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Recent carbonatitic magmatism in Angola: the dykes of the Chiva lagoon maar

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ABSTRACT

Chiva is a circular lagoon with a diameter of 900 m, located in the vicinity of the Catanda graben (Angola). This area is largely covered by Quaternary lacustrine sediments, however, a few dykes of carbonatitic composition have been found near the SE shore of the lagoon. The intrusion of these carbonatitic dykes has been linked to the occurrence of recent volcanic events in the Chiva area and interpreted as feeding channels of an eruptive centre with maar morphology. The Chiva carbonatitic dykes show a finely porphyritic texture formed by phenocrysts of subhedral calcite (60 %), apatite (30 %) and magnetite (10 %) hosted in a fine calcite-rich ground mass that also contains accessory minerals such as pyrochlore, perovskite and kimzeyite. The mineralogical and compositional features of the Chiva dykes are similar to those described in some of the carbonatitic lavas from the Catanda graben, suggesting that both volcanic events are genetically related. The finding of this new eruptive centre enlarges the carbonatitic domain of Catanda and opens the door to the discovery of new carbonatitic outcrops in this region of Western Angola.

Keywords: carbonatites, carbonatitic dykes, maar, Angola, Catanda, Chiva

RESUMEN

En las inmediaciones del graben de Catanda (Angola) se localiza una laguna circular de 900 metros de diámetro conocida como laguna de Chiva. Aunque éste área se encuentra mayoritariamente cubierta por sedimentos lacustres, hay afloramientos de un pequeño grupo de diques de composición carbonatítica en la orilla SE de la laguna. La intrusión de estos diques carbonatíticos se relaciona con el desarrollo de eventos volcánicos en la zona de Chiva y se interpreta como los antiguos canales de alimentación de un centro eruptivo que desarrolló una morfología de maar. Los diques carbonatíticos de Chiva presentan una típica textura porfídica formada por fenocristales subhedrales de calcita (60 %), apatito (30 %) y magnetita (10 %) implantados en una matriz de calcita de grano fino con cantidades accesorias de pirocloro, perovskita y kimzeyita. Las características mineralógicas y composicionales de los diques del maar de Chiva son equivalentes a las descritas en las lavas carbonatíticas del graben de Catanda, lo que sugiere que ambos procesos volcánicos están relacionados genéticamente. El hallazgo de este nuevo centro eruptivo amplía la zona de desarrollo del vulcanismo carbonatítico en Catanda y abre la puerta al posible descubrimiento de nuevos afloramientos carbonatíticos extrusivos en esta zona occidental de Angola.

Palabras clave: carbonatitas, diques carbonatíticos, maar, Angola, Catanda, Chiva.

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1. Introduction

Carbonatites are igneous rocks composed of more than 50 % of carbonates and less than 20 wt% of SiO_2 (Le Maitre, 2002). However, their classification is still under discussion and many authors consider that magmatic rocks should be classified as carbonatites when they contain more than 30 % of modal carbonates, not including the SiO_2 content (Mitchell, 2005). Carbonatites can be plutonic, hypabyssal or volcanic depending on their emplacement. They are typically associated to rift structures that are developed inside continental cratons (Le Bas, 1986), although they can be also related to other geological settings such as oceanic islands (Mangas *et al.*, 1996; Widom *et al.*, 1999) or orogenic belts (Tilton *et al.*, 1998; Chakhmouradian *et al.*, 2008).

According to experimental petrology data and isotopic studies of stable and radiogenic isotopes, it is considered that carbonatitic magmas are originated in the lithospheric mantle (Wallace *et al.*, 1988; Bell and Simonetti, 2010). Carbonatitic melts are marked by their high content in volatile elements, especially CO_2 but also F, and their enrichment in strategic elements such as Rare Earth Elements, Nb or Ta (Wyllie *et al.*, 1996; Chakhmouradian, 2006; Chakhmouradian and Wall, 2012). This point has generated a great interest in the study of carbonatites despite their rarity since only 527 carbonatite localities have been reported in the world. An important part of the research studies in carbonatites are focused on the study of extrusive localities, since volcanic products can provide very significant information about the genesis and the composition of carbonatitic magmas and the understanding of experimental petrology data.

Extrusive carbonatites are extremely rare rocks, only 47 localities have been reported in the world until present-day (Woolley and Church, 2005; Woolley and Kjarsgaard, 2008). The majority of extrusive carbonatites are located in Africa and especially in the Eastern Rift where the only occurrence of an active carbonatitic volcano is found, the Ol Doinyo Lengai in Tanzania (Daw-

son, 1962; Dawson *et al.*, 1995; Keller and Zaitsev, 2012). In the western margin of the African continent a magnificent example of carbonatitic volcanism is also reported: the volcanic area of Catanda (Kwanza Sul, Angola). It is formed by a group of small monogenetic volcanic buildings of carbonatitic composition cropping out in a 50 km² graben hosted in Archean granites and it is delimited by the intersection of three different fault systems. The Catanda carbonatitic volcanoes are largely eroded and covered by Quaternary colluvial and epiclastic sediments, which partly obscure the original morphology of the volcanic edifices. However, seven different eruptive centres were distinguished, associated to tuff ring and maar morphologies (Figure 1; Campeny *et al.*, 2014).

The emplacement of the Catanda volcanic carbonatites is associated to the extensional domain of the Lucapa belt, which is a rift corridor defined by the NE-SW trending fractures of the Quilengues-Andulo fault system (Jelsma *et al.*, 2009). The first magmatic activity reported in the Lucapa corridor is from the Paleoproterozoic (Sykes *et al.*, 1978) but magmatism was also significantly important in the Upper Cretaceous, associated to the break-up of Gondwana and the corresponding opening of the South Atlantic Ocean (Moulin *et al.*, 2010). Most of the Angolan carbonatites and kimberlites are distributed along Lucapa (Alberti *et al.*, 1999) and were emplaced during the Cretaceous magmatic event (Torró *et al.*, 2012; Calvo *et al.*, 2011a; Calvo *et al.*, 2011b; Robles-Cruz *et al.*, 2012). Silva and Pereira (1973) and Torquato and Amaral (1974) reported a Cretaceous age for the Catanda carbonatitic volcanism of 92 ± 7 Ma based on the K/Ardating of phlogopite from a tinguaita dyke outcropping in the area. However, the well-preserved morphology of the volcanic edifices, the nature of the pyroclastic materials and the existence of present-day geothermal and seismic activity in the same area suggest a recent age for these eruptive processes (Campeny *et al.*, 2014; Campeny, 2016).

During the study of satellite images of the Catanda volcanic area we localised a small circular

