si, Ti, Al, Fe, Mn and Mg measured in three different modes of portable (Niton FXL 950) and conventional (Siemens SRS 3000) XRF spectrometers.

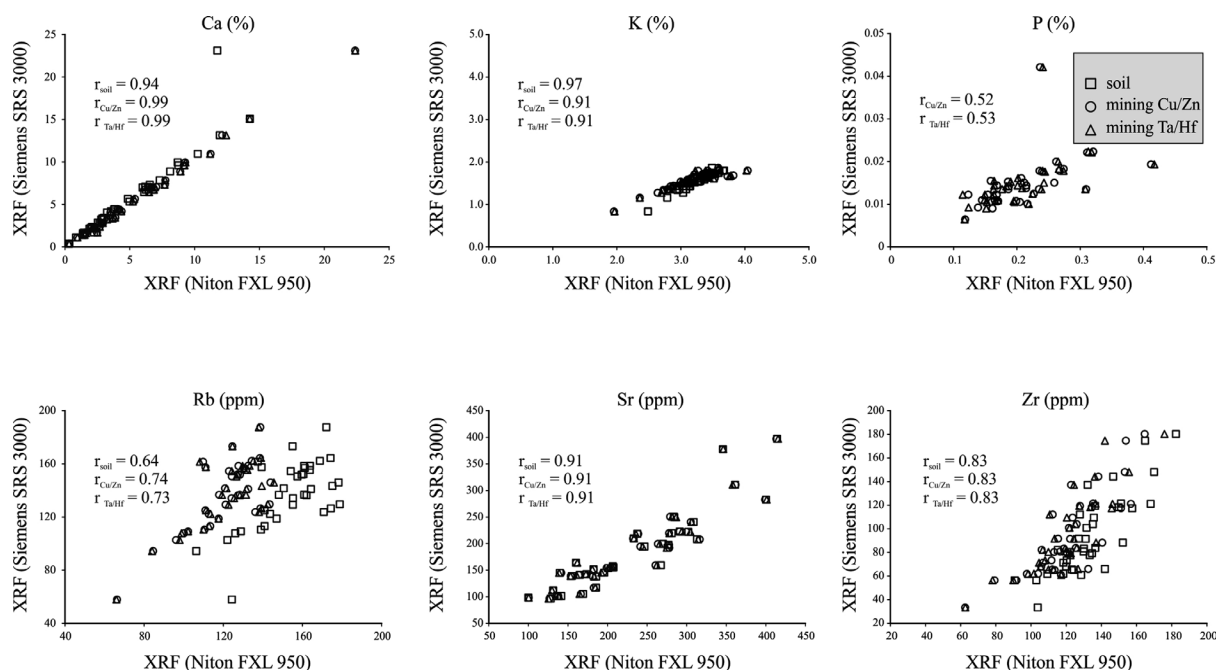


Figure 3. Comparison of concentrations of Ca, K, P, Rb, Sr and Zr measured in three different modes of portable (Niton FXL 950) and conventional (Siemens SRS 3000) XRF spectrometers.

intensity with concentration of an element is associated with both absorption and enhancement related to the matrix effect (Birks and Harris, 1962; Mitchell and Kellam, 1968). Based on the characteristics of matrix (lighter, heavy and neutral), X-ray intensity of an element can increase, decrease or remain neutral (Mitchell and Kellam, 1968). Birks and Harris (1962) studied Fe and Cr in a chromite-olivine mineral system and observed that X-ray intensity of Fe decreased as absorption coefficient of Cr increased in a matrix with more Cr compared to Fe.

We estimated the difference between results obtained using both the XRF spectrometers by calculating $\%change = (X_{FXL} - X_{SRS3000} / X_{SRS3000}) \times 100$, where X_{FXL} is elemental concentration in the portable XRF and $X_{SRS3000}$ is concentration of the same element measured in conventional XRF (Table 4). The portable Niton FXL950 XRF overestimated concentrations of some elements and underestimated concentrations of others compared to the conventional XRF. Positive values represent overestimation of an element in portable XRF (in %) and negative values suggest underestimation of an element (in %) compared to the results of conventional XRF. Irrespective of the sample matrix, estimation of the %change was similar

for each element in both the mining modes. Estimated difference for the soil mode was distinct compared to both the mining modes. For example, Ti was underestimated in the soil mode and overestimated in both the mining modes. Similarly, Rb was overestimated in soil mode and underestimated in both the mining modes. Fe and Zr were overestimated in all the three modes of analysis. However, the soil mode overestimated Fe to a lesser degree and Zr to a higher degree compared to results obtained in both the mining modes.

