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Mid-early late Albian foraminiferal assemblage from the El Abra Formation in the El Madroño locality, eastern Valles-San Luis Potosí Platform, Mexico: Paleoenvironmental and paleobiogeographical significance

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ABSTRACT

Albian-Cenomanian carbonate deposits of the El Abra Formation are widely distributed in the folded Sierra Madre Oriental. At the El Madroño locality, this stratigraphic unit contains an abundant, diverse and well-preserved fossil association composed of foraminifers and algae as well as rudists and other kinds of bivalves, gastropods, as well as corals, echinoids, and sponges. The El Madroño site is located in the NE extreme of the state of Querétaro in central-eastern Mexico. The foraminiferal assemblage is well-preserved and consists mostly of orbitolinids: *Dictyoconus walnutensis* (Carsey, 1926), *Paracoskinolina sunnilandensis* (Maync, 1955b), *Coskinolinoides texanus* (Keijzer, 1942), as well as *Buccicrenata subgoodlandensis* (Vanderpool, 1933), *Novalesia angulosa* (Magniez, 1972), *Mayncina* sp., *Istriloculina* sp., *Arenobulimina* sp., *Buccicrenata* sp. and *r* strategist planktic foraminifers such as *Favusella washitensis* (Carsey, 1926), *Muricohedbergella simplex* (Morrow, 1934), and *M. albiana* (BouDagher-Fadel, et al., 1996). Among the calcareous algae are *Cayeuxia kurdistanensis* Elliott, 1957, *Terquemella americana* (Konishi and Epis, 1962), *Acroporella* sp., *Triploporella* sp., and the encruster *Lithocodium aggregatum* Elliott, 1956. The age assignment of the deposit is based on the benthic foraminifer *Dictyoconus walnutensis*, which is considered a regional marker of middle Albian age, however this succession spans to the early Albian. The paleoenvironmental reconstruction, inferred from the lithology and benthic assemblage, suggests a warm shallow-water platform.

Keywords: Albian, foraminiferal assemblage, El Madroño, Valles-San Luis Potosí Platform, Mexico.

RESUMEN

Los depósitos carbonatados del Albiano-Cenomaniano de la Formación El Abra están ampliamente distribuidos en la Sierra Madre Oriental. En la localidad de El Madroño, esta unidad estratigráfica cuenta con una abundante, diversa y bien preservada asociación compuesta de foraminíferos y algas así como rudistas y otros bivalvos, gasterópodos, así como corales, echinoides y esponjas. El Madroño está localizado en el extremo NE del estado de Querétaro, en el centro-este de México. La asociación de foraminíferos está bien preservada; principalmente consiste de orbitolinidos: *Dictyoconus walnutensis* (Carsey, 1926), *Paracoskinolina sunnilandensis* (Maync, 1955 b), *Coskinolinoides texanus* (Keijzer, 1942), *Buccicrenata subgoodlandensis* (Vanderpool, 1933), *Novalesia angulosa* (Magniez, 1972), *Mayncina* sp., *Istriloculina* sp., *Arenobulimina* sp., *Buccicrenata* sp. y los foraminíferos planctónicos con estrategia *r* como *Favusella washitensis* (Carsey, 1926), *Muricohedbergella simplex* (Morrow, 1934), and *M. albiana* (BouDagher-Fadel, et al., 1996). Entre las algas calcáreas están *Cayeuxia kurdistanensis* Elliott, 1957, *Terquemella americana* (Konishi and Epis, 1962), *Acroporella* sp., *Triploporella* sp. y el incrustante *Lithocodium aggregatum* Elliott, 1956. La edad asignada al depósito está basada en el foraminífero bentónico *Dictyoconus walnutensis* el cual es considerado un marcador regional para la edad Albiano medio, sin embargo esta sucesión se extiende al Albiano temprano. La reconstrucción paleoambiental es inferida de la litología y la asociación bentónica que sugieren un ambiente de plataforma de aguas cálidas someras.

Palabras clave: Albiano, asociación de foraminíferos, El Madroño, Plataforma Valles-San Luis Potosí, México.

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1. Introduction

In the mid-Cretaceous, shallow-water platforms were widely distributed surrounding the Gulf of Mexico basin, particularly during the Albian, where carbonate deposition extended from the Bahamas along the northern rim of the Gulf, across Mexico to the Yucatan Peninsula, and to various Caribbean localities (Scott, 1990). The larger benthic foraminifers that inhabited such environments are an important tool for dating these rocks (Schroeder and Neumann, 1985). In this paper these microfossils are also used for age assignment. Garfias (1915) was the first to use the term El Abra Limestone for the shallow-water carbonates from the Sierra de El Abra. The El Abra Formation is a carbonate complex widely distributed in the Valles–San Luis Potosí Platform (VSLPP), cropping out in the central part of San Luis Potosí State and some areas of the states of Guanajuato, Querétaro, Hidalgo, Tamaulipas and Nuevo León (Carrillo-Bravo, 1971; Wilson and Ward, 1993). This unit comprises two different facies or members based on the lithology and faunal characteristics, which were recognized in the earliest studies. The Taninul facies consists of massive and very thick limestone ranging from light to medium gray color, containing a great variety of invertebrates, mostly rudists, other bivalves, gastropods, echinoids, foraminifers and algae. The El Abra facies, which is cream in color and contains numerous miliolids, is a lagoonal facies from the Albian-Cenomanian (Aguayo, 1998).

The stratigraphic framework for the Valles–San Luis Potosí platform has been reported by Nigra (1950), Carrillo Bravo (1971), Smith (1986), Basañez *et al.* (1993), Wilson and Ward (1993). The sedimentology of the El Abra process has been documented by Minero (1982), Enos *et al.* (1983), Enos and Stephens (1993) and Aguayo (1998).

Paleontological studies of the taxonomy of rudists from the El Madroño locality have been published by Alencaster (1987, 1998), Scott (1990), and Alencáster and García Barrera (2008); based on the study of these fossils, the authors have assigned

an age from the late Albian to early Cenomanian. Albian-Cenomanian benthic foraminifera from the El Abra Formation located in the eastern part of the VSLPP were reported by Bonet (1952, 1956), Tavitas and Solano (1984), Ornelas *et al.* (2006). Rosales-Domínguez (1998) made the first taxonomic study of the foraminifers from the El Madroño locality (state of Querétaro). Buitrón *et al.* (1995) analyzed Albian gastropods, foraminifera and algae from the Ahuacatlán region also in the state of Querétaro. The El Abra Formation has generally been assigned to the Albian and Cenomanian (Muir, 1936), based mostly on the invertebrate fossils. Later paleontological studies from the top of the El Abra Formation specify an early Cenomanian chronostratigraphic position (Young, 1977, 1984). Recent studies based on the benthic foraminifers indicate that in the west part of the VSLPP the age assigned to the El Abra Formation spans to the mid-late Cenomanian (Omaña *et al.*, 2013).

The objective of this paper is to document the occurrence of the benthic foraminifera and algae from the El Abra limestone samples collected from El Madroño in order to date this interval. The microfossils were analyzed with the aim of inferring the paleoenvironmental conditions; the paleobiogeographic distribution of the foraminifers was also examined.

2. Geological Setting

The study area is located near the town of El Madroño situated in the NE part of the state of Querétaro (Figure 1). This area lies on the southeastern edge of the VSLPP.

The VSLPP is part of an extensive carbonate platform system that rimmed the ancestral Gulf of Mexico during the mid-Cretaceous. The Early Cretaceous was a time of remarkable tectonic stability in the Gulf of Mexico basin, characterized by a decrease in terrigenous influx and development of stable shelves, ramps and platforms bordering the deep central part of the Gulf of Mexico

basin, which became the site of widespread carbonate deposition, particularly during the Albian (Salvador, 1991).