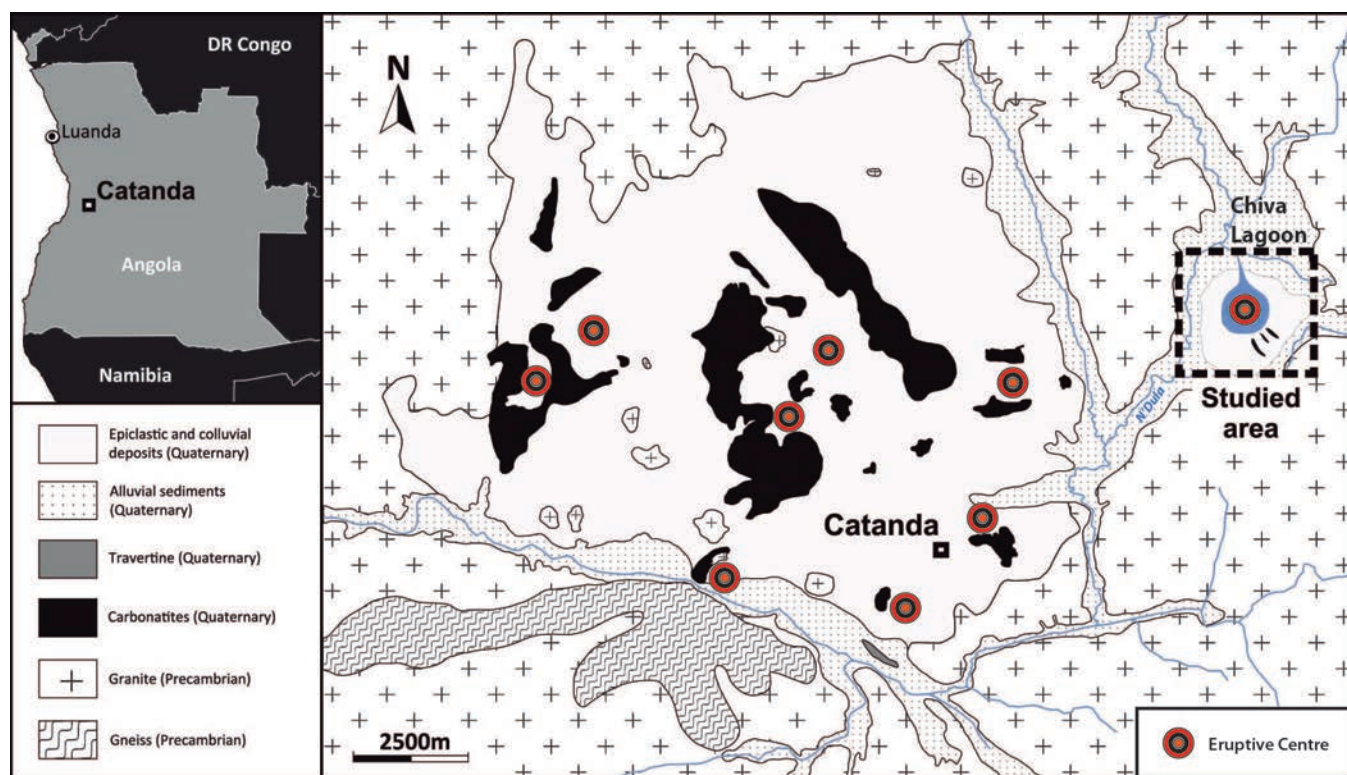


Figure 1 Geological map of the Catanda graben and the area of the Chiva lagoon (modified after Campeny *et al.*, 2014).

lagoon in the Chiva area, located outside of the Catanda graben and slightly far away from the main carbonatite outcrops. Field work carried out in the surroundings of this lagoon allowed us to distinguish carbonatitic dykes, thus indicating the possible relation of this area with the occurrence of the carbonatitic volcanism in Catanda (Campeny *et al.*, 2014; Campeny *et al.*, 2015). The aim of this work is to provide new information about the nature of the materials of the Chiva lagoon and to discern whether their formation is related to the carbonatitic volcanism of the Catanda region or whether it should be considered as a genetically independent event.

2. Methods

The petrographic study of carbonatitic rocks from the Chiva area was based on the identification of main rock-forming minerals and the description

of textural features. It has been carried out by the use of optical microscopy using transmitted and reflected light. Scanning Electron Microscopy E-SEM was also used for more detailed analyses of mineralogical relationships and textures, using a Quanta 200 FEI XTE 325/D8395BSE with a Genesis EDS microanalysis system from the Scientific and Technical Centres of the University of Barcelona (CCiTUB). The operating conditions of the SEM were 20 – 25 keV, 1 nA beam current and a working distance of 10 mm from the sample to the detector. Detail images of the textural patterns were obtained in back-scattered electron (BSE) mode.

Geochemical data of major and trace element compositions were obtained by performing X-ray fluorescence (XRF) (Norrish and Hutton, 1969) and ICP-emission spectrometry following a lithium metaborate/tetraborate fusion, respectively, at the ACTLABS Activation Laboratories Ltd. in Ancaster, Canada.

3. The Chiva Carbonatites

The Chiva lagoon, which is considered to be the source of the N'Dula river, is located 350 km SE of Luanda and 5 Km NE of the volcanic carbonatitic outcrops of Catanda (Figure 1). It forms a shallow depression of 900 metres in diameter, surrounded by a smooth relief that defines a typical ring shape (Figure 2a). In the vicinity of the lagoon, farming activity is developed and original outcrops are largely covered by alluvial and colluvial sediments. However, a few carbonatitic dykes can be distinguished on the SE edge of the lagoon, forming coherent 1.5-metre-thick bodies (Figure 2b), radially arranged from the lagoon centre to its border. Outcrops are brownish in colour when weathered and dark grey in fresh sample (Figure 2c).

3.1. PETROGRAPHY

Carbonatitic dykes from the Chiva lagoon have a porphyritic texture with 20 – 25 % of phenocrysts

and xenocrysts in a carbonate-rich aphanitic groundmass (Figure 3a), which also contains a large number of vesicles.

3.1.1. PHENOCRYSTS

The population of phenocrysts is made up of calcite, apatite and magnetite. Calcite represents around 60 % of modal phenocrysts of the Chiva dykes, such as reported in the Ipunda lavas from Catanda area (Table 1). Calcite crystals are subhedral, with a distinct tabular habit and attain up to 0.25 mm in length (Figure 3a). However, they have an internal polycrystalline texture, suggesting recrystallization and pseudomorphosis processes (Figure 3b). Apatite crystals are subhedral to euhedral, composed by prisms and hexagonal bipyramids; they are short prismatic in habit and may reach a diameter of up to 0.3 mm, representing 30 % modal of the total phenocryst population. They are slightly zoned, with rims enriched in britholite component; hence, the rims appear slightly brighter in SEM-BSE images than the cores of these crystals. Magnetite forms clearly zoned euhedral octahedral grains up to 0.4 mm in diameter

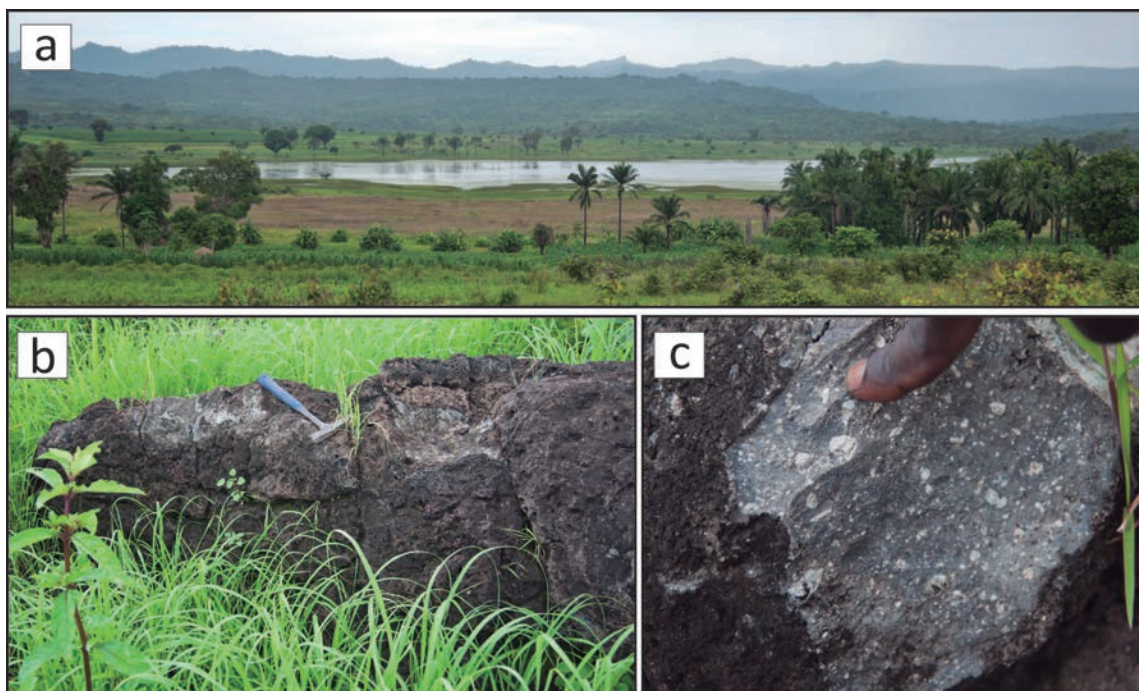


Figure 2 a. General view of the Chiva lagoon area. b. Carbonatitic dyke outcropping on the SE shore of the Chiva lagoon. c. Detail of typical porphyritic texture of the Chiva carbonatitic dykes.