Soil mode: Average values suggest underestimation of Ti (%change=5.2) and Ca (%change=11.2) and overestimation of Mn (%change=246.9) and K (%change=112.1). Concentrations of Fe, Rb, Sr and Zr were overestimated with %Change between 15 and 54.2.

MiningCu/Zn: ExceptforRb,allotherelements were overestimated. Both Ti (%change=1.5) and Ca (%change=0.4) showed the least differences and P (%change=1401.6) exhibited the highest difference. Concentrations of Al, Mn and K (%change=103.6-362.5) were more overestimated compared to Si, Mn, Sr and Zr (%change=24.1-54.6).

Mining Ta/Hf: Differences calculated for this mode is similar to the %change estimated for all the elements in the mining Cu/Zn mode (Table 4). Concentrations of Ti and Ca measured in both the mining modes of the portable XRF were almost similar to concentrations obtained in the conventional XRF. Similarly, the concentration of P obtained in portable XRF was overestimated by almost 14 times with respect to the conventional XRF.

We generated linear regression equation for each element after comparing results obtained in both spectrometers and corrected only the concentrations of elements (Si, Ti, Al, Fe, Mn, Ca, K, Sr and Zr) with stronger correlations ($r > 0.8$). Concentrations of elements (Mg, P and Rb) with lower correlations ($r \leq 0.7$) were not rectified.

Conclusions

Concentrations of twelve different commonly used elements (Si, Ti, Al, Fe, Mn, Mg, Ca, K, P, Rb, Sr and Zr) in sedimentary research were measured in sediments from the Santiaguillo Basin (state of Durango, Mexico) with variable matrix using a portable ED-XRF (Niton FXL 950) and non-destructive sample preparation technique. This study suggests that concentrations of Si, Ti, Al, Fe, Mn, Ca, K, Sr and Zr can be quickly and economically analyzed in a portable Niton FXL 950 XRF

spectrometer and subsequently corrected. Additionally, the use of non-destructive sample preparation technique can be useful to reuse the same sample for multiple other chemical analyses. The main conclusions of our research are:

Except for Mg and P, all other elements had precisions of < 2.3 . Precision remained homogeneous independently of the measuring mode or matrix. Among the less precise elements, P (%RSD=2.4-4.3) showed better precision compared to Mg (%RSD=12-35.2).

Accuracy was influenced by both the measuring mode as well as matrix. For example, accuracies for Ti (%ACC=0.7-30.2) and P (%ACC=51.9-138.8) were heterogeneous for different geological reference materials with distinct matrices in the mining Cu/Zn mode.

In sediments of the Santiaguillo Basin with variable matrix, concentrations of Mg, P, and Rb measured in the portable XRF exhibited relatively lower correlations for ($r \leq 0.7$) with the results obtained in a conventional XRF. Si, Ti, Al, Fe, Mn, Ca, K and Sr showed higher ($r \geq 0.9$) correlation coefficients compared to Zr ($r = 0.8$). We rule out the influence of abundance on lower correlation coefficients of some elements. However, the variable matrix of sediments might have caused the lower correlations for Mg, P and Rb.

Table 4. Estimation of degree of difference (%change) of elemental concentrations of 36 samples from the Santiaguillo Basin measured in portable Niton FXL 950 XRF compared to conventional XRF (Min=minimum, Max=maximum and avg= average).

% change = (XFXL950 - xSRS3000/xSRS3000) × 100									
element	soil mode			Mining Cu/Zn			Mining Ta/Hf		
	Min	Max	Avg	Min	Max	Avg	Min	Max	Avg
Si				11.4	45.8	24.1	10.9	45.5	23.7
Ti	-27.0	9.2	-5.2	-8.8	18.0	1.5	-8.0	15.8	1.6
Al				156.2	244.4	183.4	154.0	244.2	180.8
Fe	41.8	87.4	54.2	81.7	125.9	97.0	78.5	127.1	96.7
Mn	63.1	388.2	246.9	125.3	592.4	362.5	109.0	594.6	362.2
Mg				0.5	223.1	54.6	3.2	290.1	55.9
Ca	-49.1	31.9	-11.2	-21.0	45.7	-0.4	-25.4	47.5	-0.7
K	87.4	197.0	112.1	82.5	133.5	103.6	79.9	134.6	103.4
P				461.3	2184.0	1401.6	471.3	2175.2	1370.0
Rb	-14.9	114.8	15.0	-32.0	14.5	-10.1	-33.1	14.0	-9.6
Sr	-8.5	67.9	26.8	-8.5	63.7	25.7	-8.4	64.7	25.9
Zr	-5.6	210.9	54.1	-11.9	100.8	37.2	-18.7	92.4	36.5

Elements were either underestimated or overestimated in the portable XRF and we used the linear regression equations to reduce difference between results obtained in both the spectrometers. Elements (Si, Ti, Al, Fe, Mn, Ca, K, Sr and Zr) with stronger and positive correlations were rectified. However, concentrations of Mg, P and Rb with lower correlations suffered from the matrix effect and hence were not corrected.