According to Wilson and Ward (1993) the VSLPP is one of “the largest isolated carbonate platforms

(200 by 300 km.) located on the western side of the Gulf of Mexico, it began to develop in Early Cretaceous but reached maximum growth during the Albian when it evolved a rimmed shelf margin. The platform apparently formed along the

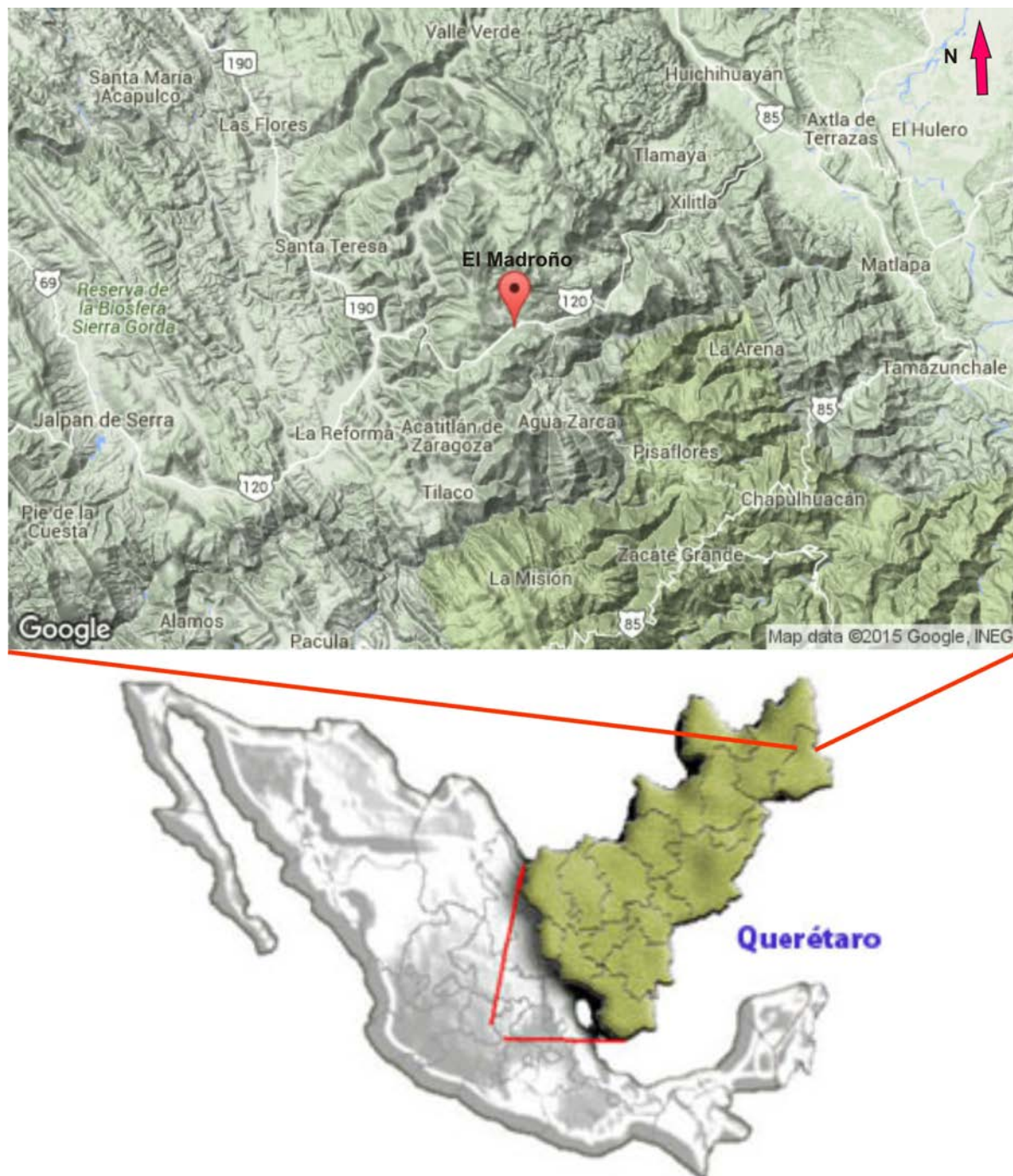


Figure 1 Image taken from Google Maps, with the location of the studied area (21° 06'–21°27' Lat N; 99° 03'–99° 22' Long W).

southwestern flank of a north-northwest-trending tectonic ridge of Precambrian gneiss and Upper Paleozoic schist which is exposed at Miquihua-na, 70 km. southwest of Ciudad Victoria on the state line of Nuevo León and Tamaulipas. Here thin Lower Cretaceous rocks rest unconformably on Jurassic red-beds, which lie unconformably on Upper Paleozoic talcose sericitic schists”

Carrillo-Bravo (1971) describes to the VSLPP as a “Barremian-Aptian rudist reef-rimmed, almost atoll-like configuration”. This carbonate platform is believed to have surrounded a lagoon filled with evaporites and terrigenous mud of the 2000 m-thick Guaxcamá Formation. The Lower Cretaceous rocks are overlain by 1500 – 2000 m of Albian-Cenomanian reef and backreef carbonate rocks termed the El Abra Formation (Garfias, 1915, Carrillo-Bravo, 1971).

The Albian–Cenomanian carbonate platform now lies within the Sierra Madre Oriental fold thrust belt (Suter 1984, 1987, 1990). Its eastern shelf margin is exposed in quarries along the Cuesta de El Abra near Valles and Antiguo Morelos. Part of its western side is seen along road cuts 30 km east of the city of San Luis Potosí.

3. Material and Methods

The samples come from the Taninul facies collected of the El Madroño site, which is located approximately 50 km from Jalpan (Querétaro) in the direction of Xilitla (San Luis Potosí) at km 233 on Federal Highway 120. The outcrops form low hills covered by red soil that are visible on both sides of the highway.

We measured a section (65 m) and collected 12 limestone samples at an average interval of 4 to 5 m (Figure 2). The top of the section is capped by red soil. For micropaleontological and microfacies analysis, the samples were prepared in thin sections 50 µm thick. The faunal preservation is good, which permitted a precise identification of foraminifers and algae for age assignment. In addition, a microfacies study was carried out for re-

constructing the paleoenvironment.

4. Results and Discussion

4.1. FORAMINIFERAL ASSEMBLAGE AND AGE

The foraminiferal assemblage is composed mostly of the orbitolinids *Dictyoconus walnutensis* (Carsey, 1926), *Paracoskinolina sunnilandensis* (Maync, 1955b), *Coskinolinoides texanus* (Keijzer, 1942), which are helpful for having a short stratigraphic range. Other foraminifers identified were *Buccicrenata subgoodlandensis* (Vanderpool, 1933), *Novalesia angulosa* (Magniez, 1972), *Mayncina* sp., *Istriloculina* sp., *Arenobulimina* sp., *Buccicrenata* sp. and the planktics *Favusella washitensis* (Carsey, 1926), *Muricohedbergella simplex* (Morrow, 1934), and *M. albiana* (BouDagher-Fadel *et al.*, 1996) (Figures 3-5).

The algae are *Cayeuxia kurdistanensis* Elliott, 1957, *Terquemella americana* (Konishi and Epis, 1962), *Triploporella* sp., *Acroporella* sp. and the encruster *Lithocodium aggregatum* Elliot, 1956 (Figure 6).

The benthic foraminifera provide a useful tool for age assignment using the stratigraphic ranges of the age diagnostic species.

The most important foraminifers used for dating the El Madroño succession are the orbitolinids *Dictyoconus walnutensis* (Carsey, 1926), *Coskinolinoides texanus* (Keijzer, 1942), *Paracoskinolina sunnilandensis* (Maync, 1955b), which support a mid-early late Albian age considering that the studied succession precedes the global sea-level rise that took place in the late Albian (Grötsch *et al.*, 1993; Filkorn and Scott, 2011). The occurrence of other species such as *Buccicrenata subgoodlandensis* (Vanderpool, 1933), *Novalesia angulosa* (Magniez, 1972); *Istriloculina* sp., *Buccicrenata* sp., *Arenobulimina* sp. and the *r* strategist planktics *Favusella washitensis* (Carsey, 1926), *Muricohedbergella albiana* (BouDagher-Fadel *et al.*, 1996) and *Muricohedbergella simplex* (Morrow, 1934) is in accord with the determined age.