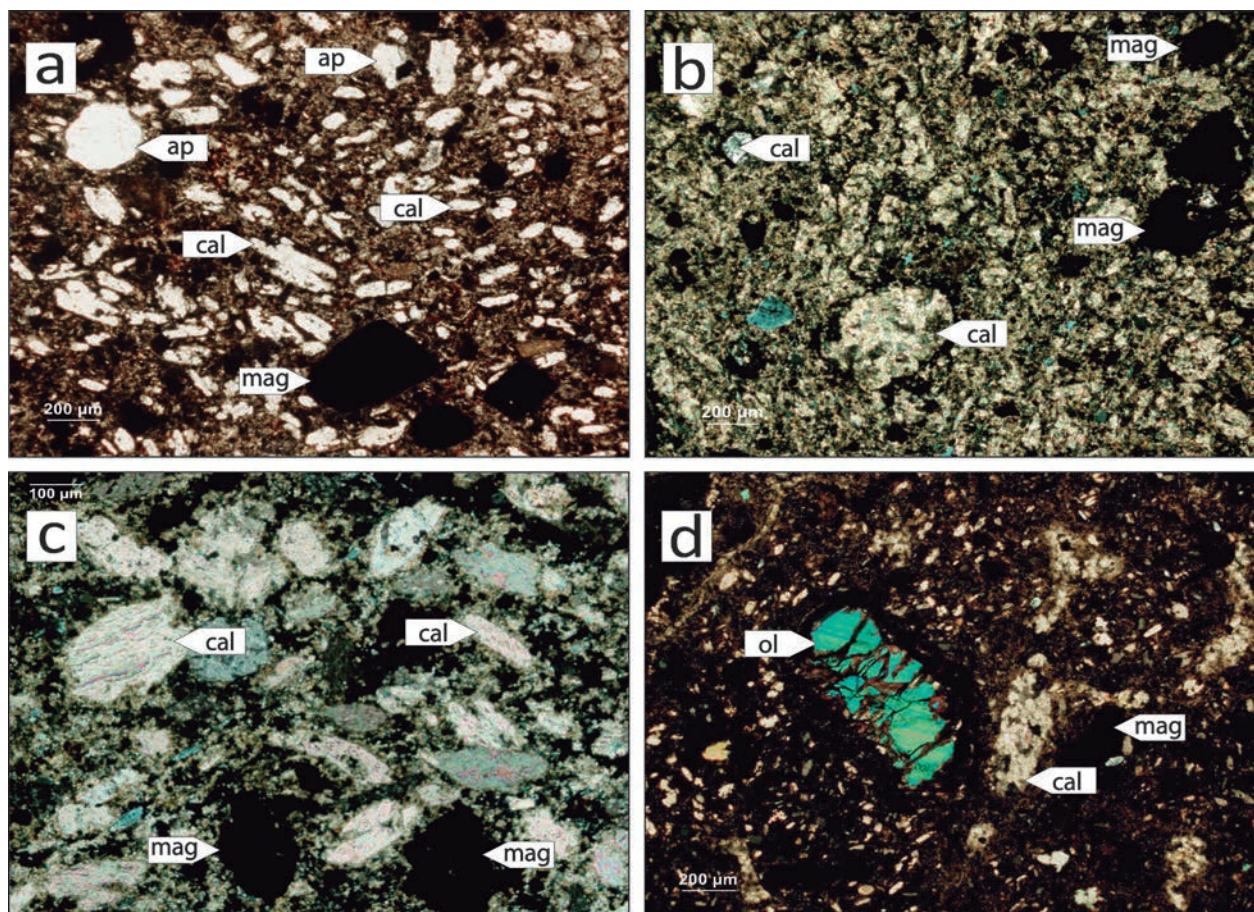


Figure 3 Plane-polarized light photomicrographs of Chiva carbonatitic dykes. a. General view (using parallel Nicols) of the carbonate-rich groundmass mainly formed by calcite (cal) and magnetite (mag). b. Crossed Nicols picture of Apatite (ap), magnetite (mag) and calcite phenocrysts hosted in a carbonatite-rich groundmass. c. Subhedral tabular phenocrysts of calcite together with magnetite (mag). Picture taken using crossed Nicols. d. Pelletal lapilli core comprised of a partially fractured and altered olivine (ol) crystal. Picture taken using crossed Nicols.

representing around 10 % modal (Figures 3a, 3b and 3c).

3.1.2. XENOCRYSTS

Rounded and corroded olivine grains, up to 0.6 mm, partly or completely replaced by serpentine minerals are common and may be interpreted as xenocrysts (Table 1; Figure 3d). In fact, olivine from the Catanda carbonatites has generally been interpreted as xenocrysts by different authors (Peres *et al.*, 1968; Campeny *et al.*, 2015).

3.1.3. GROUNDMASS

The main components of the groundmass are calcite (60 % modal), apatite (10 % modal) and

magnetite (10 % modal), presenting textural patterns similar to those described in the phenocrysts (Table 1; Figure 4a). In addition, variable proportions of magnesium-rich smectites (≈ 20 % modal) are widespread in the groundmass, forming cryptocrystalline masses that fill interstitial spaces (Figure 4). Minor proportions of accessory minerals such as pyrochlore and perovskite are also noticeable. The grain size of these minerals is very small, less than 5 μm (Figures 4a and 4b), and they are scattered in the groundmass or can also appear as inclusions inside magnetite grains. Both minerals develop euhedral to subhedral crystals, and zoning is not evident. Pyrochlore is slightly enriched in Zr while Nb content in perovskite is below the detec-

Table 1. Comparison of the mineral composition between Catanda graben lavas and carbonatitic dykes from the Chiva lagoon area.

Location	Texture	Lithology	Composition	Mineralogy
Huilala - Ungongué	Finely porphyritic	Lava	Aillikite	* Microphenocrysts: fluorapatite (35 %), titaniferous magnetite (35 %), augite (15 %), phlogopite (15 %) * Groundmass: calcite, fluorapatite, magnetite, pyrochlore, baddeleyite, cuspidine, perovskite, periclase, brucite, zirconolite, spurrite, oldhamite * Xenocrysts: olivine * Xenoliths: glimmerite
Jango	Finely porphyritic	Lava	Calciocarbonatite	* Phenocrysts: fluorapatite (60 %), magnetite (25 %), phlogopite (10 %), augite (5 %) * Groundmass: calcite, fluorapatite, magnetite, pyrochlore, baddeleyite, monticellite * Xenocrysts: amphibole
Utihohala	Finely porphyritic	Lava	Calciocarbonatite	* Microphenocrysts: fluorapatite (60 %), magnetite (20 %), phlogopite (15 %), augite (5%) * Groundmass: calcite, fluorapatite, magnetite, pyrochlore, alabandite
Ipunda	Trachytoid	Lava	Altered Natrocarbonatite	* Microphenocrysts: calcite (60 %), fluorapatite (15 %), magnetite (15 %), phlogopite (5 %), augite (5%) * Groundmass: calcite, fluorapatite, magnetite, pyrochlore, baddeleyite * Xenoliths: glimmerite
Chiva lagoon	Finely porphyritic	Dyke	Calciocarbonatite	* Microphenocrysts: calcite (60 %), fluorapatite (30 %), magnetite (10 %) * Groundmass: calcite, fluorapatite, magnetite, pyrochlore, perovskite, kimzeyite * Xenocrysts: olivine

tion limit of the EDS. More rarely, we have also found small subhedral grains of kimzeyite garnet, up to 2 μm in diameter.

3.1.4. VESICULA INFILL

Vesicular porosity is irregular in size and distribution. However, millimetre- to centimetre-sized amygdalae are rather common, and they are normally completely infilled by a supergene sequence of clay minerals, euhedral quartz and late sparitic calcite. Barite is also found in this assemblage as tiny crystals, less than 2 μm in length, commonly included in the calcite groundmass.

4. Geochemistry

Major and trace element compositions of Chiva carbonatitic dykes are compiled in Table 2. It also includes compositional values from different worldwide carbonatite localities for comparison, such as the lavas from the Catanda area (Angola), Fort Portal (Uganda), Oldoinyo Lengai volcano (Tanzania) and the aillikites from Aillik Bay (Canada) (Tappe *et al.*, 2006; Eby *et al.*, 2009; Keller and Zaitsev, 2012; Campeny *et al.*, 2015).