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References

- Arce J.L., Muñoz-Salinas E., Castillo M., Salinas I., 2015, The ~2000yr BP Jumento volcano, one of the youngest edifices of the Chichinautzin Volcanic Field, Central Mexico. *J. Volc. Geoth. Res.*, 308, 30-38.
- Armstrong-Altrin J.S., Lee Y.I., Kasper-Zubillaga J.J., Carranza-Edwards A., Garcia D., Eby G.N., Balaram V., Cruz-Ortiz N. L., 2012, Geochemistry of beach sands along the western Gulf of Mexico, Mexico: implication for provenance. *Chemie der Erde-Geochemistry*, 72, 4, 345-362.
- Birks L.S., Harris D.L., 1962, Unusual matrix effects in fluorescent X-ray spectrometry. *Analyt. Chem.*, 34, 943-945.
- Cheshire H., Thurow J., Nederbragt A.J., 2005, Late Quaternary climate change record from two long sediment cores from Guaymas Basin, Gulf of California. *J. Quat. Sci.*, 20, 5, 457-469.
- Cox R., Lowe D.R., Cullers R.L., 1995, The influence of sediment recycling and basement composition on evolution of mudrock chemistry in the southwestern United States. *Geochimica et Cosmochimica Acta*, 59, 2919-2940.
- El-Shafeiy M., El-Kammar A., El-Barkooky A., Meyers P.A., 2016, Paleo-redox depositional conditions inferred from trace metal accumulation in two Cretaceous-Paleocene organic-rich sequences from Central Egypt. *Marine and Petroleum Geology*, 73, 333-349.
- Fedo C.M., Nesbitt H.W., Young G.M., 1995, Unraveling the effects of potassium metasomatism in sedimentary rocks and paleosols, with implications for paleoweathering conditions and provenance. *Geology* 23, 921-924.
- IAG. International Association of Geoanalysts, 2015. Reference Material Data Sheet, SDAR-M2, Metal Rich Sediment.
- Kibria G., Hossain M.M., Mallick D., Lau T.C., Wu R., 2016, Trace/heavy metal pollution monitoring in estuary and coastal area of Bay of Bengal, Bangladesh and implicated impacts. *Marine pollution bulletin*, 105, 393-402.
- Lozano Santa Cruz R., Verma S.P., Girón P., Velasco F., Morán Zenteno D., Viera F., Chávez G., 1995, Calibración preliminar de Fluorescencia de Rayos-X para análisis cuantitativo de elementos mayores en rocas ígneas. *Actas INAGEQ*, 1, 203-208.
- Lozano Santa Cruz R., Bernal J.P., 2005, Assessment of a new set of geochemical reference materials for XRF major and trace element analysis. *Revista Mexicana de Ciencias Geológicas*, 22, 3, 329-344.
- Lozano-García S., Ortega B., Roy P.D., Beramendi-Orosco L., Caballero, M., 2015, Climatic variability in the northern sector of the American tropics since the latest MIS 3. *Quat. Res.*, 84, 2, 262-271.
- McLennan S.M., 1989, Rare earth elements in sedimentary rocks: influence of provenance and sedimentary processes, in: Lipin, B.R., McKay, G.A. (Eds.), *Geochemistry and Mineralogy of Rare Earth Elements*. *Rev. Mineral.*, 21, 169-200.
- Metcalf S., Jones M.D., Davies S.J., Noren A., Mackenzie A., 2010, Climate variability over the last two millennia in the North American Monsoon region, recorded in laminated lake sediments from Laguna Juanacatlán México. *The Holocene* 20, 1195-1206.
- Mitchell B.J., Kellam J.E., 1968, Unusual matrix effects in X-ray spectroscopy: A study of the range and reversal of absorption enhancement. *Appl. Spect.*, 22, 742-748.