Dictyoconus walnutensis (Figure 3a-c) was described by Carsey (1926) as *Orbitolina* from the Walnut Clay of the Cretaceous of central Texas; Silvestri

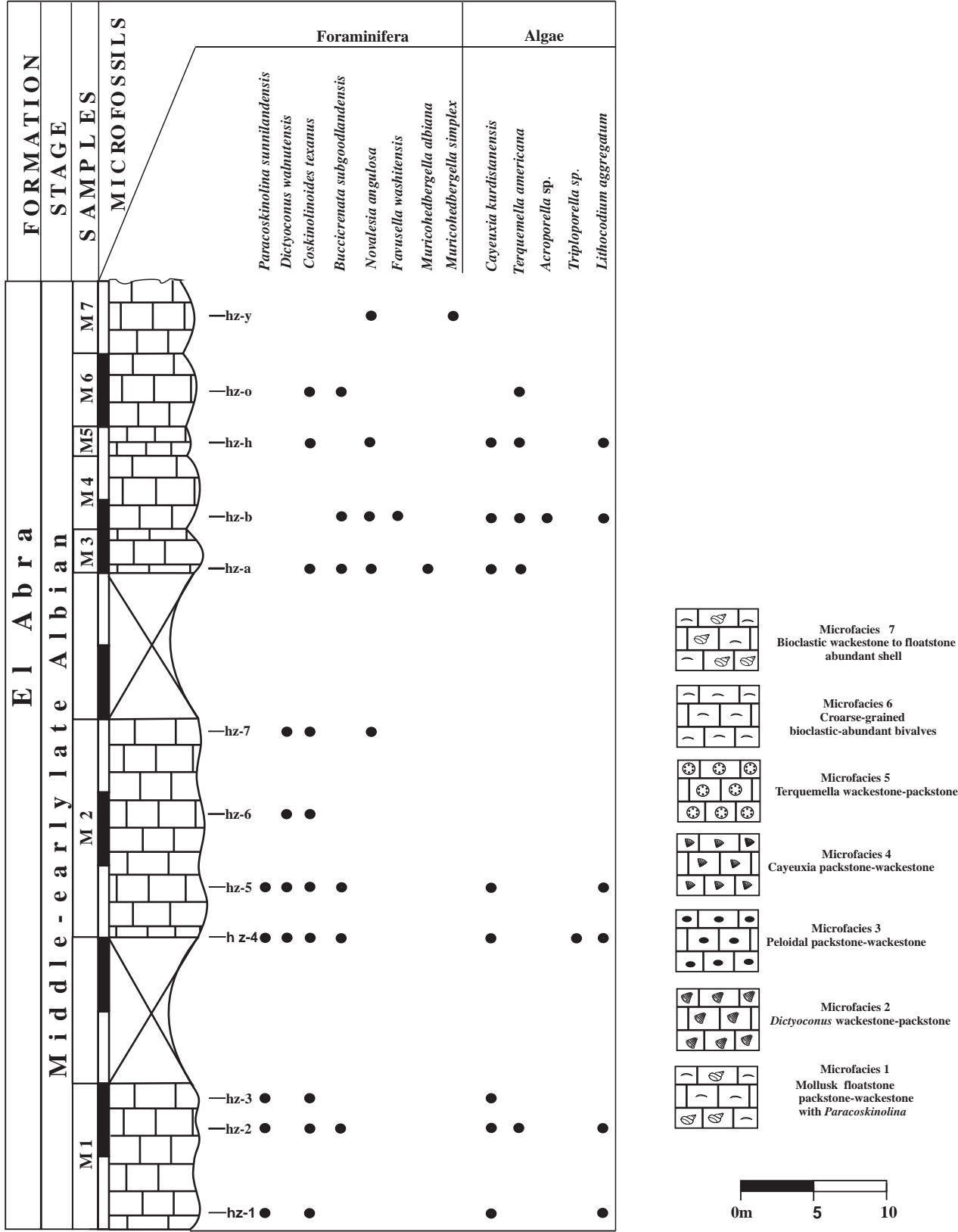


Figure 2 Composite stratigraphic column.

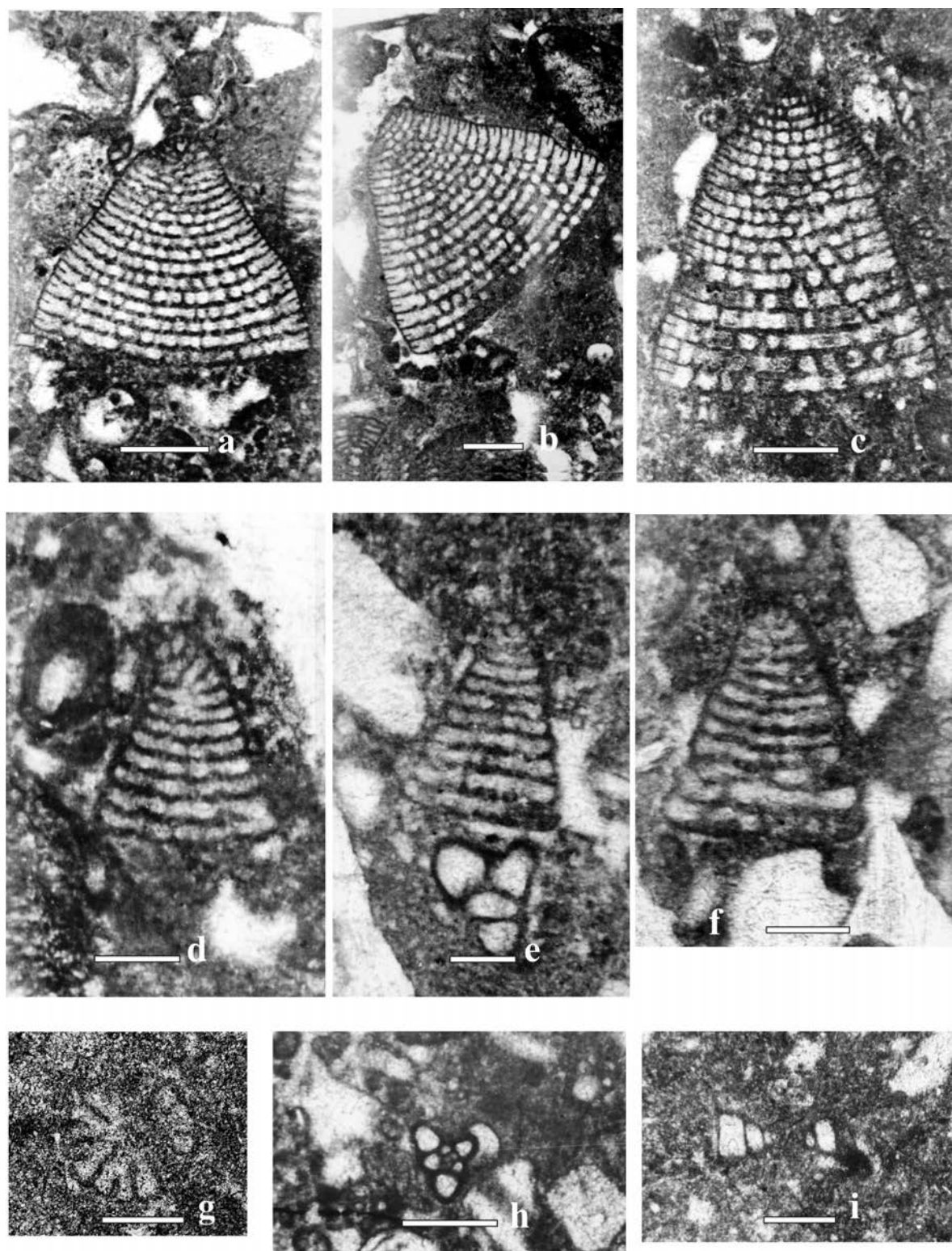


Figure 3 Mid-early late Albian foraminifera from the El Abra Formation. a- Axial section of *Dictyoconus walnutensis* (Carsey, 1926) hz-4 (scale bar 500 μ m). b- Subaxial section *Dictyoconus walnutensis* (Carsey, 1926) hz-4 (scale bar 500 μ m). c- Subaxial section *Dictyoconus walnutensis* (Carsey, 1926) hz-5 (scale bar 500 μ m). d, e, f- Axial and transverse section of *Coskinolinoides texanus* (Keijzer, 1942) hz-5 (scale bar 200 μ m). g- Horizontal section of *Coskinolinoides texanus* (Keijzer, 1942) hz-2 (scale bar 300 μ m). h- Miliolid hz-5 (scale bar 300 μ m). i- Axial section *Spiroloculina* hz-5 (scale bar 300 μ m).