The carbonatitic dykes of the Chiva area have

CaO contents of 35.52 to 34.27 wt%. These values are similar to those reported in the aillikitic and calcio carbonatitic lavas from the Catanda area and the carbonatitic lavas from Fort Portal, but significantly higher in comparison to the aillikites from Aillik Bay and the Oldoinyo Lengai fresh natrocarbonatite lavas. Despite that, altered natrocarbonatite lavas from Catanda and the Oldoinyo Lengai area present higher values of CaO, varying from 45.91 to 46.81 wt% (Table 2). In the case of MgO, Chiva dykes show values from 9.13 to 9.80 wt%, similar to the aillikitic and calciocarbonatitic lavas from Catanda and Fort Portal (Table 2). FeO_t contents range from 8.40 to 8.34 and are also similar from those which are presented by the aillikitic and calciocarbonatitic lavas from Catanda, but they are significantly lower than those reported in the Fort Portal lavas and, especially, in comparison to aillikites from Aillik Bay, that present values from 12.2 to 16.0 wt% FeO_t (Table 2).

SiO_2 totals vary between 12.7 and 13.5 wt% in Chiva dykes (Table 2), with similar contents to those reported in the calciocarbonatite lavas from the Catanda area, as well as in the lavas from Fort Portal (Table 2). In the case of the aillikitic rocks from Catanda and Aillik Bay, SiO_2 values are significantly higher than in Chiva, varying from

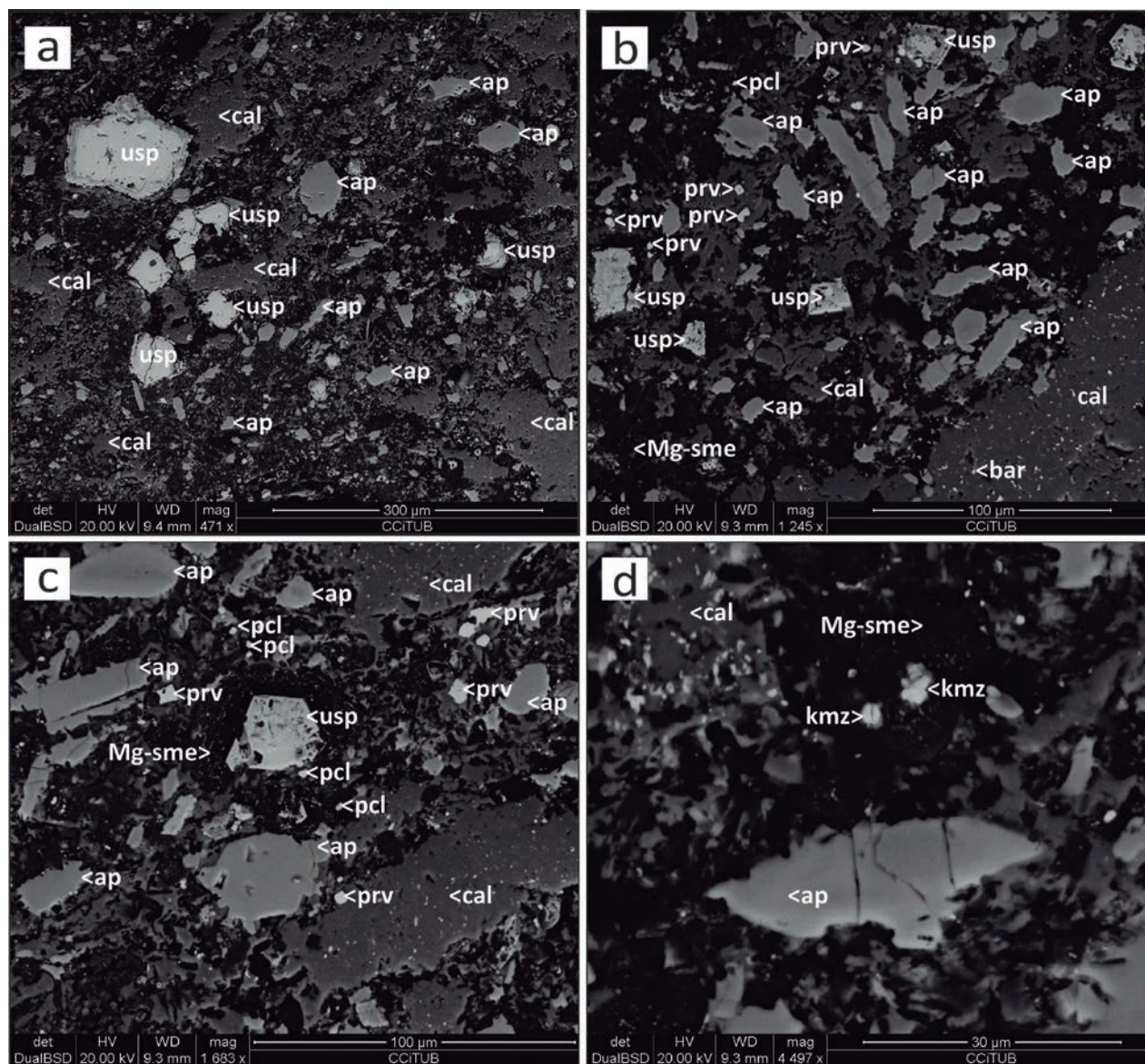


Figure 4 Backscattered scanning electron microscope (SEM) images of Chiva carbonatitic dyke samples. a. General view of typical porphyritic texture with magnetite (mgt), calcite (cal) and apatite (ap) microphenocrysts. b. Apatite (ap) prismatic crystals along side small grains of perovskite (prv) and magnetite (mgt), present in a calcite-rich (cal) groundmass, with abundant small inclusions of barite (bar) and magnesium-rich smectite (Mg-sme) in interstitial spaces. c. Detail of the groundmass with pyrochlore (pcl), perovskite (prv) and microphenocrysts of apatite (ap), magnetite (mgt) and calcite (cal) and interstitial magnesium-rich smectite (Mg-sme). d. Subhedral grains of kimzeyite (kmz) hosted in magnesium-rich smectite (Mg-sme) and associated to a calcite groundmass (cal) and a subhedral prismatic grain of apatite (ap).

14.29 to 23.33 and from 18.2 to 26.8, respectively. We also carried out a general comparison between SiO_2 contents and the proportion of different elements found in Chiva dykes, that also include data from different carbonatitic rocks worldwide (Figure 5). In these diagrams it is possible to distinguish that the main compositional features of

Chiva dykes are especially similar to those reported in the calciocarbonatitic lavas of the Catanda area (Figure 5).

Chiva dykes have a significant enrichment in REE, presenting LREE contents from 100 to 1000 times higher to those reported in C1 chondrites and 10 to 100 times enriched in the case of HREE (Figure

Table 2. Major element composition of Chiva carbonatitic dykes compared to the composition of Catanda graben lavas. Values from Campeny *et al.* (2015). AL (aillikite lavas), NL (natrocarbonatite lavas), CCL (calcicarbonatitelavas), CD (carbonatitic dykes).