- Munguía-Rojas P., García-Padilla J.L., Armenta-Román R., Cruz-Pérez R., Camacho J. M., Céspedes J.S., 1998, Carta Geológico-Minera, Durango G13-11, escala 1:250000 con texto explicativo. Secretaría de Comercio y Fomento Industrial, Servicio Geológico Mexicano.
- Nagarajan R., Roy P.D., Jonathan M.P., Lozano R., Kessler F.L., Prasanna M.V., 2014, Geochemistry of Neogene sedimentary rocks from Borneo Basin, East Malaysia: Paleo-weathering, provenance and tectonic setting. *Chemie der Erde-Geochemistry*, 74, 139-146.
- National Institute of Standards and Technology NIST, 2009. Certificate of Analysis, Standard Reference Material 2709a, San Joaquin Soil.
- NIST, National Institute of Standards and Technology, 2016. Certificate of Analysis, Standard Reference Material 2702, Inorganics in Marine Sediments.
- Nesbitt H.W., Young G.M., 1982, Early Proterozoic climates and plate motions inferred from major element chemistry of lutites. *Nature*, 299, 715-717.
- Nieto-Samaniego A.F., Barajas-Gea C.I., Gómez-González J.M., Rojas A., Alaniz-Álvarez S.A., Xu S., 2012, Geología, evolución estructural (Eoceno al actual) y eventos sísmicos del Graben de Santiaguillo, Durango, México. *Revista Mexicana de Ciencias Geológicas*, 29, 115-130.
- Noble P.J., Ball G.I., Zimmerman S.H., Maloney J., Smith S.B., Kent G., Adams K.D., Karlin R.E., Driscoll N., 2016, Holocene paleoclimate history of Fallen Leaf Lake, CA., from geochemistry and sedimentology of well-dated sediment cores. *Quat. Sci. Rev.*, 131, 193-210.
- Ontiveros-Cuadras J.F., Ruiz-Fernández A.C., Sanchez-Cabeza J.A., Pérez-Bernal L.H., Sericano J.L., Preda M., Wee Kwong L.L., Páez-Osuna F., 2014, Trace element fluxes and natural potential risks from 210 Pb-dated sediment cores in lacustrine environments at the Central Mexican Plateau. *Sci. Total Environ.*, 468, 677-687.
- Quiroz Jiménez J.D., Roy P.D., Beramendi-Orosco L., Lozano-García S., Vázquez-Selem L., 2017, Orbital-scale droughts in the central-northern Mexico during the late Quaternary and comparison with other subtropical and tropical records. *Geolog. J.*, DOI: 10.1002/gj.2888.
- Retama I., Jonathan M.P., Roy P.D., Rodríguez-Espinosa P.F., Nagarajan R., Sarkar S.K., Morales-García S.S., Muñoz-Sevilla N.P., 2016, Metal concentrations in sediments from tourist beaches of Huatulco, Oaxaca, Mexico: an evaluation of post-Easter week vacation. *Environ. Earth Sci.*, 75, 375.
- Roy P.D., Arce J.S., Lozano R., Jonathan M.P., Centeno-García E., Lozano-García S., 2012, Geochemistry of a tephra-paleosol sequence from north-eastern Basin of Mexico: implications to tephrochronology and chemical weathering during Quaternary. *Revista Mexicana de Ciencias Geológicas*, 29, 24-38.
- Roy P.D., Quiroz-Jiménez J.D., Pérez-Cruz L.L., Lozano-García S., Metcalfe S.E., Lozano-Santacruz R., López-Balbiaux N., Sánchez-Zavala J.L., Romero F.M., 2013, Late Quaternary paleohydrological conditions in the drylands of northern Mexico: a summer precipitation proxy record of the last 80 cal ka BP. *Quat. Sci. Rev.*, 78, 342-354.
- Ruiz-Fernández A.C., Hillaire-Marcel C., Páez-Osuna F., Ghaleb B., Caballero-Miranda M., 2007, 210 Pb chronology and trace metal geochemistry at Los Tuxtlas, Mexico, as evidenced by a sedimentary record from the Lago Verde crater lake. *Chin. J. Geochem.*, 67, 181-192.
- Verma S.P., Lozano-Santa Cruz, R., Girón-García, P., Velasco, F., 1996, Calibración preliminar de Fluorescencia de Rayos-X para análisis cuantitativo de elementos traza en rocas ígneas. *Actas INAGEQ*, 2, 237-242.
- Zhang X., Zhao G., Eizenhöfer P.R., Sun M., Han Y., Hou W., Zhu C.Y., 2016, Tectonic transition from late carboniferous subduction to early Permian post-collisional extension in the eastern Tianshan, NW China: Insights from geochronology and geochemistry of mafic-intermediate intrusions. *Lithos*, 256-257, 269-281.