(1932, p. 377 in Douglass, 1960) subsequently re-assigned it to the genus *Dictyoconus*. Barker (1944, p. 206) found this species only in the Comanche Peak of the Austin region, it also was recorded from the Fredericksburg (Scott, 1990; Scott *et al.*, 2003) in Texas. In addition this species has been identified in the subsurface of Florida (Applin and Applin, 1965). Maync (1955a) recorded the species in the Lower Cretaceous (Guácharo limestone) of eastern Venezuela and prepared a detailed description. *Dictyoconus walnutensis* is widespread in the Gulf of Mexico region and has been regarded as a stratigraphic marker restricted to the mid-Albian (Douglass, 1960; Maync, 1955a; Bonet, 1956; Scott, 1990; Scott and González-León, 1991; Buitrón *et al.*, 1995; Omaña and Alencaster, 2001; Monreal and Longoria, 1999; Fourcade *et al.*, 1999; Waite *et al.*, 2007); however, it could have ranged to the early late Albian (Phelps *et al.*, 2015). In North Texas this species ranges to the top of Goodland Formation in which the basal upper Albian ammonite *Dipoloceras cristatum* occurs; so *D. walnutensis* ranges into basal upper Albian (written communication Scott, 2016).

In Mexico, *Dictyoconus walnutensis* (Carsey, 1926) is broadly distributed. It has been reported from the following formations: Finlay (Monreal and Longoria, 1999); Espinazo del Diablo (Scott and González-León, 1991); El Abra (Bonet, 1956, Buitrón *et al.*, 1995; Rosales-Domínguez, 1998; Omaña and Alencáster, 2001); Escamela (Bonet, 1956), and Morelos (Flores de Dios *et al.*, 2004) in the states of Chihuahua, Sonora, Querétaro, Veracruz and Guerrero respectively.

Paracoskinolina sunnilandensis (Figure 4a-g) has been described from the Lower Cretaceous (Albian) Sunniland limestone of Florida, as *Coskinolina* by Maync (1955b), and later transferred to the genus *Paracoskinolina* by Moullade (1965). Maync (1955b) indicated that he observed this foraminifer from the *Dictyoconus*-bearing rock of Guácharo, a member of the Chimama Formation in eastern Venezuela which was dated as middle Albian (Rod and Maync, 1954). Scott and González-León (2004) identified *P. sunnilandensis* (Maync, 1955b) from the

Mural Limestone (Sonora); in Chiapas it has been recorded as Albian from the Sierra Madre Formation (Michaud, 1987). Arnaud-Vanneau and Sliter (1995 p. 544) also found the species at Site 866 on Resolution Guyot in the Pacific Ocean. According to Maync (1955b, p. 106) the species is known from Switzerland. In France the stratigraphic distribution of *Paracoskinolina* cf. *P. sunnilandensis* is early Barremian from the Hugii and Nicklesi Zones (Clavel *et al.*, 2010).

Coskinolinoides texanus (Figure 3 d-f) was originally described from the Walnut clay, of Lower Cretaceous (Fredericksburg) west of Austin, Texas by Keijzer in 1942, associated with *Dictyoconus walnutensis* (Carsey, 1926) is considered to characterize the Comanche Peak in central Texas (Lozo, 1944, p. 570). Applin and Applin (1965) indicated that in Florida “the species seems to be limited in its vertical range to the beds of Fredericksburg age, and specimens generally occur in greatest abundance in the upper part of the unit. Frizzell (1954, p. 76), reported that in Texas, *C. texanus* ranges upward from the Glen Rose Limestone of the Trinity Group into the progressively younger Walnut Clay, Comanche Peak Limestone, and Goodland Limestone of the Fredericksburg Group”. Phelps *et al.* (2015) indicated that this species ranges from the mid-to-upper Albian from west Texas. In Mexico, *C. texanus* has been reported by Scott and González-León (1991), Rosales-Domínguez (1998), Omaña and Alencaster (2001), Flores de Dios *et al.* (2004); and in the Dominican Republic the occurrence of *C. texanus* suggests that the Hattillo Limestone spans to the middle Albian (Bonilla-Rodríguez, 2013).

We identified other foraminifers such as *Ammobaculites subgoodlandensis* (Figure 5a), which was defined by Vanderpool (1933) from the mid-Albian Comanchean “formation” in southern Oklahoma, later placed in the genus *Buccicrenata* by Loeblich and Tappan (1949).

Filkorn and Scott (2011) found *Buccicrenata subgoodlandensis* (Vanderpool, 1933) in the Mal Paso Formation in Guerrero state (Mexico) with a foraminiferal association dated as late Albian; they