Sample wt%	Chiva lagoon		Catanda graben (Angola)										Fort Portal (Uganda)		Aillik Bay (Canada)										Oldoinyo Lengai (Tanzania)			
	CD		AL		ANL		CCL		CCL		CCL		CCL		Aillikites UML		NL		ANL		NL		ANL		NL		ANL	
	CM=14	CM=15	AC=19	AC=20	AC=21	AC=24	AC=25	AC=35	AC=36	CAI=79	AC=39	AC=41	AC=17	UG=6	UG=15	L60	ST109	ST164	ST198A	ST12001	ST1225	ST1250A	OL.234	OL.321.4	OL.329	OL.328.3	OL.239H	OL.328
SiO ₂	12.66	13.45	14.64	14.29	14.39	14.58	23.33	14.81	19.60	6.25	12.78	11.99	9.76	11.03	18.45	23.90	18.20	24.10	26.80	26.10	26.40	22.50	0.34	0.26	0.69	0.4	1.67	0.97
TiO ₂	0.80	0.75	1.45	1.41	1.45	1.49	2.35	1.51	2.21	0.38	0.68	0.76	0.24	1.62	2.40	3.80	3.60	3.30	3.80	4.80	3.30	2.50	0.01	0.01	0.04	0.03	0.04	0.03
Al ₂ O ₃	2.70	2.84	3.11	3.22	3.16	3.24	5.92	3.30	4.75	1.73	2.43	2.58	2.33	2.69	4.90	3.00	2.80	3.30	3.50	3.20	2.50	3.10	0.04	0.05	0.05	0.02	0.04	0.02
FeO _i	8.40	8.34	8.94	8.61	8.71	9.24	11.28	9.04	10.58	4.15	8.12	8.09	6.32	11.66	10.83	13.50	16.00	13.20	13.70	15.00	13.20	12.20	0.65	0.58	1.88	1.2	0.19	0.95
MnO	0.45	0.47	0.55	0.53	0.52	0.58	0.44	0.56	0.48	0.32	0.31	0.28	0.63	0.45	0.35	0.30	0.30	0.20	0.20	0.20	0.20	0.46	0.44	1.25	1.11	0.8	1.34	
MgO	9.13	9.80	8.59	8.61	8.17	9.09	9.40	8.97	9.16	1.36	9.73	9.45	9.18	8.24	8.95	19.10	15.90	18.60	17.30	16.50	20.90	20.40	0.42	0.44	0.83	0.52	1.2	1.1
CaO	35.52	34.27	38.08	37.26	34.57	37.27	29.36	33.78	31.34	46.53	36.53	37.32	36.52	34.86	28.31	17.60	19.70	18.10	14.30	16.10	16.40	15.60	16.4	15.83	46.07	45.91	46.81	45.62
Na ₂ O	0.01	0.02	0.24	0.22	0.27	0.17	0.29	0.23	0.20	0.25	0.04	0.08	0.07	0.37	0.45	0.30	0.20	0.20	0.70	0.70	0.20	0.20	33.68	33.08	1.73	2.91	2.36	2.44
K ₂ O	0.01	0.01	0.13	0.15	0.10	0.14	0.57	0.14	0.14	0.22	0.03	0.02	0.11	0.26	0.63	1.60	1.80	2.01	1.70	1.50	1.40	2.10	7.04	7.43	0.19	0.2	0.16	0.25
suma alk	0.02	0.03	0.37	0.37	0.37	0.31	0.86	0.37	0.34	0.47	0.07	0.10	0.18	0.63	1.08	1.90	2.00	2.21	2.40	2.20	1.60	2.30	40.72	40.51	1.92	3.11	2.52	2.69
P ₂ O ₅	3.99	4.14	4.23	4.12	4.08	4.28	3.06	4.38	3.58	2.28	2.79	2.73	4.71	3.17	2.34	2.10	3.10	3.00	1.70	2.30	2.00	1.50	1.13	0.9	2.68	1.51	0.66	1.71
LOI	24.92	24.54	18.5	20.2	23.1	18.4	12.9	21.9	16.8	35.8	25.7	25.7	28.6	22.73	20.17	14.00	17.30	12.80	15.00	12.50	12.30	18.50	32.63	32.97	34.02	34.61	30.52	30.09
Sum	98.58	98.63	99.53	99.57	99.54	99.50	99.54	99.50	99.54	99.83	99.68	99.70	99.57	97.08	97.78	98.90	98.70	98.70	98.60	98.80	98.90	98.70	100.2	99.66	97.62	98.94	98.37	98.44
ppm																												
V	173	165	239	215	220	225	262	234	257	76	208	191	126	291	321	205	267	204	211	242	194	197	130	155	246	218	174	294
Co	16	15	25.3	23.4	38.3	26.1	31.8	20.6	28.3	10.5	22.4	22.9	9.9	30.20	33.20	67	62	66	73	75	78	63	nd	nd	nd	nd	nd	nd
Cu	20	20	29.67	32.03	33.88	34.55	58.65	30.58	53.45	8.41	47.49	58.67	4.74	nd	nd	nd	nd	nd	nd	nd	nd	nd	7	6	8	6	8	7
Ni	<20	<20	7	7.7	9.4	17.4	7.6	11.6	20	20	1.9	3.0	1	28	73	441	211	483	509	493	719	487	6	6	9	8	13	10
Zn	110	120	120.1	118.3	125.2	131.8	110.5	118.1	114.7	76.1	64.5	66.5	122	179	122	nd	nd	nd	nd	nd	nd	nd	137	126	373	318	117	295
Rb	<2	<2	13.4	13.1	9.8	16.9	31.3	4.9	8.8	7.6	0.8	0.8	5.5	nd	10	55	61	66	49	38	29	75	nd	nd	nd	nd	nd	nd
Cs	<0.5	<0.5	0.8	0.8	0.6	1	0.5	0.2	0.5	0.1	0.4	0.2	0.2	0.80	1.10	1.56	1.53	1.94	1.23	2.29	nd	6.13	nd	nd	nd	nd	nd	nd
Sr	3650	3520	5446	5186	5147	5401	2427	3849	3249	2826	2743	2903	5491	5830	4387	2685	1678	2784	1318	2511	2868	1312	nd	nd	nd	nd	nd	nd
Ba	1620	1540	1413	1367	2010	1430	1357	1433	1147	888	936	1552	2416	2717	2441	1860	2777	1540	1028	714	1018	1366	nd	nd	nd	nd	nd	nd
La	623	652	756.4	779.1	743	811	674.6	797.7	762.2	405.4	260	252.5	718.5	454.50	319	367	390	235	182	272	335	242	571	623	1651	1604	1382	1647
Ce	1070	1120	1297.2	1319.5	1288.8	1388.5	1109.5	1390.1	1255.8	693.8	491.9	507.7	1211.6	896.50	606	744	819	469	391	572	700	476	768	770	2012	1874	1449	1924
Pr	104	108	133.49	133.30	127.99	141.67	108.71	138.01	124.47	76.80	57.39	58.34	122.70	nd	nd	86.30	95.80	55.10	45.90	65.50	82.30	52.40	nd	nd	nd	nd	nd	nd
Nd	354	365	451.5	447.2	434.3	475.1	358.4	472.2	416.3	254.6	217.7	220.1	406.4	354.00	257.60	322	364	215	181	251	319	188	153	135	424	341	170	311
Sm	47.80	48	55.67	56.34	53.89	58.91	42.86	57.24	49.87	33.50	30.26	30.38	52.81	46.00	33.50	49.80	54.60	36.80	31.30	45.40	47.20	28.10	nd	nd	nd	nd	nd	nd
Eu	12.40	12.70	14.56	14.74	13.84	15.30	11.15	15.04	13	9.10	7.68	8.03	13.70	13.10	8.57	12.54	13.46	10.11	8.45	12.43	12.57	7.14	nd	nd	nd	nd	nd	nd
Gd	28.40	28.40	41.75	41.35	38.90	43.97	31.68	41.99	37.31	19.20	19.44	19.84	39.23	30.20	22.00	30.82	33.37	26.47	21.52	30.98	30.68	17.37	nd	nd	nd	nd	nd	nd
Tb	3.10	3.20	4.09	4.17	3.96	4.30	3.11	4.13	3.62	2.70	1.86	2	3.78	3.48	2.41	3.14	3.30	3.02	2.43	3.64	3.38	1.91	nd	nd	nd	nd	nd	nd
Dy	14.60	14.90	18.92	19.33	18.71	19.69	14.83	18.93	17.14	11.50	8.35	8.55	18.79	nd	nd	12.29	12.82	12.51	9.80	15.13	13.23	7.70	nd	nd	nd	nd	nd	nd
Ho	2.30	2.30	2.74	2.75	2.82	2.97	2.14	2.85	2.58	1.90	1.28	1.20	2.50	nd	nd	1.59	1.65	1.74	1.30	2.10	1.83	1.11	nd	nd	nd	nd	nd	nd
Er	5.20	5.20	6.62	7.08	6.35	6.88	4.99	6.75	5.38	4.20	2.56	2.71	6.20	nd	nd	3.44	3.63	3.98	2.95	4.84	4.00	2.44	nd	nd	nd	nd	nd	nd
Tm	0.65	0.67	0.85	0.86	0.79	0.88	0.66	0.90	0.74	0.70	0.34	0.37	0.82	0.61	0.45	0.38	0.39	0.44	0.34	0.54	0.42	0.29	nd	nd	nd	nd	nd	nd
Yb	3.70	3.70	5.44	5.31	4.29	5.86	3.94	5.18	4.10	3.90	2.07	2.07	4.89	3.76	2.72	1.95	2.05	2.26	1.69	2.77	2.19	1.47	nd	nd	nd	nd	nd	nd
Lu	0.52	0.52	0.78	0.74	0.62	0.75	0.55	0.75	0.62	0.60	0.30	0.30	0.63	0.51	0.37	0.22	0.23	0.27	0.20	0.30	0.25	0.17	nd	nd	nd	nd	nd	nd
Y	61	61	76.6	76.9	72.8	85.4	59.9	78.2	70.5	49	33.9	34	71.7	53	38	41	43	46	35	55	43	27	nd	nd	nd	nd	nd	nd
Pb	9	8	7.09	6.89	7	5.46	6.17	7.33	5.82	15.09	10.54	7.38	12.76	nd	nd	13.00	11.90	8.50	5.40	7.50	9.60	10.80	91	101	476	170	119	371
Nb	418	190	590	577	588	607	441	624	504	281	411	365	706	509	381	166	203	153	131	177	159	184	64	64	122	58	104	59
Zr	833	288	725	717	709	700	599	731	626	401	592	555	463	667	457	466	591	608	350	359	306	303	<10	12	80	43	86	91
Hf	11	3.60	7.3	8.7	7.8	8.7	8.2	7.8	8.4	5.7	9.4	8.9	5.7	8.90	7.00	11.45	13.95	14.20	7.90	8.46	7.30	8.62	nd	nd	nd	nd	nd	nd
Ta	10.90	5	20																									

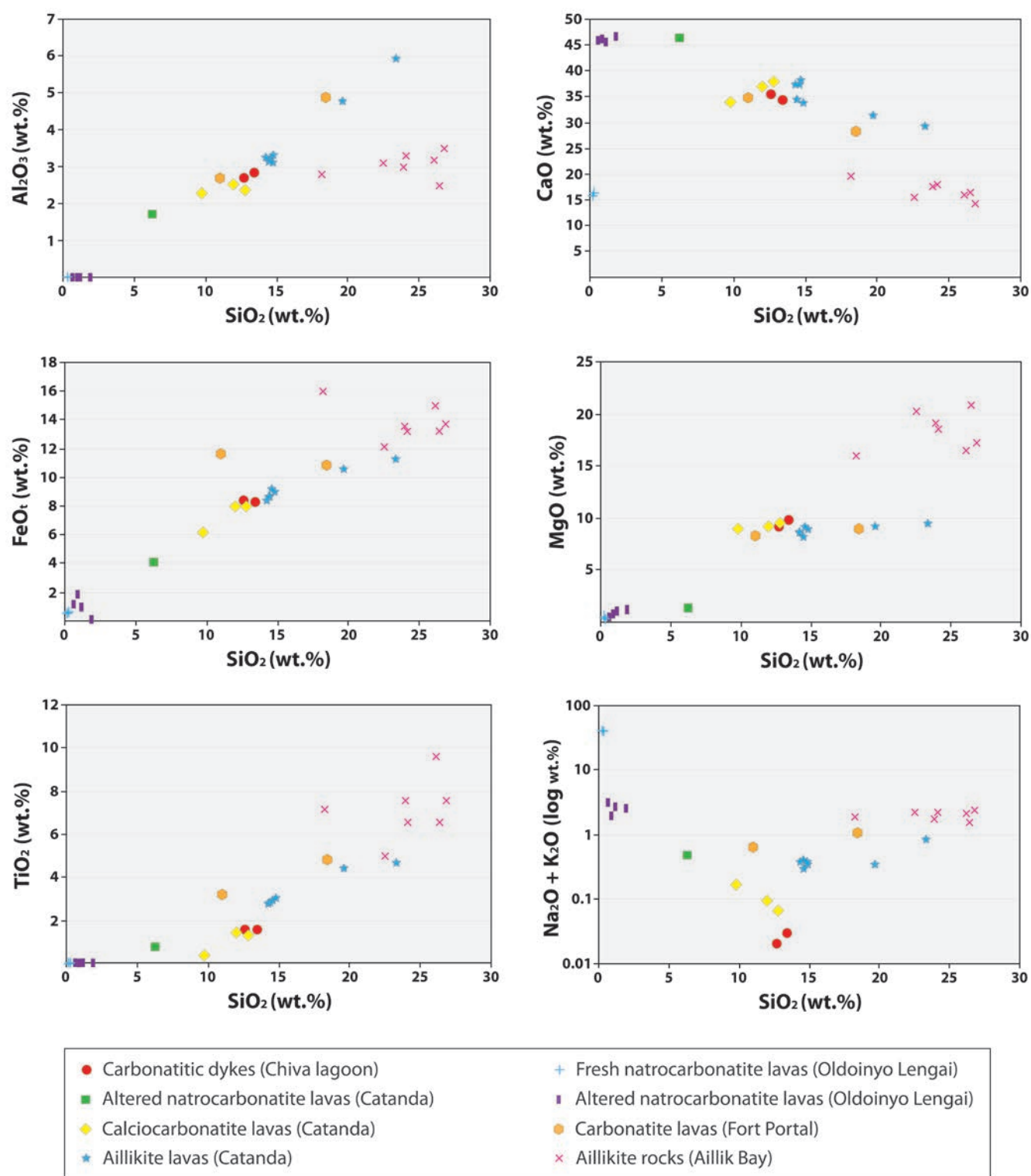


Figure 5 Compositional relation between SiO_2 and different major elements of Chiva carbonatitic dykes. Values are compared to those reported for Catanda lavas and other carbonatitic rocks worldwide. Values obtained from: Tappe *et al.* (2006); Eby *et al.* (2009); Keller and Zaitsev (2012); Campeny *et al.* (2015).

6a). General values are similar to those reported in other carbonatitic rocks (Table 2) but they are not significant enough to be evaluated as an economic resource. REE patterns of the studied carbonatitic dykes from Chiva show clear negative slopes with a significant enrichment in LREE relative to MREE and HREE. These are similar patterns to those presented by the Catanda graben lavas (Figure 6a).

Multi-element diagrams also show very similar patterns between Chiva dykes and Catanda graben lavas. The carbonatitic dykes are systematically enriched in all plotted elements (including LILE, HFSE and REE) to PM (Figure 6b).

5. Discussion

The dykes cropping out in the SE border of the Chiva lagoon indicate the development of carbonatitic magmatism events in this area. Furthermore, considering the proximity of Chiva to the Catanda graben, where volcanic carbonatitic activity is well known (Silva Pereira, 1973; Campeny *et al.*, 2014; Campeny *et al.*, 2015), both carbonatitic magmatic areas are probably related.