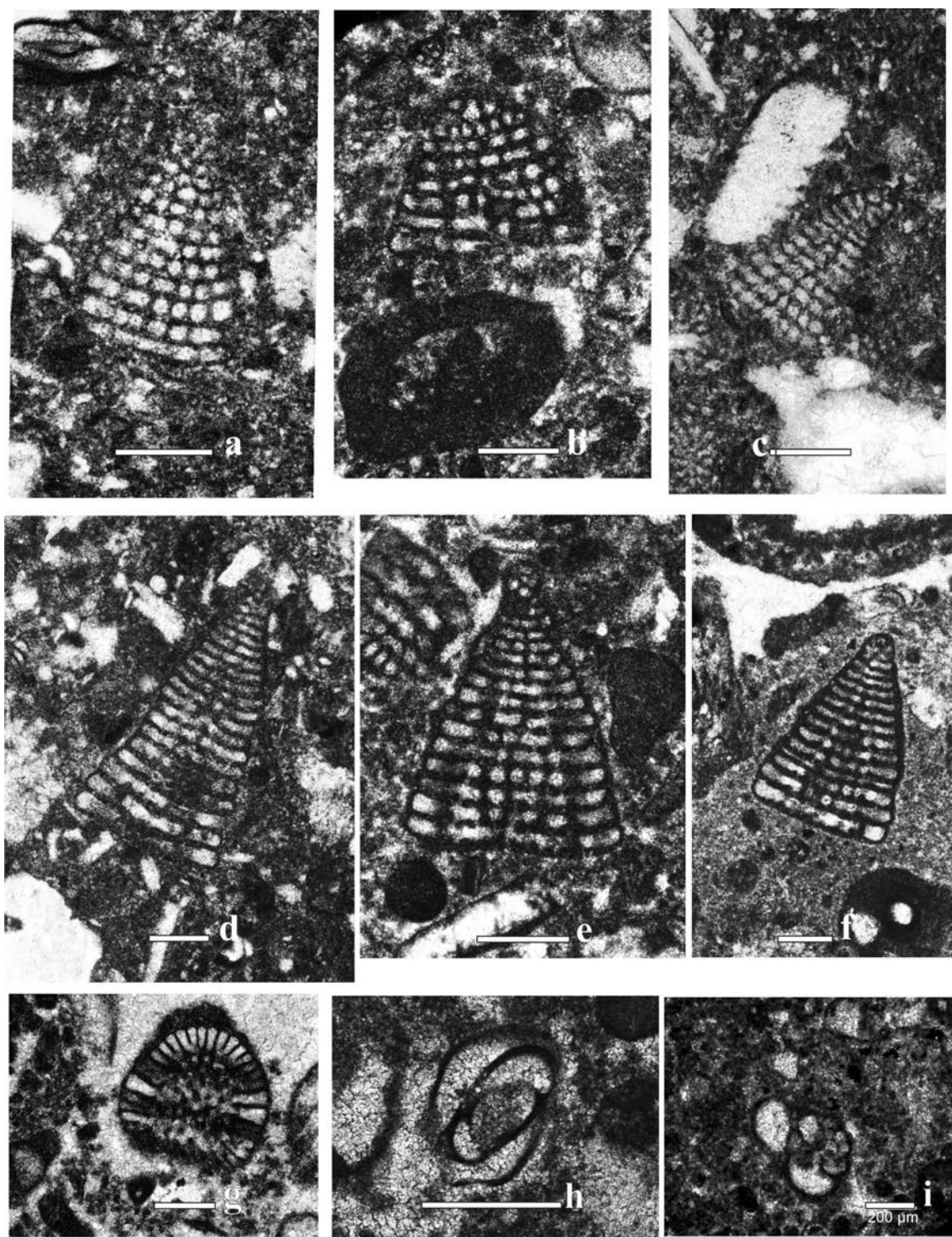


Figure 4 Mid-early late Albian foraminifera from the El Abra Formation. a- Subaxial section *Paracoskinolina sunnilandensis* (Maync, 1955b) hz-3 (scale bar 300 µm). b- Subaxial section of *Paracoskinolina sunnilandensis* (Maync, 1955b) hz-2 (scale bar 300 µm). c- Subaxial section of *Paracoskinolina sunnilandensis* (Maync, 1955b) hz-3 (scale bar 300 µm). d- Subaxial section of *Paracoskinolina sunnilandensis* (Maync, 1955b) hz-3 (scale bar 500 µm). e- Axial section of *Paracoskinolina sunnilandensis* (Maync, 1955b) hz-1 (scale bar 500 µm). f- Axial section of *Paracoskinolina sunnilandensis* (Maync, 1955b) hz-2 (scale bar 500 µm). g- Transverse section of *Paracoskinolina sunnilandensis* (Maync, 1955b) (scale bar 500 µm). h- *Istriloculina* sp. Sample hz-2 (scale bar 200 µm). i- *Arenobulimina* sp. Sample hz-b (scale bar 200 µm).

indicated that the stratigraphic range of this species extends throughout the Albian in Tethyan sections from North America and the Middle East. In Texas, it occurs in the lower Albian Glen Rose Formation and the lower to upper Albian Walnut and Goodland formations of the Fredericksburg Group. These authors note that *B. subgoodlandensis* has been reported as *P. hedbergi* from mid-upper Albian strata in Mexico (Scott and González-León, 1991) and Honduras (Scott and Finch, 1999). In Florida it was identified as *Lituola subgoodlandensis*, which is associated with *Coskinoloides texanus* (Keijzer) in the Fredericksburg Group (Applin and Applin, 1965). This species was recognized from the lower-upper Albian Nahr Umr Formation in the Arab Emirates and Oman (Banner and Highton, 1990; BouDagher-Fadel, 2001). In our material it is associated with mid-Albian-early late Albian microfauna.

Novalesia angulosa (Figure 5d, e) was described from the Albian of the Spanish Pyrenees as *Spiroplectamminoides* by Magniez 1972, and later assigned to *Novalesia* (Magniez, 1974). The known range is late Aptian to Albian (Arnaud-Vanneau and Sliter, 1995). This species is common in the El Madroño samples.

We identified a specimen of *Favusella washitensis* (Figure 5g), which was reported by Carsey (1926) as *Globigerina washitensis* from the Del Rio Formation of Texas, USA, and later placed in the subgenus *Favusella* by Michael (1972). The stratigraphic range of the species spans from the Albian to early Cenomanian (Premoli-Silva and Sliter, 2002).

Muricohedbergella albiana (BouDagher-Fadel *et al.*, 1996) (Figure 5i) extends from the latest early Albian to late Albian *Ticinella primula* Zone to the *Rotalipora appenninica* Zone (Premoli-Silva and Verga, 2004; Huber, 2006); however, it has been also reported by Magniez-Jannin (1975) from mid-Albian of the Aube, France.

We also recognized *Muricohedbergella simplex* (Morrow, 1934) (Figure 5h), which ranges from the upper part of the *Ticinella primula* Zone in the mid-Albian to Coniacian (Premoli-Silva and Verga, 2004). Although these planktic foraminifera

identified have a wide stratigraphic distribution they are in the interval of the assigned age.

4.2. PALEOECOLOGY

The Recent larger foraminifera are distributed in tropical and subtropical regions and they are most abundant in nutrient-deficient, warm, shallow-water seas (Boudagher-Fadel, 2008). Murray (2006) stated that “the controls on global distribution are thought to be water temperature, nutrient content, light intensity and hydrodynamic energy.” The light required by their symbionts led to the establishment of the depth distribution of the larger foraminiferal assemblage as a water-depth proxy indicating a shallow-water setting.

The ecology of larger foraminifera without equivalent recent forms could be inferred from the size, morphostructure, and test architecture. The dictyoconids such as *Dictyoconus*, *Paracoskinolina*, and *Coskinoloides* are characterized by having a conical test which suggests a free, epifaunal mode of life on soft substrates. This is confirmed by the wackestone textured rock where they were deposited, which permitted the inference that it occurred in a warm shallow-water platform with low energy in the euphotic zone, because they are associated with algae.

4.3. MICROFACIES AND PALEOENVIRONMENT

The paleoenvironmental interpretation is based on the foraminiferal-algae association and the microfacies study.

The thin section analysis of the textural features and faunal assemblage of the limestone enabled us to recognise several microfacies. They are described below from the base towards the top of the section. The succession is characterized by the occurrence of larger foraminifera such as orbitolinids, other small benthic foraminifera as *Novalesia* sp., *Istriloculina* sp., *Arenobulimina* sp. and planktic organisms, as well as fragments of rudists, gastropods, echinoid, and algae and encrusters.

Microfacies 1 (Sample hz-1-3) Mollusk floatstone and foraminiferal packstone-wackestone is cha-