Mineralogical features of the Chiva carbonatitic dykes are very similar to those reported in the carbonatitic lavas from Catanda (Campeny *et al.*,

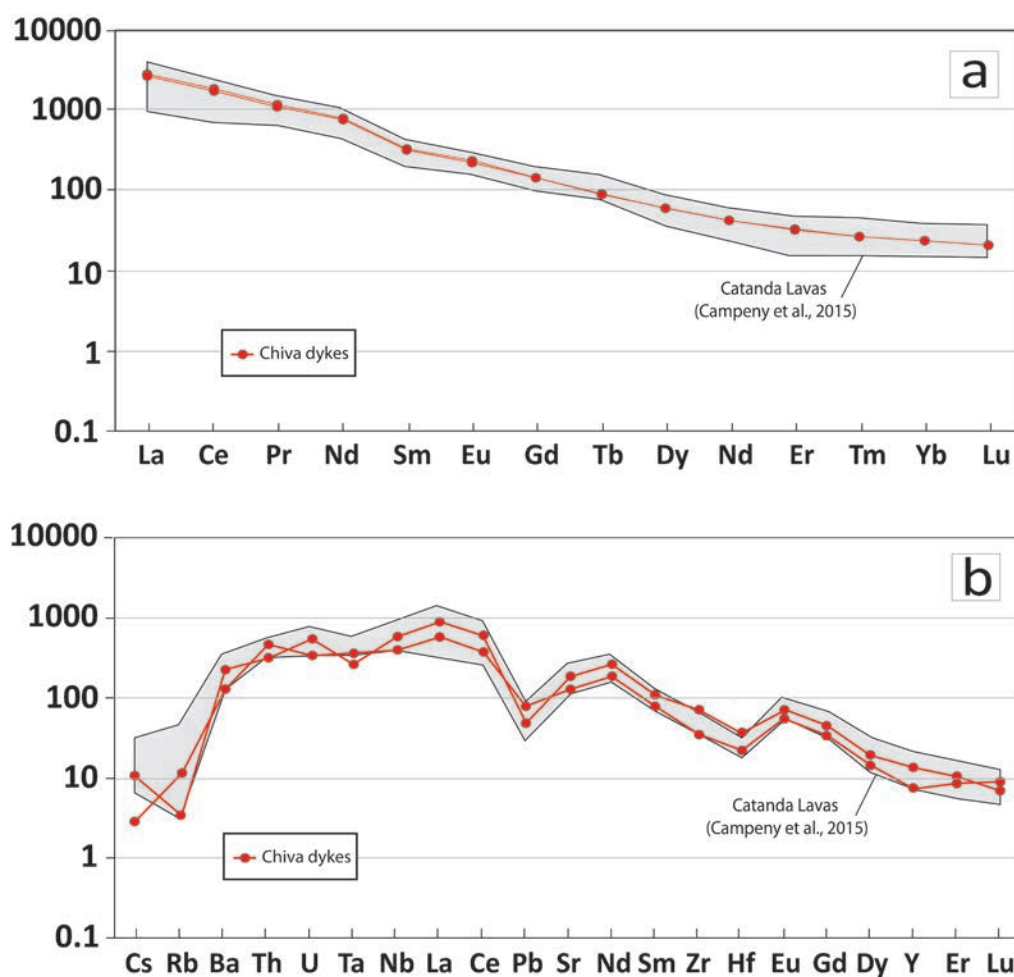


Figure 6 a. REE plot of Chiva carbonatitic dykes normalized to chondrite C1. b. Multi-elemental trace element composition of Chiva dykes normalized to primitive mantle (PM). Normalization values used in both diagrams are from Sun and McDonough (1989). Chiva carbonatitic dyke patterns are compared to Catanda lavas in both diagrams. Values from Campeny *et al.* (2015).

2014; Campeny *et al.*, 2015). Minerals such as calcite, apatite and magnetite are the main phenocryst phases in both localities, while the ground-mass mineralogy is also equivalent. The latter is comprised of sparitic calcite and accessory minerals such as pyrochlore or perovskite; both reported in Chiva dykes and also in Catanda graben lavas (Table 1).

The comparison of major and trace element compositions between Chiva dykes and Catanda lavas also indicates a clear genetic relation. Multi-elemental and REE diagrams of Chiva carbonatitic dykes present equivalent patterns to those reported for Catanda lavas, suggesting that all these carbonatitic products were generated from the same parental melt. Major element composition is very heterogeneous in the case of Catanda lavas but the Chiva carbonatitic dykes are compositionally equivalent to the calciocarbonatitic lavas from Jango and Utihohala monogenetic cones. These volcanic edifices are also located in the most external part of the Catanda graben region (Campeny *et al.*, 2014; Campeny *et al.*, 2015). In addition, we can also conclude that carbonatitic dykes from Chiva present compositional and mineralogical features akin to other carbonatitic volcanic rocks worldwide such as Fort Portal (Uganda) (Eby *et al.*, 2009).

Hence, considering the mineralogical and compositional features of the Chiva dykes, we conclude that these carbonatitic outcrops are related to the volcanic activity in Catanda despite the fact that Chiva is located outside of the Catanda graben. Therefore, we assume that carbonatitic volcanic activity in this area was also developed outside the limits of the Catanda graben and the discovery of new carbonatitic volcanic outcrops in this region should not be ruled out.

The radial distribution of the Chiva dykes, oriented from the lagoon centre to its border, and the morphology of the Chiva area, which is defined by a circular depressed ring, suggest that an eruptive centre was once located in the region. Hypabyssal intrusions related to volcanic edifices are common and broadly described in several volcanic areas of

different composition worldwide (Gautneb and Guddmundsson, 1992; Geshi, 2005; Pasquarè and Tibaldi, 2007). Dykes are generated by magma intrusions during volcanic events and represent the magma path towards the surface (Corazzato *et al.*, 2008).

Despite the fact that the morphology of the Chiva area is broadly eroded and covered by sediments and vegetation, the topographic position of the area, which is below the average altitude of surrounding terrains, suggests that the Chiva volcanic edifice has a maar morphology. Maars are volcanic constructions related to explosive activity and are commonly associated to recent carbonatitic volcanism such as in the Eifel region (Germany; Keller, 1981), San Venzano (Italy; Stoppa, 1996), the East African Rift (Deans and Roberts, 1984; Woolley and Church, 2005) or even the volcanic edifices that stand out inside the Catanda graben (Campeny *et al.*, 2014).

6. Conclusions

The main conclusions of our work can be summarized as follows.

- Dykes of carbonatitic composition outcrop in the SE border of the Chiva lagoon area (Angola). They have been interpreted as hypabyssal intrusive sheets related to the occurrence of recent volcanic activity, producing a single eruptive centre with typical maar morphology.
- Mineralogical and compositional features of the Chiva dykes are similar to those reported in the carbonatitic lavas of the Catanda graben area, suggesting that both volcanic events are genetically related. Moreover, the Chiva dykes also present compositional similarities to the volcanic silica-rich carbonatites of the Eastern African rift, such as Fort Portal (Uganda).
- The genetic relation between Chiva and the Catanda graben suggests that the recent carbonatitic volcanism of the Lucapa domain has a broader extension than previously sup-

posed. The discovery of new carbonatitic volcanic localities outside the main Catanda graben should not be discarded in this region of Angola.

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8. References

- Alberti, A., Castorina, F., Censi, P., Comin-Chiaramoni, P., Gomes, C.B., 1999, Geochemical characteristics of Cretaceous carbonatites from Angola: *Journal of African Earth Sciences*, 29, 735–759.
- Bell, K., Simonetti, A., 2010, Source of parental melts to carbonatites—critical isotopic constraints: *Mineralogy and Petrology*, 98, 77–89.
- Calvo, A., Melgarejo, J.C., Bambi, A.C.J.M., Gonçalves, A.O., Alfonso, P., 2011a, Nb and REE at the Bailundo carbonatite, Angola, *in* Barra, F. (ed.), *Let's talk ore deposits: Antofagasta*, Universidad Católica del Norte, 675–677.
- Calvo, A., Melgarejo, J.C., Alfonso, P., Bambi, A.C.J.M., Gonçalves, A.O., 2011b, Nb and REE mobilization at the Longonjo carbonatite, Angola, *in* Barra, F. (ed.), *Let's talk ore deposits: Antofagasta*, Universidad Católica del Norte, 681–683.
- Campany, M., 2016, *Caracterización del vulcanismo carbonatítico de Catanda (Angola)*: Barcelona, University of Barcelona, PhD thesis, 151 p.
- Campany, M., Mangas, J., Melgarejo, J.C., Bambi, A., Alfonso, P., Manuel, J., 2014, The Catanda extrusive carbonatites (Kwanza Sul, Angola): an example of explosive carbonatitic volcanism: *Bulletin of Volcanology*, 76, 818–832.
- Campany, M., Kamenetsky, V.S., Melgarejo, J.C., Mangas, J., Manuel, J., Alfonso, P., Kamenetsky, M.B., Bambi, A., Gonçalves, A.O., 2015, Carbonatitic lavas in Catanda (Kwanza Sul, Angola): mineralogical and geochemical constraints on the parental melt: *Lithos*, 232, 1–11.
- Chakhmouradian, A.R., 2006, High-field-strength elements in carbonatitic rocks: Geochemistry, crystal chemistry and significance for constraining the sources of carbonatites: *Chemical Geology*, 235, 138–160.
- Chakhmouradian, A.R., Mumin, A.H., Demény, A., Elliott, B., 2008, Postorogenic carbonatites at Eden Lake, Trans-Hudson Orogen (northern Manitoba, Canada): geological setting, mineralogy and geochemistry: *Lithos*, 103, 503–526.
- Chakhmouradian, A.R., Wall, F., 2012, Rare earth elements: minerals, mines, magnets (and more): *Elements*, 8 (5), 333–340.
- Corazzato, C., Francalanci, L., Menna, M., Petrone, C.M., Renzulli, A., Tibaldi, A., Vezzoli, L., 2008, What controls sheet intrusion in volcanoes? Structure and petrology of the Stromboli sheet complex, Italy: *Journal of Volcanology and Geothermal Research*, 173, 26–54.
- Dawson, J.B., 1962, Sodium carbonate lavas from Oldoinyo Lengai, Tanganyika: *Nature*, 195, 1075–1076.
- Dawson, J.B., Pinkerton, H., Norton, G.E., Pyle, D.M., Browning, P., Jackson, D., Fallick, A.E., 1995, *Petrology and geochemistry*