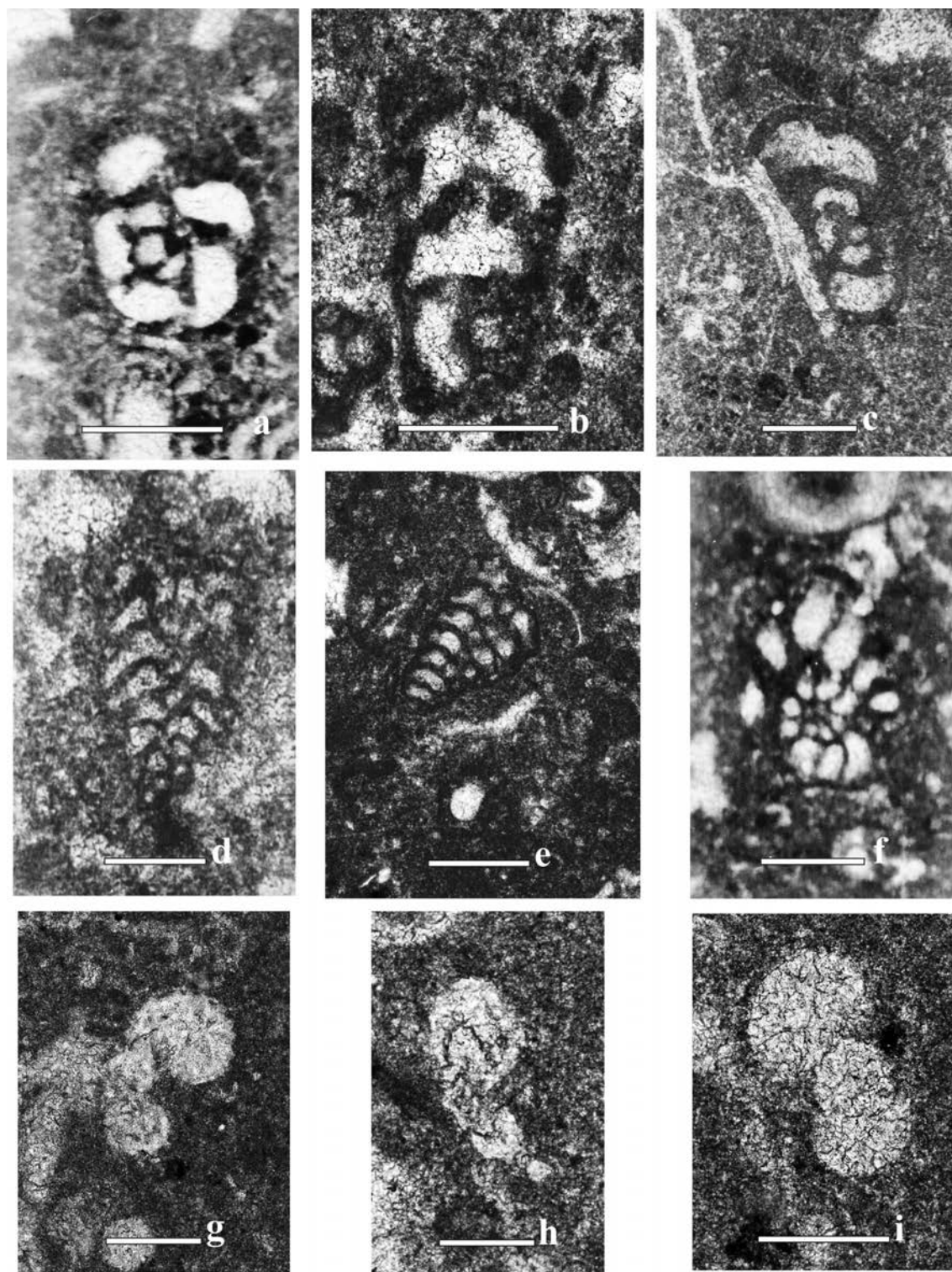


Figure 5 Mid-early late Albian foraminifera from the El Abra Formation. a- Equatorial section of *Buccicrenata subgoodlandensis* (Vanderpool, 1933) hz-4 (scale bar 300 µm). b- Subequatorial section of *Buccicrenata* sp. hz-2 (scale bar 300 µm). c- Subaxial section of *Buccicrenata* sp. hz-a (2) (scale bar 300 µm). d- *Novallesia angulosa* (Magniez, 1972) hz-a (1) (scale bar 300 µm). e- *Novallesia angulosa* (Magniez, 1972) hz-6 (scale bar 500 µm). f- Equatorial section of *Mayncina* sp. hz-6 (scale bar 300 µm). g- Axial section of *Favusella washitensis* (Carsey, 1924) hzb (2) (scale bar 200 µm). h- Axial section of *Muricohedbergella simplex* (Morrow, 1934) hzy (2) (scale bar 100 µm). i- Subaxial section of *Muricohedbergella albiana* (BouDagher-Fadel *et al.*, 1996) hza (2) (scale bar 200 µm).

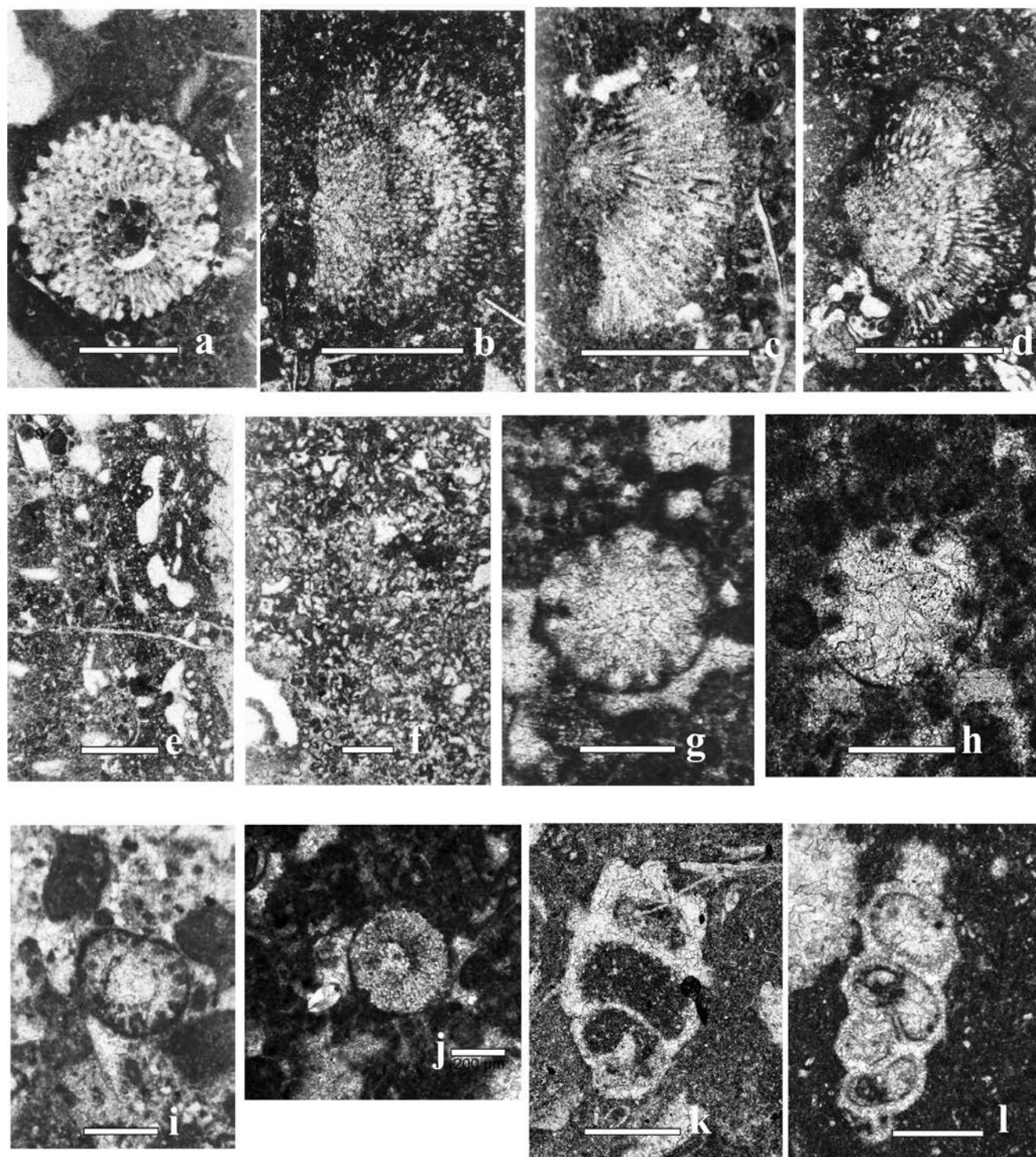


Figure 6 Mid-early late Albian algae from the El Abra Formation. a- *Triploporella* sp. hz-4 (scale bar 1 mm). b- Longitudinal section of *Cayeuxia kurdistanensis* Elliott, 1957 hza (3) (scale bar 1 mm). c- Longitudinal section of *Cayeuxia kurdistanensis* Elliott, 1957 hz-5 (scale bar 1 mm). d- Longitudinal section of *Cayeuxia kurdistanensis* Elliott, 1957 hzb (a) (scale bar 1 mm). e- *Lithocodium aggregatum* Elliott, 1956, hzb (A) (scale bar 1 mm). f- *Lithocodium aggregatum* Elliott, 1956, hzb (A) rojo (scale bar 2 mm). g- Transverse section of *Terquemella americana* (Konishi and Epis, 1962) hza (2) (scale bar 300 μ m). h- Transverse section of *Terquemella americana* (Konishi and Epis, 1962) hz-2 (scale bar 300 μ m). i- Transverse section of *Acroporella* sp. hzb (3) (scale bar 300 μ m). j- Echinoderm hzb (1) (2) (scale bar 200 μ m). k- Gastropod hzo (2) (scale bar 500 μ m). l- Gastropod hzh (2) (scale bar 500 μ m).

characterized by a foraminiferal association composed mostly of *Paracoskinolina sunnilandensis* (Maync, 1955b). Other foraminifers present are *Novalesia angulosa* (Magniez, 1972), *Buccicrenata* sp. The algae are *Cayeuxia kurdistanensis* Elliott, 1957; *Terquemella americana* (Konishi and Epis, 1962), and the encruster *Lithocodium aggregatum* Elliott, 1956. This interval also contains gastropods, and corals (Figure 7a). Microfacies 2 (Samples hz-4-7) is a foraminiferal packstone–wackestone showing a more diversified association, contains *Dictyoconus walnutensis* (Carsey, 1926) and *Coskinolinoides texanus* (Keijzer, 1942). Also present are *Paracoskinolina sunnilandensis* (Maync, 1955b); *Buccicrenata subgoodlandensis* (Vanderpool, 1933); *Novalesia angulosa* (Magniez, 1972); the algae *Cayeuxia kurdistanensis* Elliott, 1957, *Triploporella* sp. and the encruster *Lithocodium aggregatum* Elliott, 1956 as well as gastropods and fragments of bivalves (Figure 7b,c).