- of Oldoinyo Lengai lavas extruded in November 1988: magma source, ascent and crystallization, *in* Bell, K., Keller, J. (eds.), Carbonatite Volcanism. Oldoinyo Lengai and the petrogenesis of natrocarbonatite IAVCEI Proceedings in Volcanology 4, 47–69.
- Deans, T., Roberts, B., 1984, Carbonatite tuffs and lava clasts of the Tinderet foothills, western Kenya: a study of calcified natrocarbonatites: *Journal of the Geological Society of London* 141, 563–580.
- Eby, G.N., Lloyd, F.E., Woolley, A.R., 2009, Geochemistry and petrogenesis of the Fort Portal, Uganda, extrusive carbonatite: *Lithos* 113, 785–800.
- Gautneb, H., Gudmundsson, A., 1992, Effect of local and regional on sheet emplacement in West Iceland: *Journal of Volcanology and Geothermal Research*, 51, 339–356.
- Geshi, N., 2005, Structural development of dike swarms controlled by the change of magma supply rate: the cone sheets and parallel dike swarms of the Miocene Otoge igneous complex, Central Japan: *Journal of Volcanology and Geothermal Research*, 141, 267–281.
- Jelsma, H., Barnett, W., Richards, S., Lister, G., 2009, Tectonic setting of kimberlites: *Lithos*, 112S, 155–165.
- Keller, J., 1981, Carbonatitic volcanism in the Kaiserstuhl alkaline complex: evidence for highly fluid carbonatitic melts at the Earth's surface: *Journal of Volcanology and Geothermal Research*, 9, 423–431.
- Keller, J., Zaitsev, A., 2012, Geochemistry and petrogenetic significance of natrocarbonatites at Oldoinyo Lengai, Tanzania: Composition of lavas from 1988 to 2007: *Lithos*, 148, 45–53.
- Le Bas, M.J., Rex, D.C., Stillman, C.J., 1986, The early magmatic chronology of Fuerteventura: *Geological Magazine*, 123, 287–298.
- Le Maitre, R.W., 2002, *Igneous Rocks: a Classification and Glossary of Terms*. Cambridge, Cambridge University Press, 236 p.
- Mangas, J., Pérez Torrado, F.J., Martín Izard, A., Reguilón Bragado, R.M., 1996, Minerales de tierras raras en las carbonatitas del Complejo Basal de Fuerteventura (Islas Canarias): *Geogaceta*, 20 (7), 1511–1513.
- Mitchell, R.H., 2005, Carbonatites and carbonatites and carbonatites: *The Canadian Mineralogist*, 43, 2049–2068.
- Moulin, M., Aslanian, D., Unternehr, P., 2010, A new starting point for the South and Equatorial Atlantic Ocean: *Earth-Science Reviews*, 98, 1–37.
- Norrish, K., Hutton, J.T., 1969, An accurate X-ray spectrographic method for the analysis of a wide range of geological samples: *Geochimica et Cosmochimica Acta* 33, 431–453.
- Pasquarè, F., Tibaldi, A., 2007, Structure of a sheet-laccolith system revealing the interplay between tectonic and magma stresses at Stardalur Volcano, Iceland: *Journal of Volcanology and Geothermal Research*, 161, 131–150.
- Peres, A.M., Gomes, C.F., Cardoso-Simões, M.V., 1968, As variedades semi-preciosas de olivina da Catanda-Angola: *Boletim dos Serviços de Geologia e Minas de Angola*, 18, 5–15.
- Robles-Cruz, S.E., Escayola, M., Jackson, S., Galí, S., Pervov, V., Watangua, M., Gonçalves, A., Melgarejo, J.C., 2012, U-Pb SHRIMP geochronology of zircon from the Catoca kimberlite, Angola: Implications for diamond exploration: *Chemical Geology*, 310–311, 137–147.
- Silva, M.V.S., Pereira, E., 1973, Estrutura Vulcânica-Carbonatítica da Catanda (Angola): *Boletim dos Serviços de Geologia e Minas de Angola*, 24, 5–14.
- Stoppa, F., 1996, The San Venanzo maar and tuff ring, Umbria, Italy: eruptive behaviour of a carbonate-melilitite volcano: *Bulletin of Volcanology*, 57, 563–577.

- Sun, S.S., McDonough, W.F., 1989, Chemical and isotopic systematics of oceanic basalts; implications for mantle composition and processes, *in* Saunders, A.D., Norry, M.J. (eds.), *Magmatism in the ocean basins*: London, Geological Society of London, 42, 313–345.
- Sykes, L., 1978, Intraplate seismicity, reactivation of pre-existing zones of weakness, alkaline magmatism, and other tectonism postdating continental fragmentation: *Reviews of Geophysics*, 16, 621–688.
- Tappe, S., Foley, S.F., Jenner, G.A., Heaman, L.M., Kjarsgaard, B.A., Romer, R.L., Stracke, A., Joyce, N., Hoefs, J., 2006, Genesis of ultramafic lamprophyres and carbonatites at Aillik Bay, Labrador: a consequence of incipient lithospheric thinning beneath the North Atlantic Craton: *Journal of Petrology*, 47(7), 1261–1315.
- Tilton, G.R., Bryce, J.G., Mateen, A., 1998, Pb–Sr–Nd isotope data from 30 and 300 Ma collision zone carbonatites in northwest Pakistan: *Journal of Petrology*, 39, 1855–1874.
- Torquato, J.R., Amaral, G., 1974, Algumas idades K/Ar do magmatismo mesozóico de Angola e sua correlação com o correspondente do sul do Brasil: *Boletim do Instituto de Investigação Científica de Angola*, 10, 31–38.
- Torró, L., Villanova, C., Castillo, M., Campeny, M., Gonçalves, A.O., Melgarejo, J.C., 2012, Niobium and rare earth minerals from the Virulundo carbonatite, Namibe, Angola: *Mineralogical Magazine*, 76, 393–409.
- Wallace, M.E., Green, D.H., 1988, An experimental determination of primary carbonatite magma composition: *Nature*, 335, 343–346.
- Widom, E., Hoernle, K.A., Shirey, S.B., Schminke, H.U., 1999, Os isotope systematics in the Canary Islands and Madeira: Lithospheric contamination and mantle plume signatures: *Journal of Petrology*, 40 (2), 279–296.
- Woolley, A.R., Church, A.A., 2005, Extrusive carbonatites: a brief review: *Lithos*, 85, 1–14.
- Woolley, A.R., Kjarsgaard, B.A., 2008, Carbonatites of the world: map and database: *Mineralogical Magazine*, 71, 718.
- Wyllie, P.J., Jones, A.P., Dent, J., 1996, Rare earth elements in carbonate-rich melts from mantle to crust, *in* Jones, A.P., Wall, F., Williams, C.T. (eds.), *Rare earth minerals: Chemistry, origin and ore deposits*: London, Chapman & Hall, 77–102.