Microfacies 3 (Sample hz-a) Peloidal packstone–wackestone containing benthic foraminifers *Paracoskinolina sunnilandensis* (Maync, 1955b), *Coskinolinoides texanus* Keijzer, 1942, *Novalesia angulosa* (Magniez, 1972), *Buccicrenata* sp. and algae *Terquemella americana* (Konishi and Epis, 1962), *Cayeuxia kurdistanensis* Elliott, 1957 and the planktic foraminifer *Muricohedbergella albiana* (BouDagher-Fadel *et al.*, 1996) and fragments of mollusks (Figure 7d).

Microfacies 4 (Sample hz-b) Algal packstone–wackestone containing *Paracoskinolina*, *Arenobulimina* sp., *Cayeuxia kurdistanensis* Elliott, 1957 *Lithocodium aggregatum* Elliott, 1956 and the planktic foraminifer *Favusella washitensis* (Carsey, 1926) (Figure 8a).

The depositional environment of microfacies 1 – 4 suggests a shallow marine environment within the euphotic zone in the open marine interior platform, corresponding to Facies Zone 7 (FZ 7) (Flügel, 2004, p. 663). The fossil assemblage indicates deposition in the subtidal zone with normal salinity, stable temperature conditions and good oxygenation of the seawater.

Microfacies 5-(Sample hz-h) Shell wackestone–packstone with foraminifers *Coskinolinoides texanus* Keijzer, 1942, *Novalesia* sp. and algae *Terquemella americana* (Konishi and Epis, 1962), and *Lithocodium*

aggregatum Elliott, 1956 (Figure 8b).

Microfacies 6-(Sample hz-o) Coarse-grained bioclastic-wackestone, partly floatstone, containing abundant bivalve shell fragments of mollusks, gastropods, echinoids, with scarce foraminifers, and algae (Figure 8c).

Microfacies 7-(Sample hz-y) Bioclastic wackestone to floatstone containing abundant mollusk fragments, gastropods, echinoderms, and benthic foraminifers such as *Novalesia* sp. and planktic *Muricohedbergella simplex* (Morrow, 1934) (Figure 8d).

Microfacies 5 – 7 are characterized by floatstone composed of accumulations of abundant shell fragments probably caused by current concentrations, indicating deposition in an open-shelf setting with moderate energy and could correspond to Facies Zone 5 (FZ 5) of Flügel (2004).

The microfacies and the faunal association including Dictyoconinae are essentially linked to fine carbonate muddy deposits in shallow environments as has been proposed by Masse (1992) which is confirmed in the studied samples. On the other hand, the occurrence of gastropods and broken mollusk fragments in the wackestone–packstone textured limestone suggests a shallow, open-marine platform with two events, one of low energy and the other of moderate energy. The presence of algae indicates that the deposit took place within the photic zone. The existence of the r strategist planktic species in the association demonstrates their adaptation to a shallow-water environment, although the presence of several levels of floatstone and the scarce occurrence of these planktic foraminifers also could indicate a trend in the environment's evolution towards the platform border.

4.4. PALEOBIOGEOGRAPHY

Dictyoconus walnutensis (Carsey, 1926), *Coskinolinoides texanus* (Keijzer, 1942) and *Paracoskinolina sunnilandensis* (Maync, 1955b) are common species in the Gulf of Mexico area; for instance, in the USA, in Texas, New Mexico and Florida from the Fredericksburg Group, Finlay and Sunniland formations (Scott, 1990; Scott, *et al.*, 2003; Lucas *et al.*, 2010),

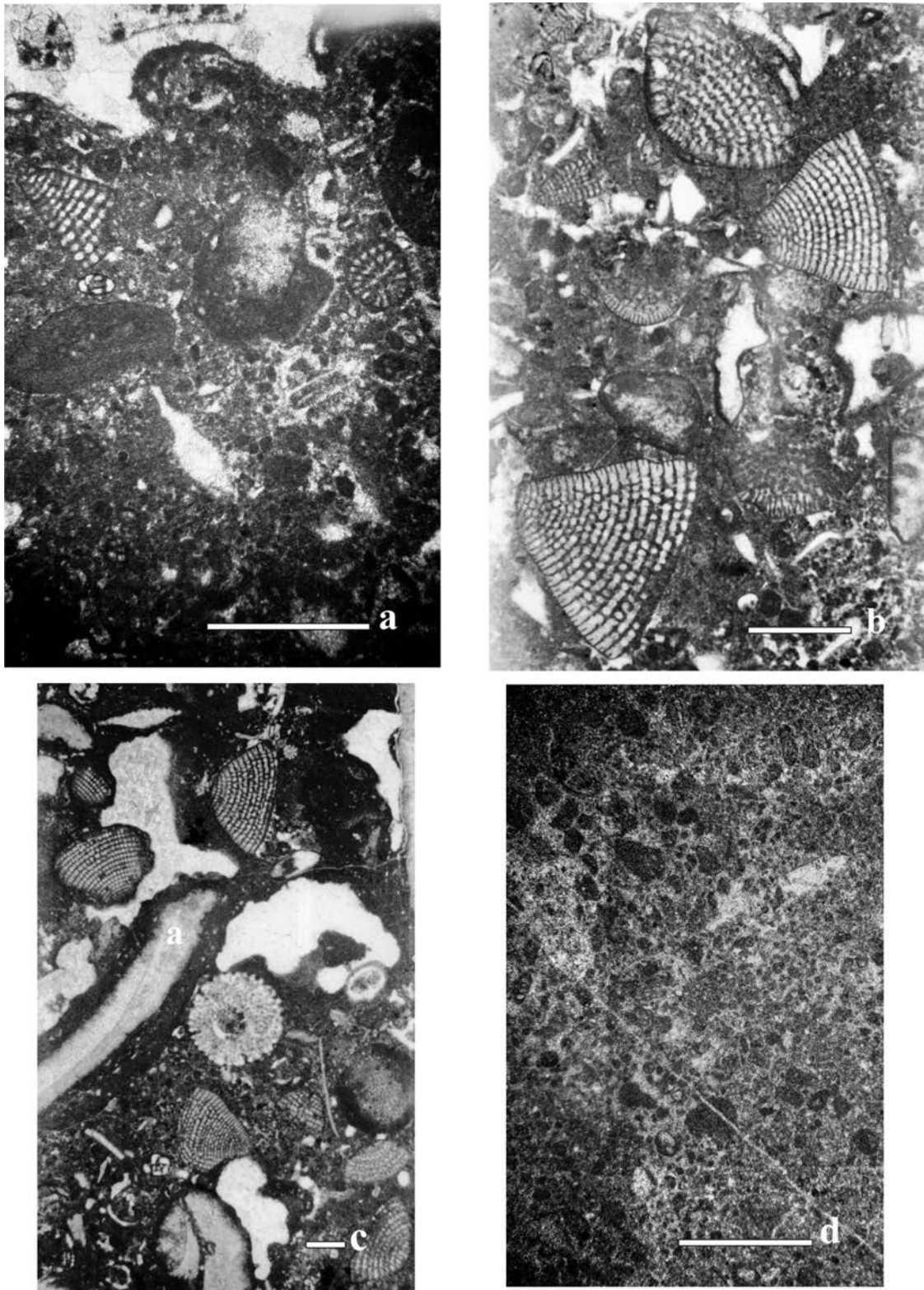


Figure 7 Microfacies from the El Abra Formation in the El Madroño locality. a– Microfacies 1 Mollusk floatstone and foraminiferal packstone–wackestone (hz–1) with *Paracoskinolina sunnilandensis* (scale bar 1 mm). b, c Microfacies 2 (Samples Sample hz–4–7) Foraminiferal packstone–wackestone with *Dictyoconus walnutensis*, *Coskinolinoides texanus*, *Buccicrenata subgoodlandensis* and *Triploporella* sp. (scale bar 1 mm). d– Microfacies 3 (Sample hz–a) Peloidal packstone–wackestone with foraminifers (scale bar 1 mm).

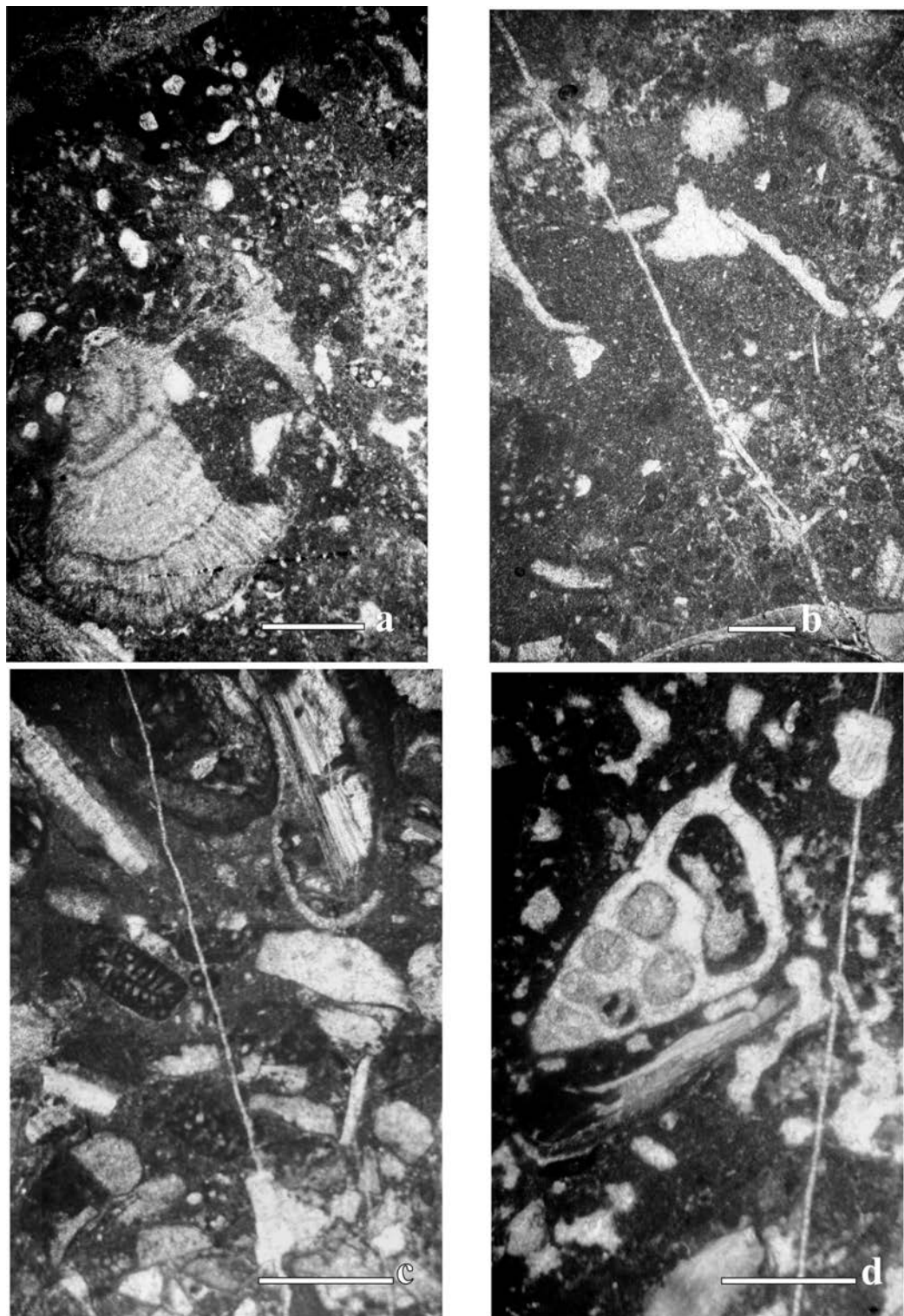


Figure 8 Microfacies from the El Abra Formation in the El Madroño locality. a- Microfacies 4 (Sample hz-b) Algal packstone-wackestone containing *Paracoskinolina*, *Arenobulimina* sp., *Cayeuxia kurdistanensis* Elliott, 1957 (scale bar 500 μ m). b- Microfacies 5 (Sample hzh) Shell wackestone-packstone with *Coskinolinoides texanus*, *Novalesia* sp. and *Terquemella americana* (Konishi and Epis, 1962) (scale bar 500 μ m). c- Microfacies 6 (Sample Sample hz-o) Coarse-grained bioclastic wackestone, partly floatstone with scarce foraminifers (scale bar 1 mm). d- Microfacies 7-(Sample hz-y-1) Bioclastic wackestone to floatstone containing abundant mollusk fragments (scale bar 1 mm).

and the subsurface in the Dollar Florida Bay Formation (Bozkurt, 2015). In Mexico, this orbitolinid association is extensively distributed in the Finlay, Espinazo del Diablo, El Abra, Escamela, Morelos and the Sierra Madre formations. In addition, it is known in the Caribbean region of Guatemala in the Cobán Formation (Fourcade *et al.*, 1999); in the Palenque Formation in Cuba (Meyerhoff and Hatten, 1974, p. 440); in Honduras in the Atima limestone (Scott and Finch, 1999); in Venezuela from the Guácharo limestone (Maync, 1955 a, b) and the Dominican Republic in the Hatillo Formation (Bonilla-Rodríguez, 2013; Bonilla *et al.*, 2014). These species have been also identified by Granier (2005) from samples collected from the Bahamas Escarpment. Cherchi and Schroeder (1984) reported *Dictyoconus walnutensis* (Carsey, 1926) and *Paracoskinolina sunnilandensis* (Maync, 1955b) in the southwestern part of the Gulf of Mexico (Site 77) as reworked material.

According to the geographical distribution of these orbitolinids in the mid-Cretaceous (Albian) we consider that they are endemic forms for the Gulf Coast and Caribbean region into the western part of Tethys. This is in agreement with the view of Cherchi (2004) who “indicated that *Dictyoconus walnutensis* (Carsey, 1926) and *Paracoskinolina sunnilandensis* (Maync, 1955b) are restricted to the Gulf region; thus records of these taxa in the European Tethyan realm are erroneous”. However *Paracoskinolina* cf. *P. sunnilandensis* (Maync, 1955b) have been recorded in older rocks dated as early Barremian from the Urganian platform of SE France (Clavel *et al.*, 2010).

In addition, the other foraminifers *Buccicrenata subgoodlandensis* (Vanderpool, 1933), *Novalesia angulosa* (Magniez, 1972), *Arenobulimina* sp. and the planktic foraminifers have a wide distribution in the Tethys Realm.

5. Conclusions

The foraminiferal assemblage contained in the El Abra succession sampled from the El Madroño site

consists of the benthic foraminifers *Dictyoconus walnutensis* (Carsey, 1926), *Paracoskinolina sunnilandensis* (Maync, 1955b45), *Coskinolinoides texanus* (Keijzer, 1942), *Buccicrenata subgoodlandensis* (Vanderpool, 1933), *Novalesia angulosa* (Magniez, 1972), *Arenobulimina* sp. and the algae species, *Cayeuxia kurdistanensis* Elliot, 1957, *Terquemella americana* (Konishi and Epis, 1962), *Triploporella* sp., *Acroporella* sp. and the encruster *Lithocodium aggregatum* Elliot, 1956. This assemblage is dated as mid-early late Albian based mostly on the *Dictyoconus*, *Paracoskinolina* and *Coskinolinoides* assemblage.

The microfacies change from wackstone-packstone to grainstone toward a floatstone indicating a shallow, open-marine environment with two phases, one of low energy and the other of moderate energy, which is inferred on the basis of the limestone texture and the benthic foraminiferal and algal association.

The foraminifers analyzed in this study are widely distributed in the shallow-water deposits of Mexico, the Gulf Coast of the United States and the Caribbean region in the western part of the Tethys Realm.

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